

# **Modeling of geochemical processes**

**Equilibrium modeling**

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# Modeling of geochemical processes

## Equilibrium modeling

A set of functions,  $f$ , are defined that must be solved simultaneously to determine equilibrium for a given set of conditions.

Many of these functions are derived from

- *mole-balance equations* for *each element* or *element valence state*, *exchange site*, and *surface site*
- *mass-action equations* for pure phases and solid solutions.
- *aqueous charge balance*

Additional functions are derived for

- *alkalinity*
- *activity of water*
- *gas-phase equilibria*
- *ionic strength*
- *surface-complexation equilibria.*

Each function is reduced to contain a minimum number of variables, such that *the number of functions equals the number of variables.*

The program uses a modified *Newton-Raphson method* to solve the simultaneous nonlinear equations.

# Modeling of geochemical processes

## Equilibrium modeling

In the following equations, lack of a subscript or the subscript “(aq)” will refer to entities in the aqueous phase, “(e)” refers to exchangers, “(g)” refers to gases, “(s)” refers to surfaces, “(ss)” refers to solid solutions, and “(p)” refers to phases.

The unknowns for each aqueous species  $i$  are the activity,  $a_i$ , activity coefficient,  $\gamma_i$ , molality,  $m_i$ , and moles in solution,  $n_i$ .

The activity of water,  $a_{H_2O}$ , the ionic strength,  $\mu$ , and the mass of solvent water in an aqueous solution,  $W_{aq}$ .

The following relationships apply

$$a_i = \gamma_i m_i$$

$$n_i = m_i W_{aq}$$

# Modeling of geochemical processes

## Equilibrium modeling

### Components of base:

- $A_{\text{aq}}$  – water
- $A_i$  – dissolved species, base components
- $A_p$  – mineral phases in equilibrium
- $A_g$  – gases of known fugacity

### *Another species:*

- $A_j$  – dissolved species (secondary)
- $A_{\text{pj}}$  – secondary minerals, (they can be virtual)

Formation of mineral,  $A_j$ , from the base components  $\mathbf{B} = (A_{\text{aq}}, A_i, A_p, A_g)$

$$v_p A_p + v_{\text{aq}} A_{\text{aq}} + v_g A_g + v_i A_i + v_j A_j = 0$$

$$\sum_u v_u A_u = 0$$

$$v_j A_j = v_p A_p + v_{\text{aq}} A_{\text{aq}} + v_g A_g + v_i A_i$$

$$A_j = \frac{v_p}{v_j} A_p + \frac{v_{\text{aq}}}{v_j} A_{\text{aq}} + \frac{v_g}{v_j} A_g + \frac{v_i}{v_j} A_i$$

$$A_j = v_{\text{pj}} A_p + v_{\text{aq},j} A_{\text{aq}} + v_{\text{gj}} A_g + v_{\text{ij}} A_i$$

# Modeling of geochemical processes

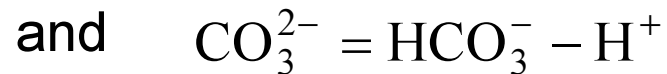
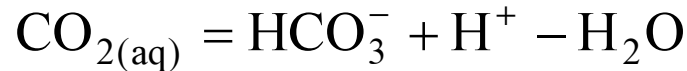
## Equilibrium modeling

### Example

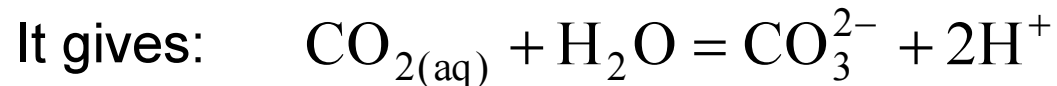
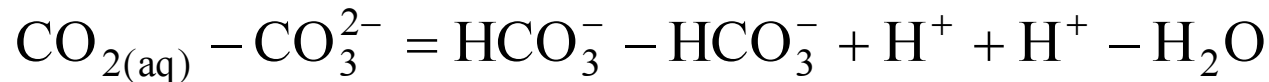
Let us have the base  $\mathbf{B} = (\text{H}_2\text{O}, \text{H}^+, \text{HCO}_3^-)$ ,

Secondary species are  $\text{CO}_{2(\text{aq})}$  and  $\text{CO}_3^{2-}$ .

Build the secondary species from the base components:



Reaction between the secondary species can be expressed by linear combination of both equations (e.g. by subtracting of second equation from the first one):



# Modeling of geochemical processes

## Equilibrium modeling

### Mass-action equations

In general, mass-action equations can be written as

$$K_i = a_i \prod_m^{M_{aq}} a_m^{v_{m,i}}$$

where  $K_i$  is a temperature-dependent equilibrium constant,  $v_{m,i}$  is the stoichiometric coefficient of master species  $m$  in species  $i$  and  $M_{aq}$  is the total number of aqueous master species. The values of  $v_{m,i}$  may be positive or negative.

The total moles of an aqueous species  $i$  can be derived from the mass-action expression:

$$n_i = m_i W_{aq} = K_i W_{aq} \frac{\prod_m^{M_{aq}} a_m^{-v_{m,i}}}{\gamma_i}$$

# Modeling of geochemical processes

## Equilibrium modeling

Activity coefficients of aqueous species are defined with the Davies equation:

$$\log \gamma_i = -A z_i^2 \left( \frac{\sqrt{\mu}}{1 + \sqrt{\mu}} - 0,3 \mu \right)$$

or the extended Debye-Hückel equation:

$$\log \gamma_i = -\frac{A z_i^2 \sqrt{\mu}}{1 + B a_i^o \sqrt{\mu}} + b_i \mu$$

where  $z_i$  is the ionic charge of aqueous species  $i$ , and  $A$  and  $B$  are constants dependent only on temperature.  $a_i^o$  and  $b_i$  are ion-specific parameters in the Debye-Hückel equation.

The ionic strength of the aqueous solution is

$$\mu = \frac{1}{2} \sum_i^{N_{\text{aq}}} z_i^2 \frac{n_i}{W_{\text{aq}}}$$

# Modeling of geochemical processes

## Equilibrium modeling

### Example

Calcite dissolution



$$K_p = \frac{a_{\text{Ca}^{2+}} a_{\text{HCO}_3^-}^2}{a_{\text{CaCO}_3} a_{\text{H}_2\text{O}} a_{\text{CO}_2}} \quad a_{\text{Ca}^{2+}} = \frac{K_p a_{\text{CaCO}_3} a_{\text{H}_2\text{O}} a_{\text{CO}_2}}{a_{\text{HCO}_3^-}^2}$$

$$m_{\text{Ca}^{2+}} = \frac{K_p m_{\text{CaCO}_3} \gamma_{\text{CaCO}_3} m_{\text{H}_2\text{O}} \gamma_{\text{H}_2\text{O}} m_{\text{CO}_2} \gamma_{\text{CO}_2}}{\gamma_{\text{Ca}^{2+}} m_{\text{HCO}_3^-}^2 \gamma_{\text{HCO}_3^-}^2}$$



# Modeling of geochemical processes

## Equilibrium modeling

### Mole-balance equation

The total moles of an element in the system are the sum of the moles initially present in the

- *pure-phase*
- *solid-solution assemblages,*
- *aqueous phase,*
- *exchange assemblage,*
- *surface assemblage,*
- *gas phase,* and
- *diffuse layers* of the surfaces.

The following function is the general **mole-balance equation**:

# Modeling of geochemical processes

## Equilibrium modeling

$$f_m = \left( T_m - \sum_p^{N_p} b_{m,p} n_p - \sum_{ss}^{SS} \sum_{p_{ss}}^{N_{ss}} b_{m,p_{ss}} n_{p_{ss}} \right) -$$

$$- \sum_i^{N_{aq}} b_{m,i} n_i - \sum_e^E \sum_{i_e}^{N_e} b_{m,i_e} n_{i_e} -$$

$$- \sum_s^S \sum_k^{K_s} \sum_{i_{(sk)}}^{N_{sk}} b_{m,i_{(sk)}} n_{i_{(sk)}} -$$

$$- \sum_g^{N_g} b_{m,g} n_g - \sum_s^S \sum_i^{N_{aq}} b_{m,i} n_{i,s}$$

$f_m$  is zero when mole-balance is achieved

$T_m$  is the total moles of the element

$N_p$  is the number of phases

$SS$  is the number of solid solutions

$N_{ss}$  is the number of components in solid solution  $ss$

$N_{aq}$  is the number of aqueous species

$E$  is the number of exchangers

$N_e$  is the number of exchange species for exchange site  $e$

$S$  is the number of surfaces

$K_s$  is the number of surface types for surface  $s$

$N_{sk}$  is the number of surface species for surface type  $s_k$

$N_g$  is the number of gas-phase components

The moles of each entity are represented by  $n_p$  for phases,  $n_{p_{ss}}$  for components in a solid solution,  $n_i$  for aqueous species,  $n_{i_e}$  for the exchange species of exchange site  $e$ ,  $n_{i_{(sk)}}$  for surface species for surface site type  $s_k$ ,  $n_g$  for the gas components, and  $n_{i,s}$  for aqueous species in the diffuse layer of surface  $s$ .

The moles of element  $m$  per mole of each entity are represented by  $b_{m,}$  with an additional subscript to define the relevant entity

# Modeling of geochemical processes

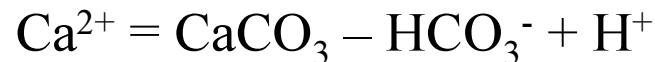
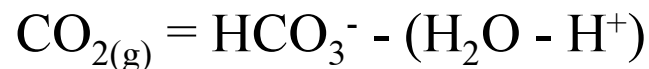
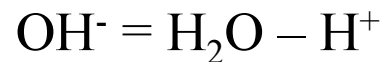
## Equilibrium modeling

### Example

- Distribution of carbonate species in closed carbonate system

Components of base,  $\mathbf{B} = [\text{CaCO}_3, \text{H}^+, \text{H}_2\text{O}, \text{HCO}_3^-]$

Secondary species,  $\text{Ca}^{2+}, \text{OH}^-, \text{CO}_3^{2-}, \text{H}_2\text{CO}_3, \text{CO}_{2(\text{g})} \dots$



Total mass of carbonate species,  $m_{\text{TOT}} = 10^{-3} \text{ mol}$

$$f_m = 10^{-3} - m_{\text{CaCO}_3} - m_{\text{HCO}_3^-} - m_{\text{CO}_3^{2-}} - m_{\text{H}_2\text{CO}_3} - m_{\text{CO}_{2(\text{g})}}$$

# Modeling of geochemical processes

## Equilibrium modeling

### Aqueous Charge Balance

$$f_z = T_z - \sum_i^{N_{aq}} z_i n_i - \left( \sum_s^S \sum_k^{K_s} \sum_{i(s_k)}^{N_{s_k}} z_{i(s_k)} n_{i(s_k)} + \sum_s^S \sum_i^{N_{aq}} z_i n_{i,s} \right) - \sum_e^E \sum_{i_e}^{N_e} z_{i_e} n_{i_e}$$

where  $f_z$  is zero when charge balance has been achieved. If the diffuse-layer composition is explicitly calculated, a separate charge-balance equation is included for each surface and the sum of the terms in the parentheses will be zero when surface charge balance is achieved. If the diffuse-layer composition is not calculated, the second term inside the parentheses is zero.

# Modeling of geochemical processes

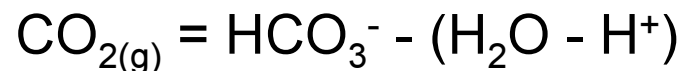
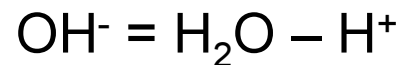
## Equilibrium modeling

### Example

#### *Charge balance of carbonate system*

Components of base,  $\mathbf{B} = [\text{Ca}^{2+}, \text{H}^+, \text{H}_2\text{O}, \text{HCO}_3^-]$

Secondary species  $\text{OH}^-, \text{CO}_3^{2-}, \text{H}_2\text{CO}_3, \text{CO}_{2(\text{g})} \dots$



Charge balance

$$f_z = 2m_{\text{Ca}^{2+}} + m_{\text{H}^+} - m_{\text{OH}^-} - m_{\text{HCO}_3^-} - 2m_{\text{CO}_3^{2-}}$$