

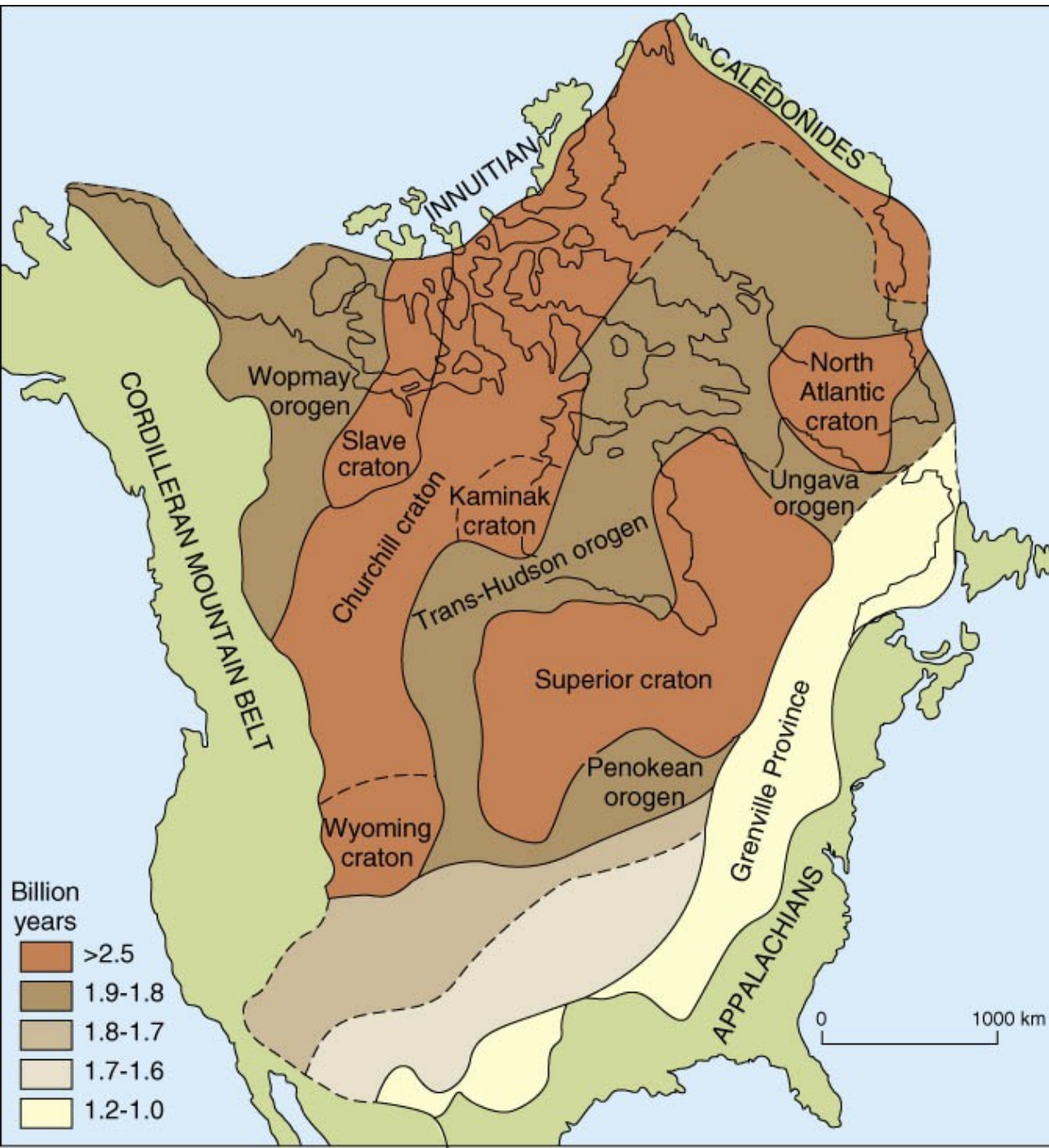
4

Canadian shield – not covered during Paleozoic, re-mainder of craton (light brown) was periodically covered by epeiric (shallow, inland) seas. Green (basins) & yellow (domes) = local areas of gentle warpage of platform sedimentary rocks.



Appalachian Basin





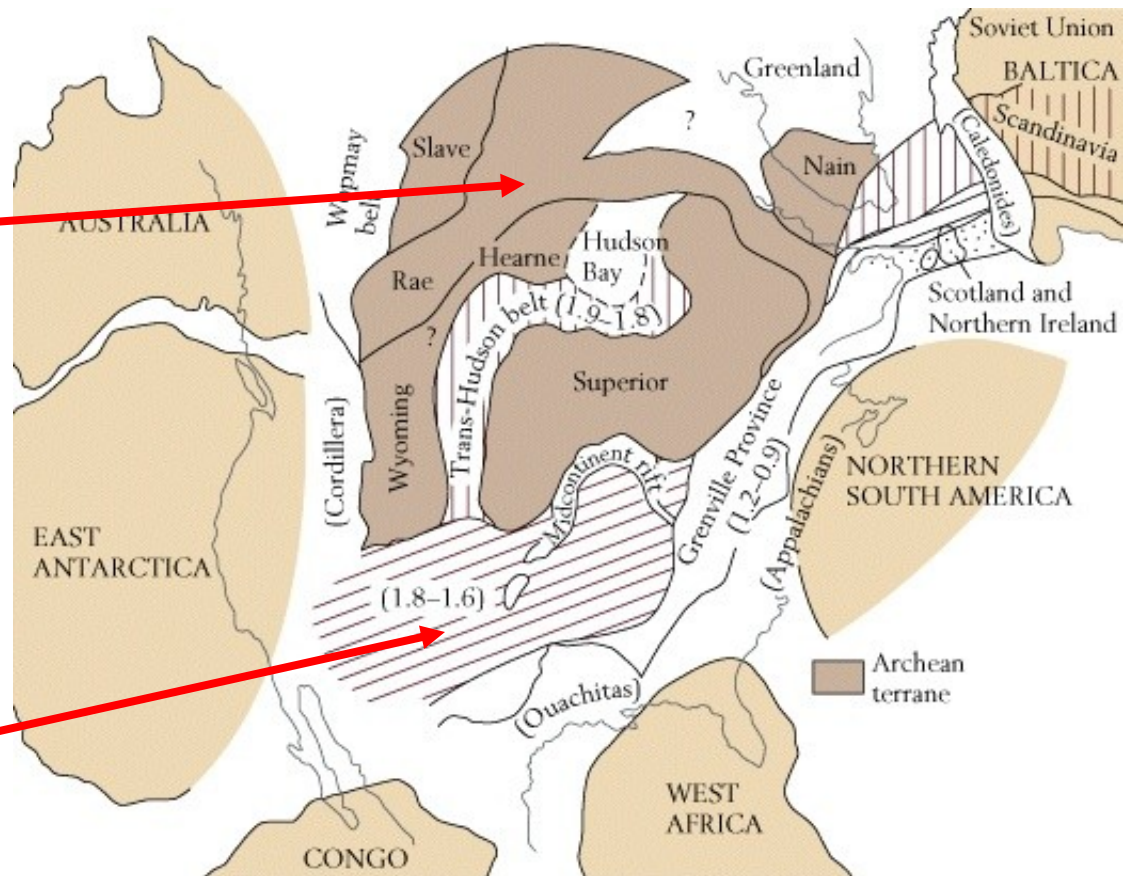
Hlavní část **laurenského štítu** byla vytvořena při **transhudsonské** orogenezi (1,8-1,9 Ga), menší části byly přiřčleněny při **mazatzalské** a **penocké** orogenezi (1,5-1,8). Nejmladší proterozoickou orogenezi je **grenvilská**, která skončila před 1,0 Ga.

Continental Accretion


- During Proterozoic, Laurentia was growing by accretion

Phase 1:
1.95-1.85 bya
Suturing of ~six
microcontinents

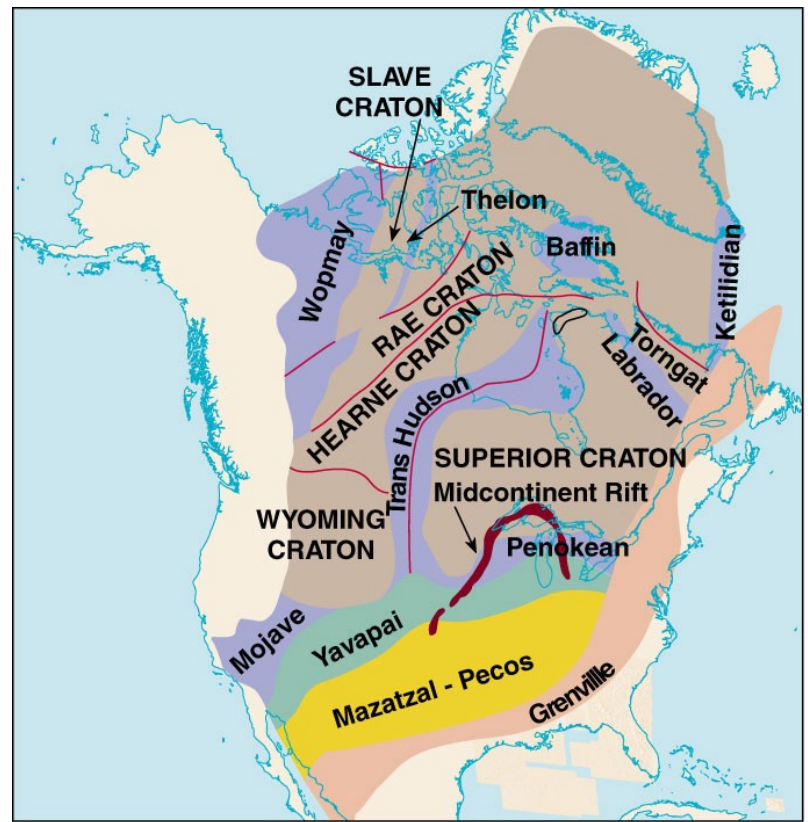
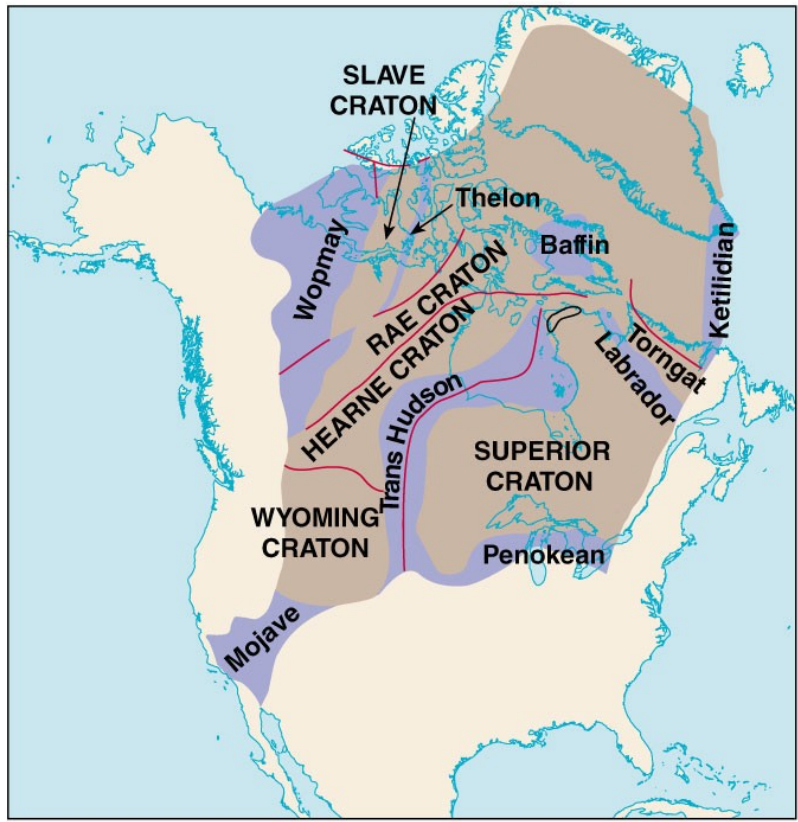
Phase 2:
1.8-1.6 bya
Suturing of island
arc to south

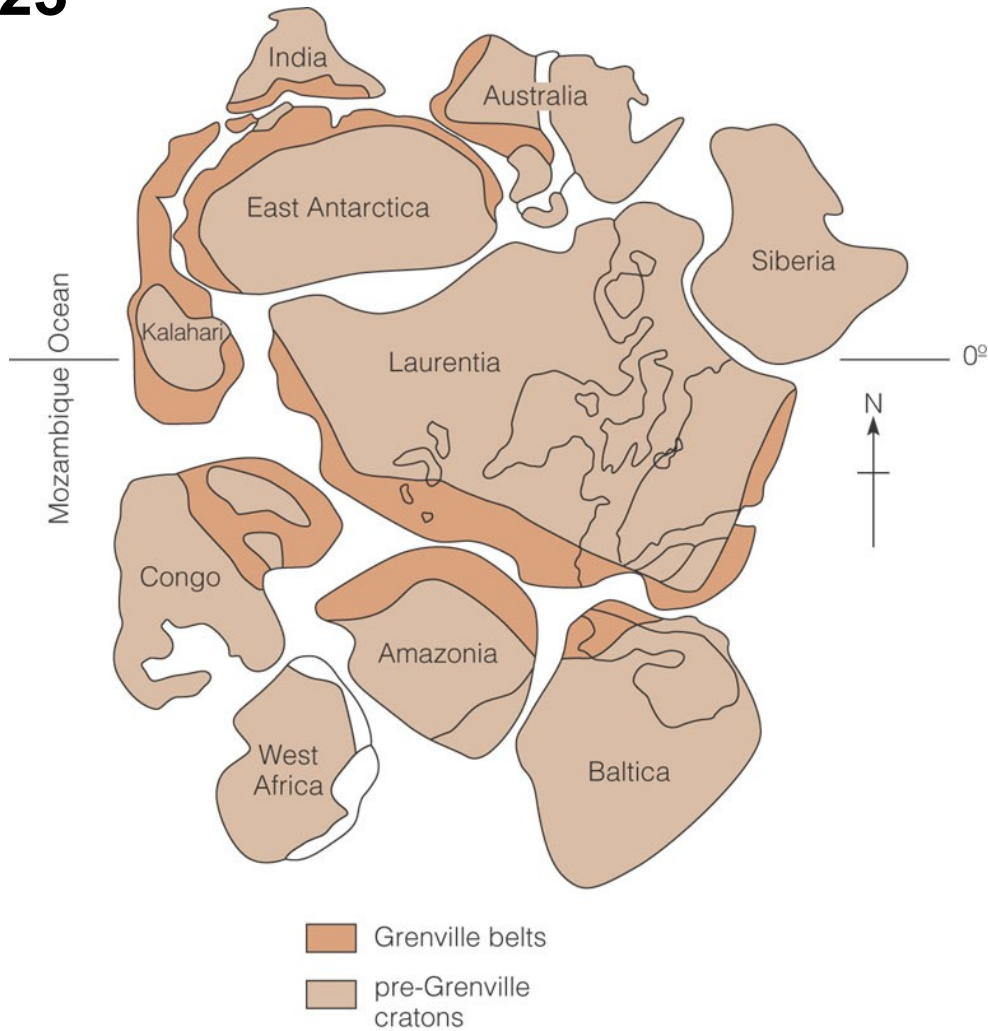


4

-  900 million - 1.2 billion
-  1.6 billion- 1.75 billion
-  1.75 billion- 1.8 billion
-  1.8 billion - 2.0 billion
-  2.5 billion - 3.0 billion

During Proterozoic, Laurentia grew to southeast by accretion of other cratons. Collision zones = orogenic belts. Brown masses – Archean age.





**Rodinia – final
“assembly” during
Grenville Orogeny
(1.3 to 1.0 b.y. ago),
then fragmentation
began 750 m.y. ago.**

Fanerozoické mobilní zóny

Kaledonsko – apalačská mobilní zóna

Kordilerská mobilní zóna

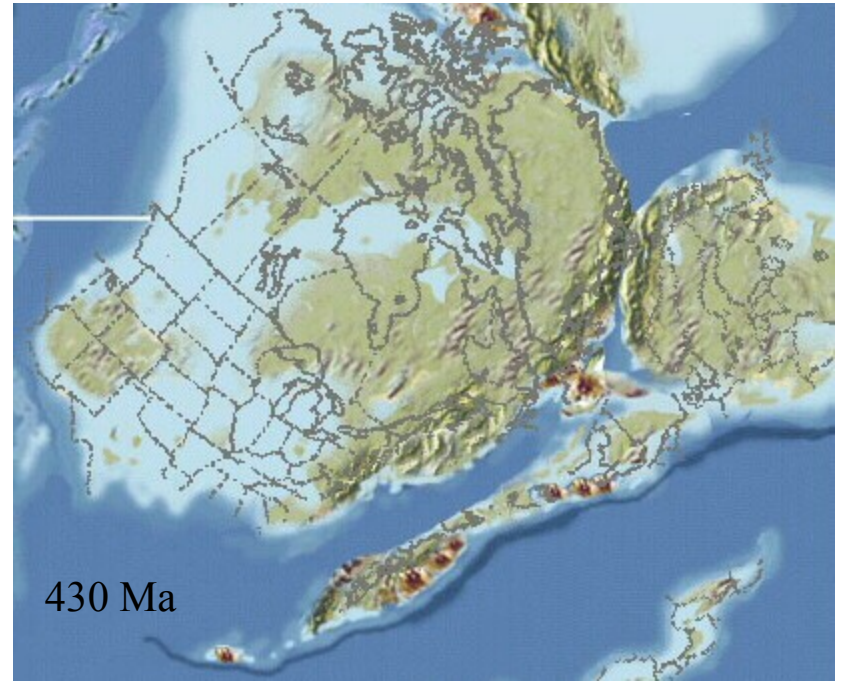
Franklinsko-inuitská mobilní zóna

KALEDONSKO-VARISKÁ OROGENEZE

Grónské kaledonidy

**Apalače – terání stavba, kaledonské (takonská, akadská)
i variské (alleghanská) fáze**

**Pásmo Quatchita-Marathon – alleghanská fáze, kolize s
jihoamerickou částí Gondwany**



APALAČE

Laurentia – Humber, Valley and Ridge, Blue-Ridge terranes

**Centrální zóna – hlavně vulkanické oblouky (Notre Dam, Dunnage, Exploit, Piedmont aj.)
a akreční melanz Iapetu**

Gondwanské terány – Avalonia, Carolina, Meguma, Gondwana

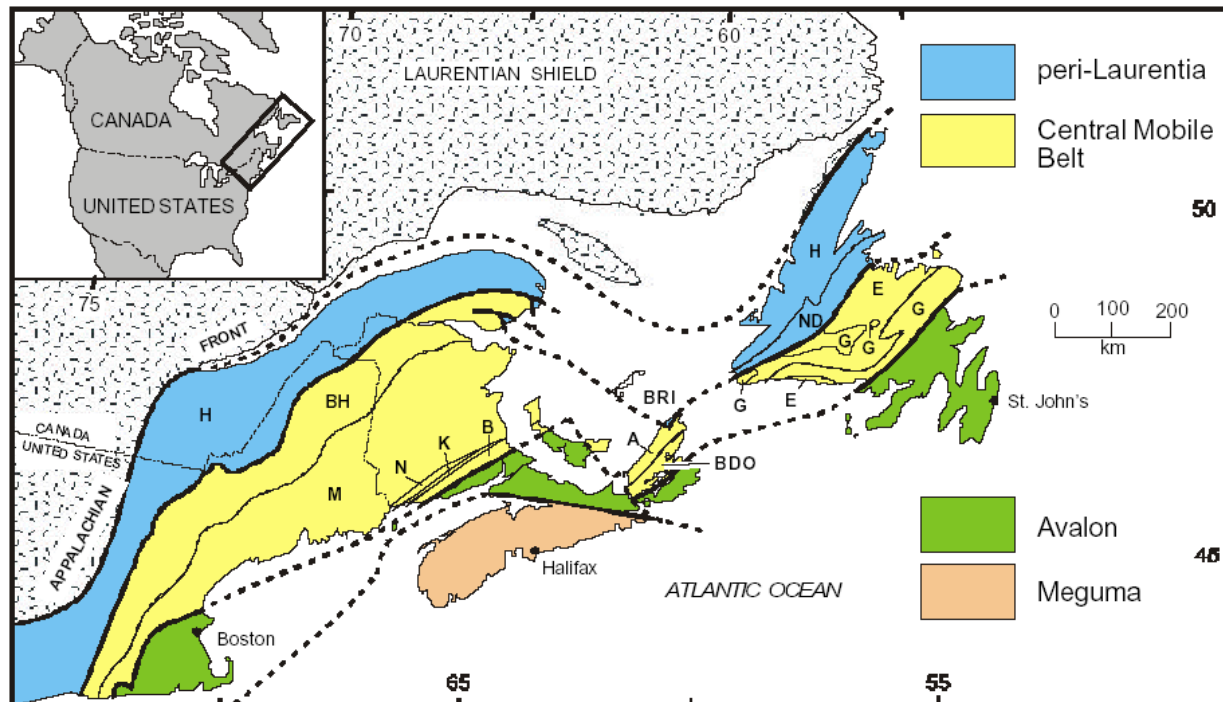
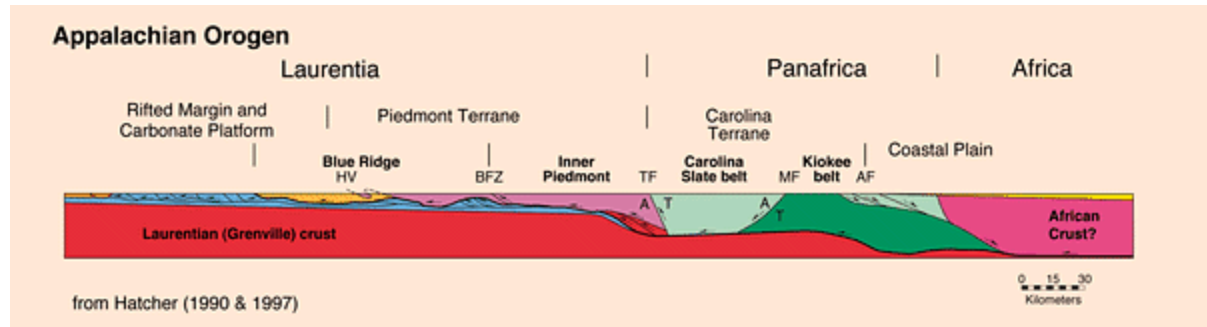
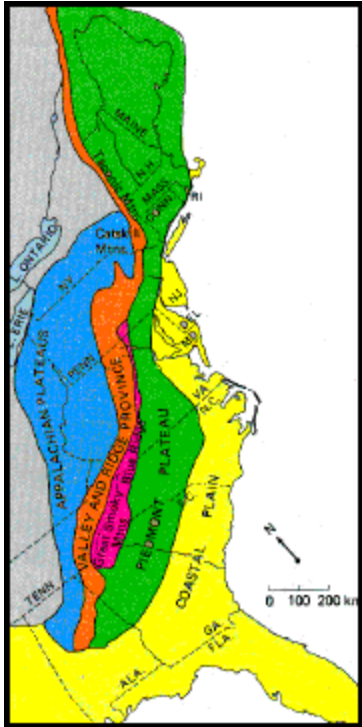
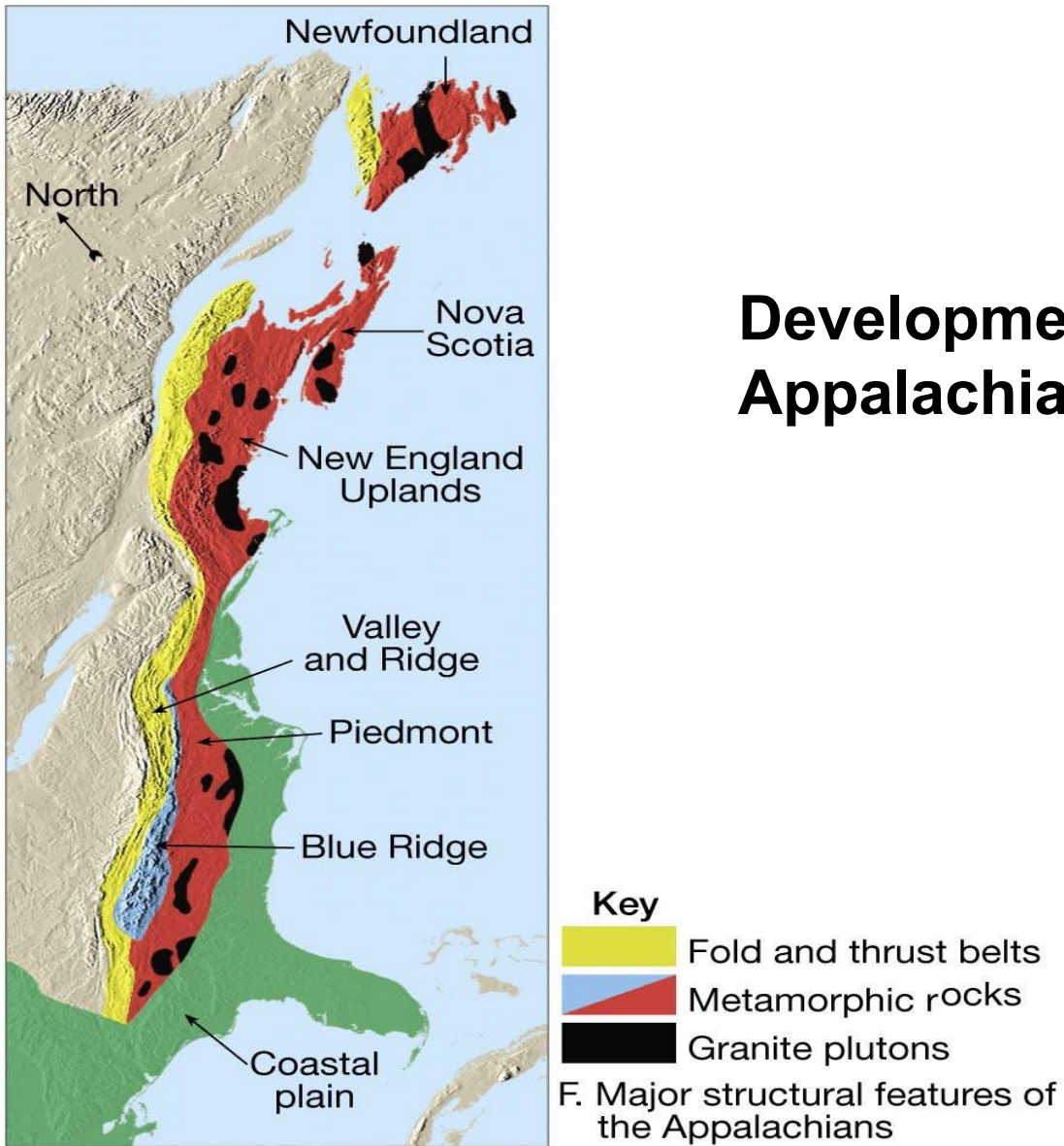


Figure 1. Simplified map of the northern Appalachian orogen showing crustal blocks and terranes (A, Aspy; B, Brookville; BDO, Bras d'Or; BH, Bronson Hill; BRI, Blair River inlier; E, Exploits; G, Gander; H, Humber; K, Kingston; M, Miramichi; N, New River; ND, Notre Dame;



Development of the Appalachian Mountains



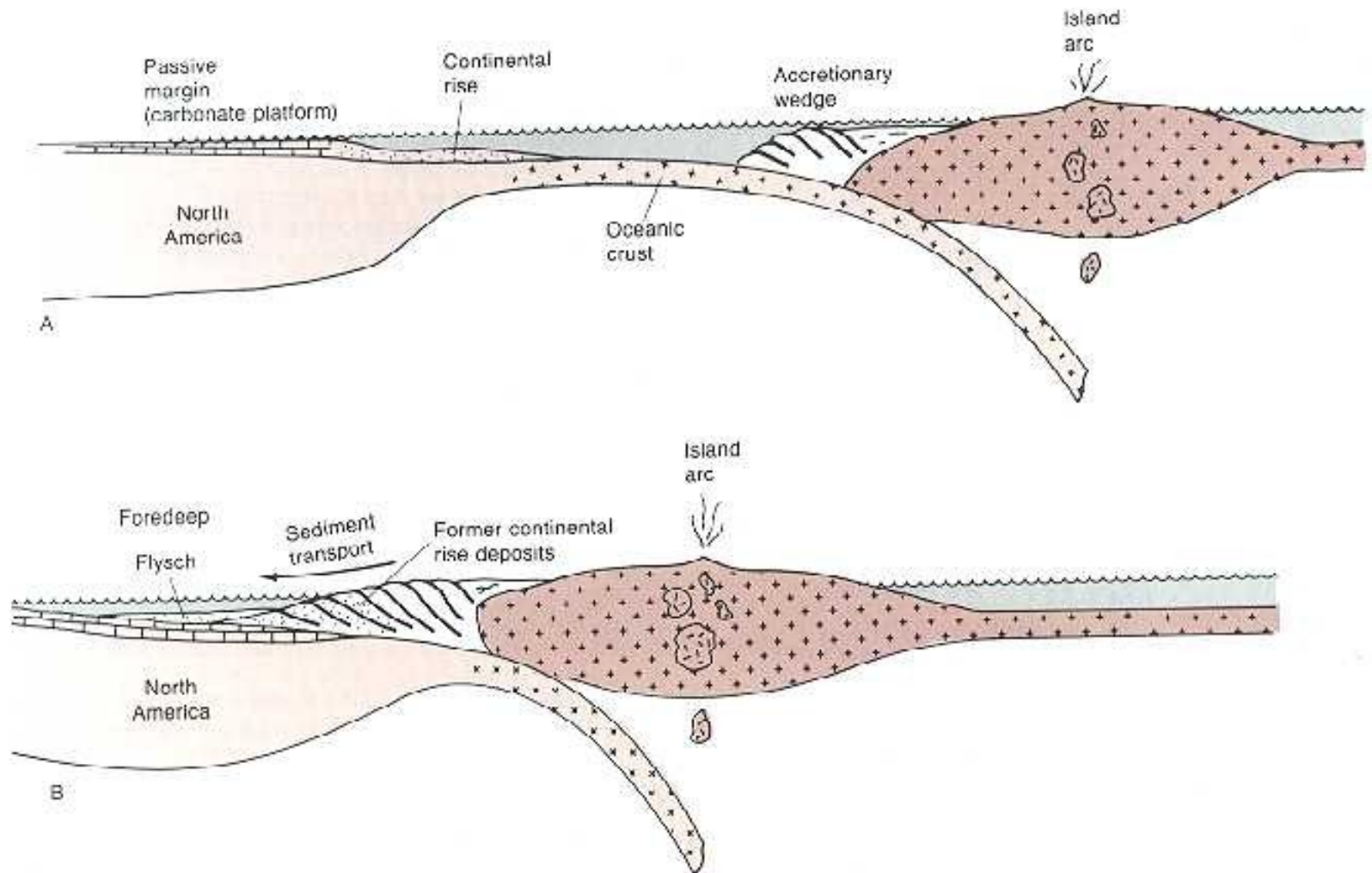
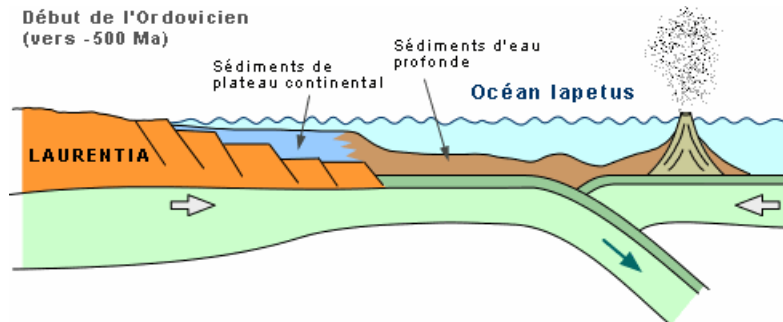


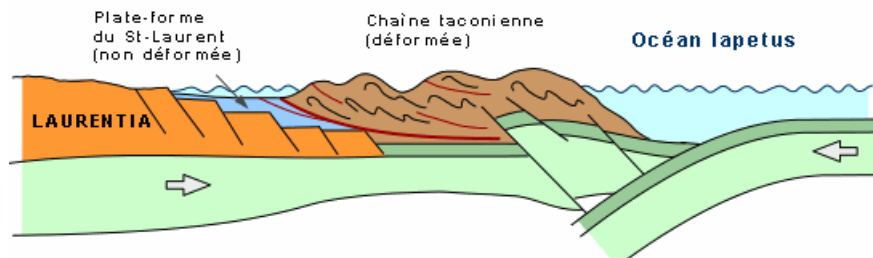
FIGURE 6-29 Diagrammatic illustration of the plate movements that in mid-Ordovician time transformed the passive margin of northeastern North America (A) into a foredeep (B). This happened when the passive margin, which supported a carbonate bank, encountered an island arc. The accretionary wedge

bordering the island arc was thrust over the continental margin, as were deep-water deposits that had accumulated along the continental rise of North America, in front of the carbonate bank. The island arc thus joined North America as an exotic terrane.

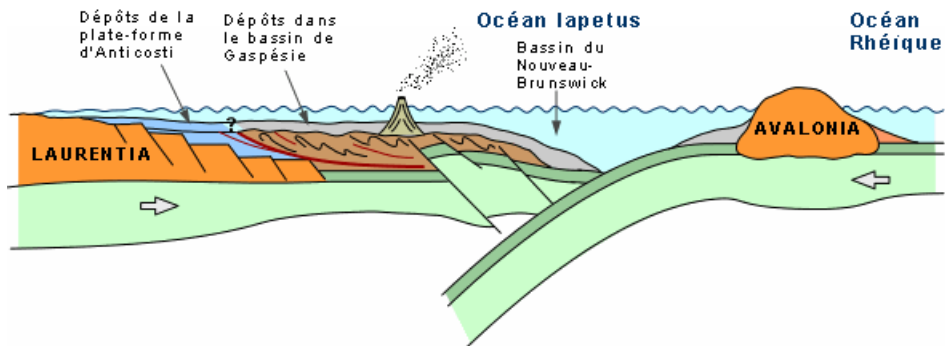
Début de l'Ordovicien
(vers -500 Ma)



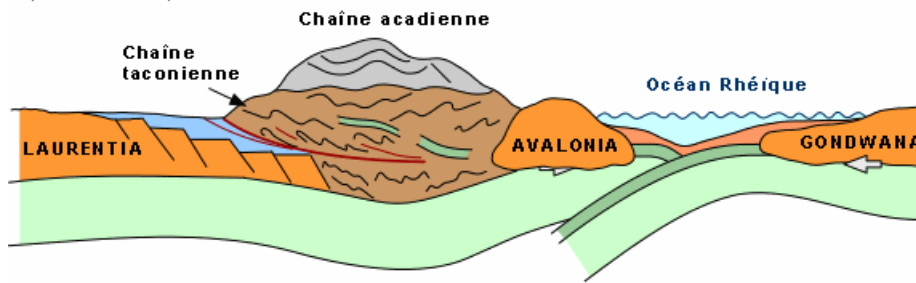
Milieu de l'Ordovicien
(vers -460 Ma)



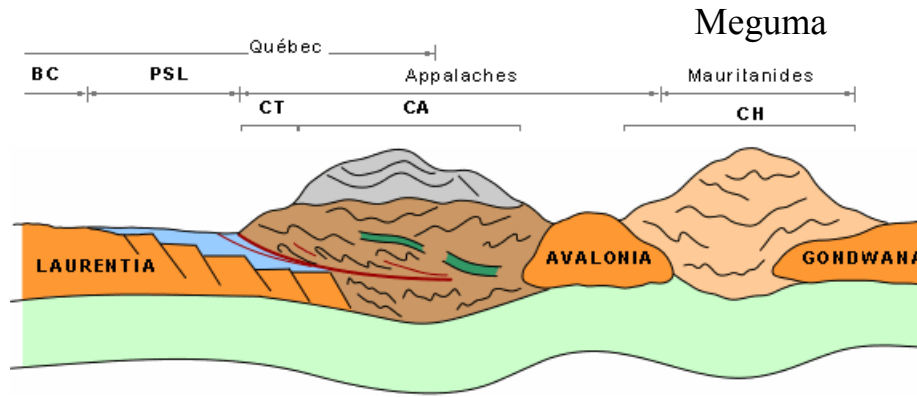
Milieu du Silurien
(vers -420 Ma)



Milieu du Dévonien
(vers -390 Ma)

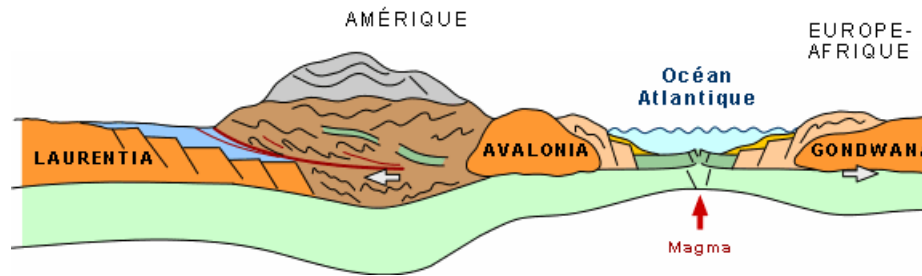


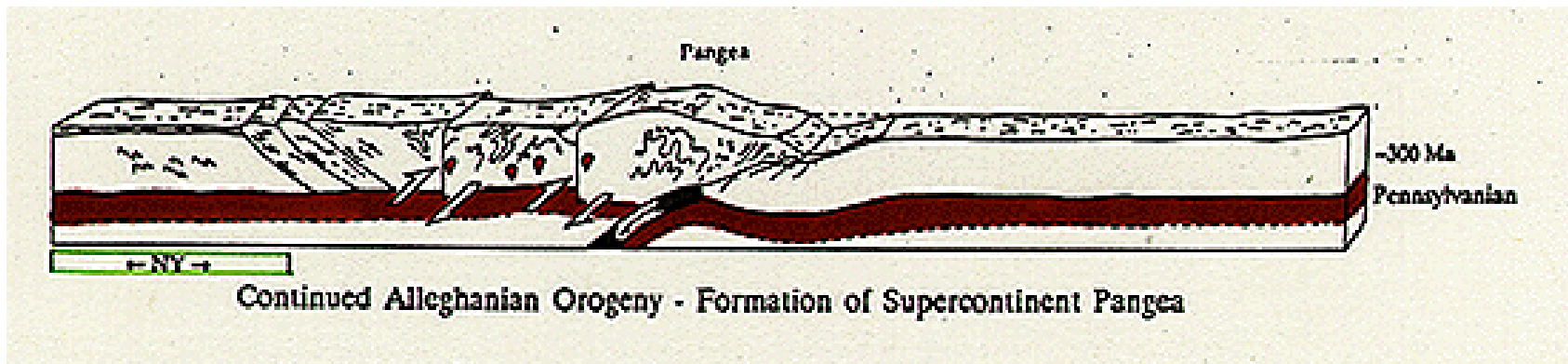
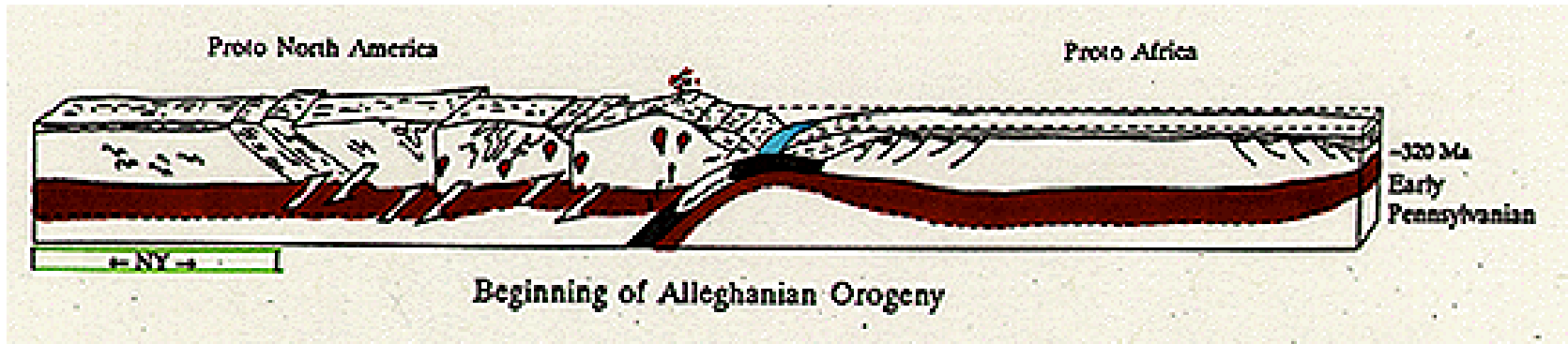
Fin du Carbonifère
(vers -310 Ma)

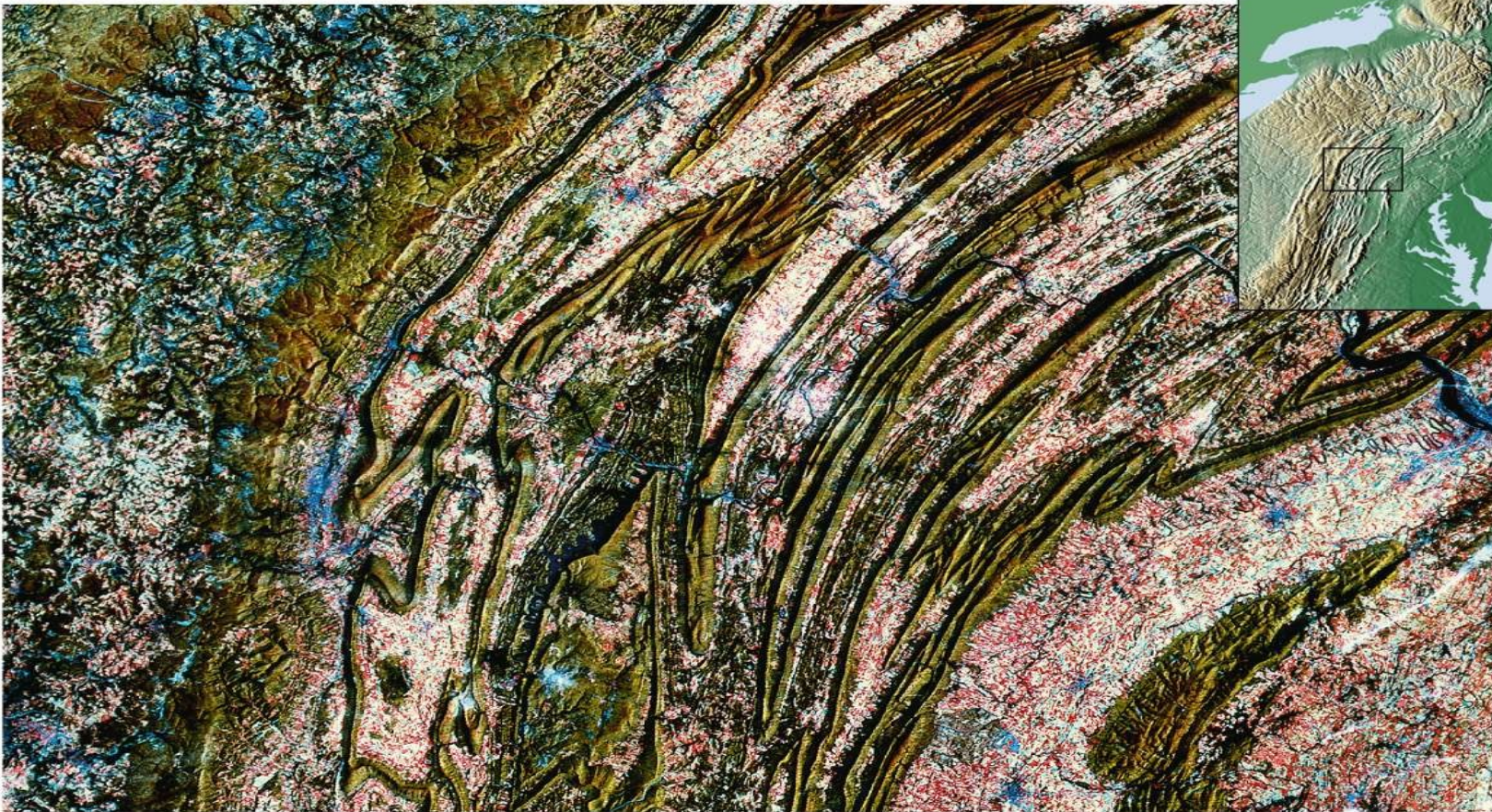


BC Bouclier canadien
Jurassique
(vers -160 Ma)

CT Chaîne taconienne





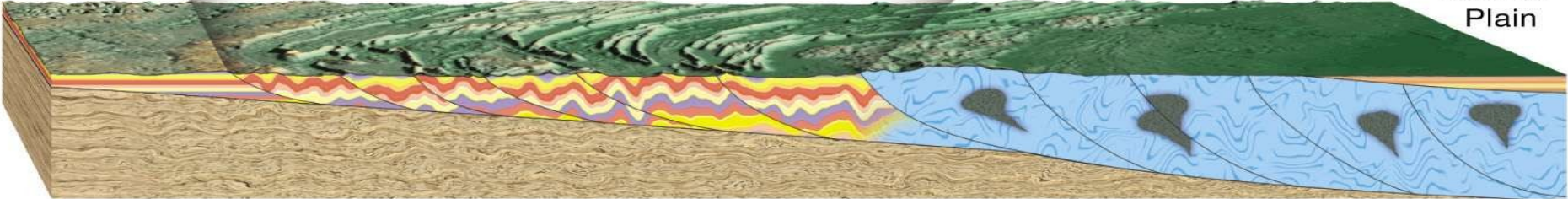


Valley and Ridge

Blue Ridge

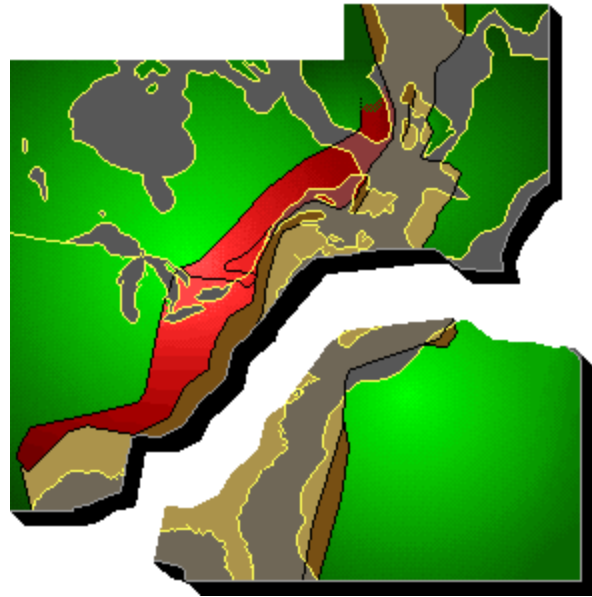
Piedmont

Coastal
Plain



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Folded mountains in the Valley and Ridge Province



Apalače představují složený orogen, jehož tvorba probíhala jak během kaledonských fází tak během variských fází . V jejich rámci můžeme rozlišit několik teránů.

Údaje seismiky ukazují, že celá jižní část Apalačí je pravděpodobně podstýlána velkou zónou odlepení a celý horský hřeben je alochtonní. Podobné struktury se dají pozorovat v seismických řezech i v severních Apalačích. Pro geologickou stavbu jsou významné alochtonní terány.

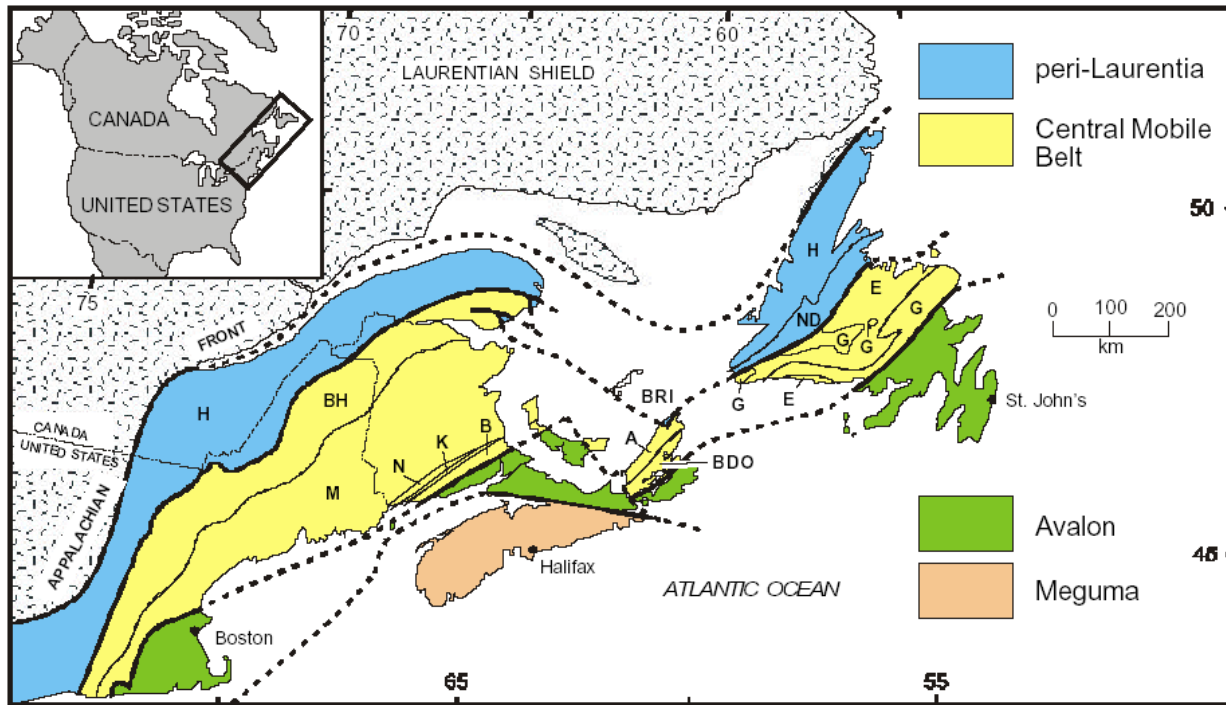


Figure 1. Simplified map of the northern Appalachian orogen showing crustal blocks and terranes (A, Aspy; B, Brookville; BDO, Bras d'Or; BH, Bronson Hill; BRI, Blair River inlier; E, Exploits; G, Gander; H, Humber; K, Kingston; M, Miramichi; N, New River; ND, Notre Dame);

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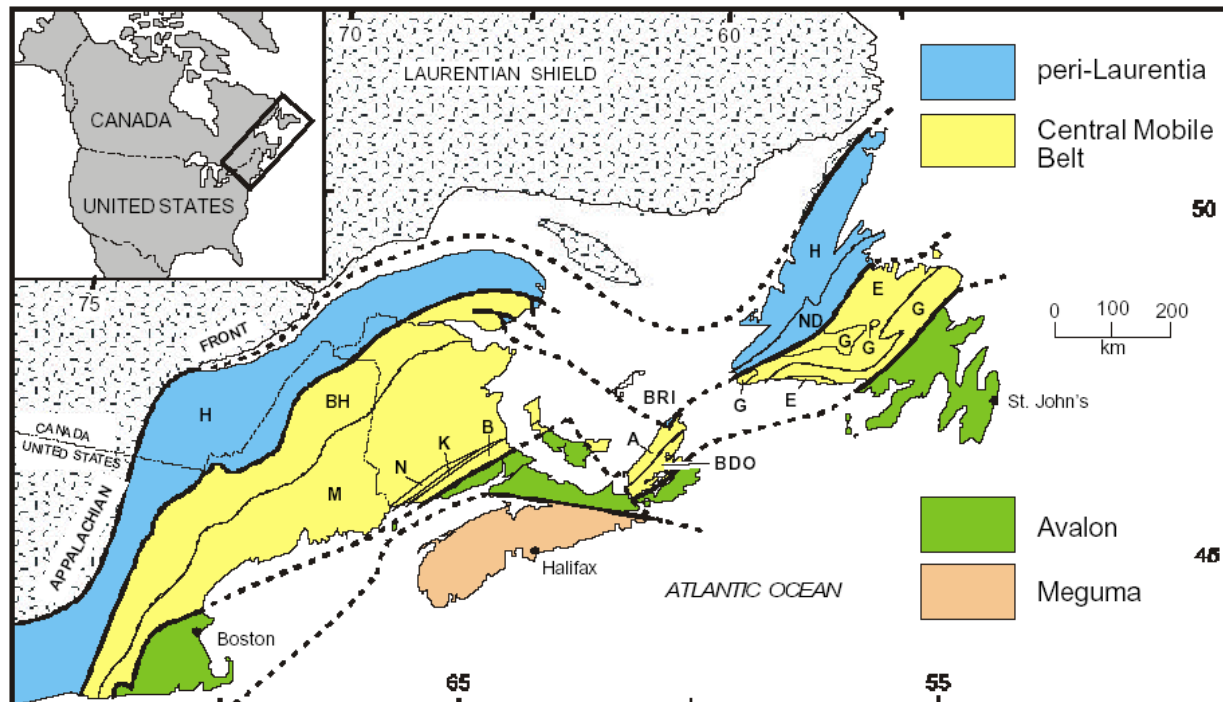


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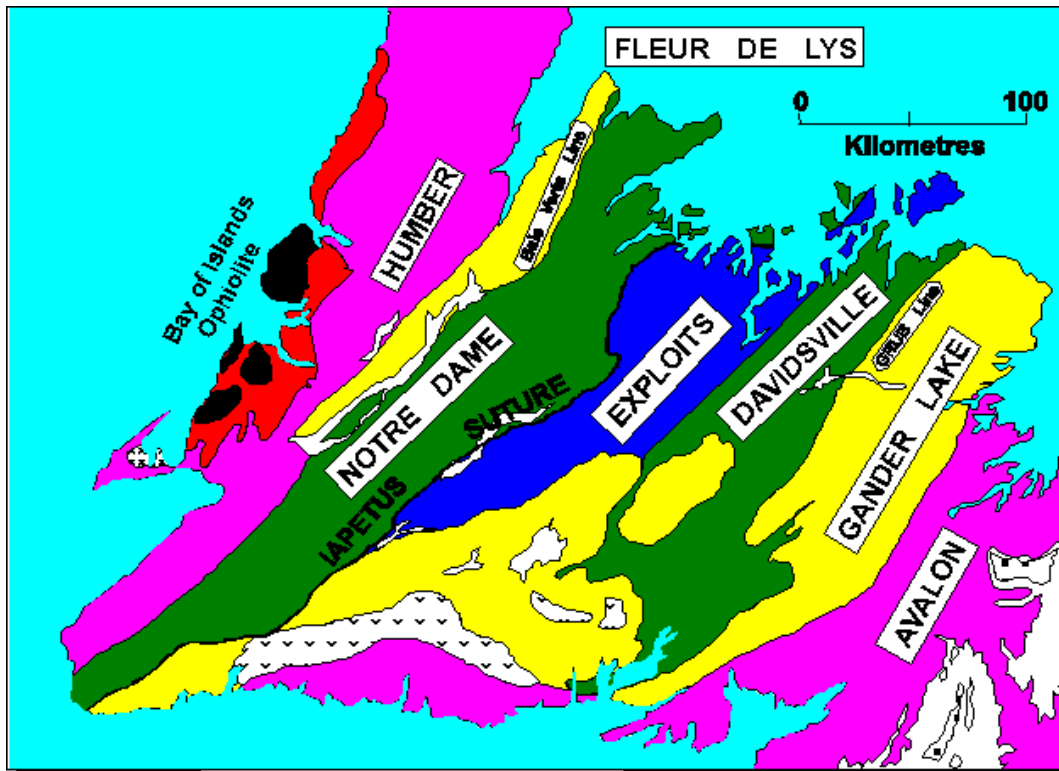
Terány skupiny I je tvořen sedimenty nejvyššího proterozoika a spodního paleozoika a svrchnoproterozoickým krystalinikem. Reprezentuje **fragments Laurentie** nebo **mikrokontinenty** původně situované blízko jejího kontinentálního okraje.

Terán Humber

Tento terán zahrnuje miogeosynklinálu situovanou na laurentinském pasivním kontinentálním okraji a takonské alochtony. Táhne se od SZ Newfoundlandu do zóny **Valley and Ridge a Blue Ridge na jihu**. Basement se skládá z **grenvilských rul** (1,0Ma) na kterých spočívají klastické a karbonátové riftové **sedimenty** kambria až spodního ordoviku

Na kontinentální okraj jsou obdukovány **dva typy teránů** – **Fleur de Lys Supergroup** (metapelite a meetapsamity interpretované jako sedimenty kontinentálního svahu na jv od karbonátové lavice) a **velké alochtony** obsahující **ofiolity** jako jsou **Bay of Islands Complex**, který byl obdukován v Ilnvirnu.

První výskyt **flyšových** úlomků v slepencích karbonátového okraje datuje přičlenění prvních teránů k americkému kontinentálnímu okraji v nejvyšším **arenigu-Ilnvirnu**.



Terány II a III reprezentuje **ostrovní oblouky**, jedná se o **ofiolity** a přidruženou **akreční melánž**, Zahnuje **takonskou** sekvenci v USA a Kanadě (Notre Dame) a terány **Dunnage a Exploit a Gander** v Kanadě a terán **Piedmont** v jižních Apalačích.. Seismické údaje ukazují, že terány Dunnage a Gander jsou alochtonní nad spodní kontinentální kůrou a že severoamerický okraj pokračuje 70 km pod terán Tyto terány představují jenom velmi hrubé rozdělení, protože **každý z nich se skládá ze značného počtu fragmentů** (nebo dílčích teránů?) různého původu. Obsahují mnoho ofiolitů a oblouků odvozených z lapetu a deformovaných během kolize kontinentálních okrajů Laurentie a Avalonie.

Terán Piedmont (nebo Vnitřní Piedmont) je složený z několika různých **teránů a fragmentů**. Jedná se o nakupení na západ sunutých příkrovů obsahujících krystalické břidlice, ruly, amfibolity a mafické-ultramafické tělesa, které mohou reprezentovat zbytky ofiolitů.. Reflexní profily ukazují, že piedmontský terán a přilehlé pásy **břidlic Blue Ridge, Charlotte a Carolina** náležejí 6-15 km mocnému příkrovu nad autochtonním basementem. Zdá se, že byl transportován přinejmenším **260km přes** kontinentální okraj Laurentie během alleghanské orogeneze produkované kolizí Severní Ameriky a Afriky.

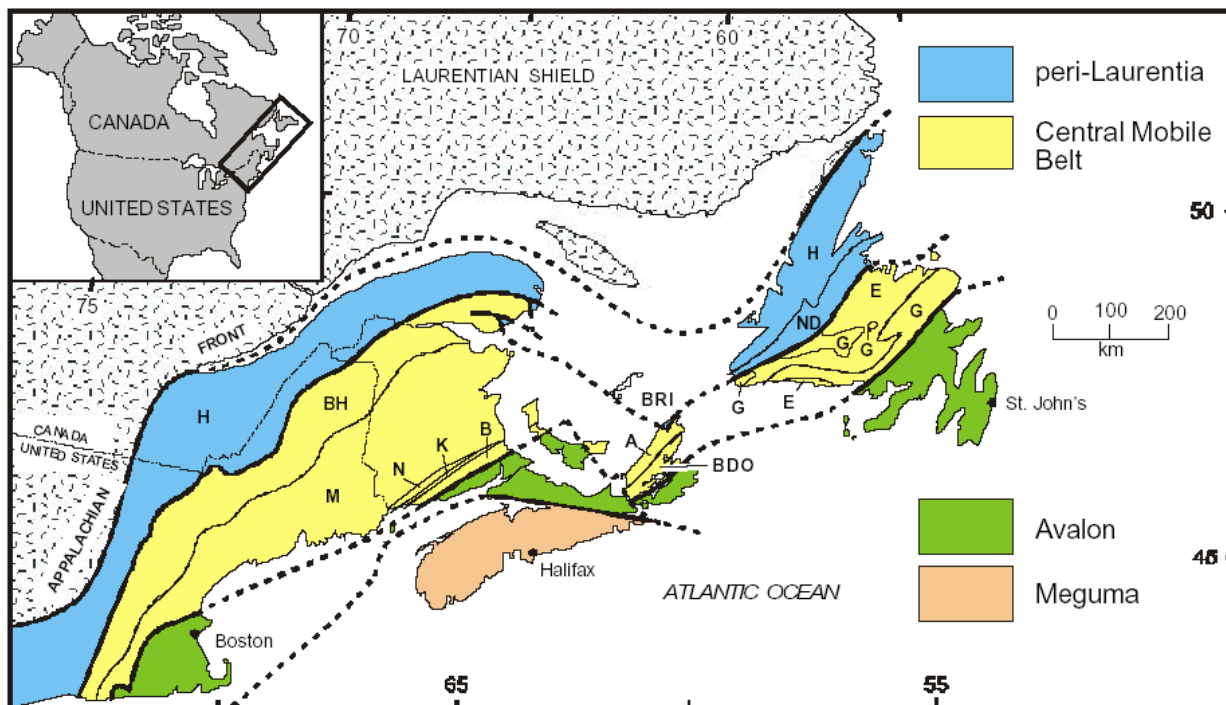
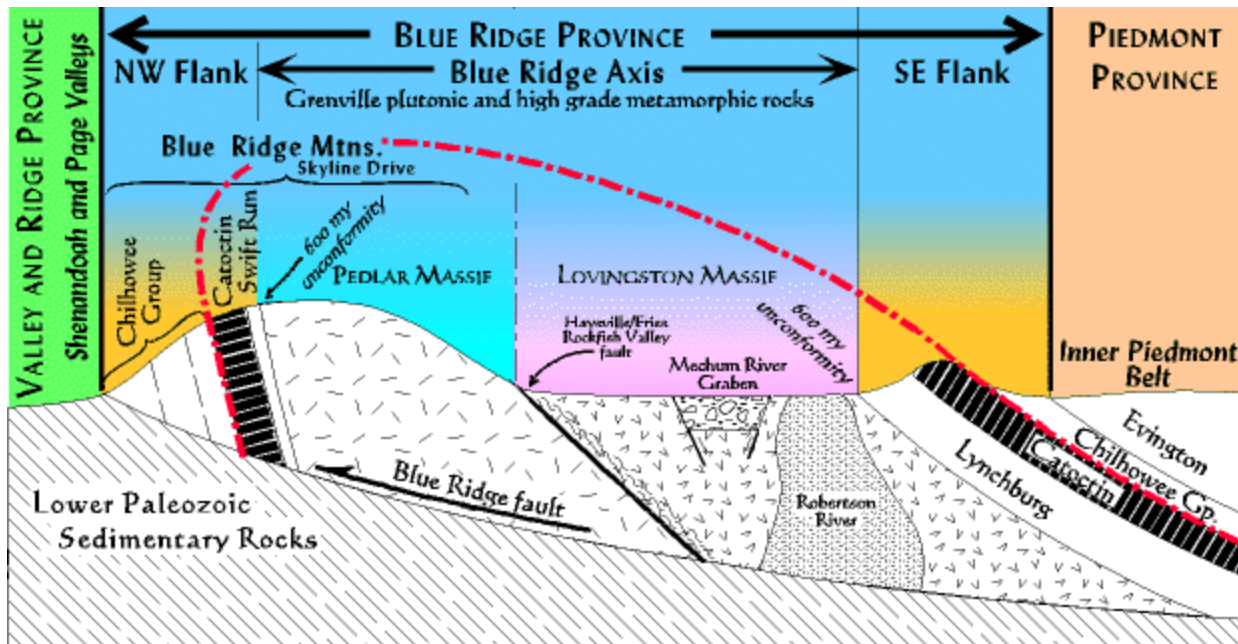
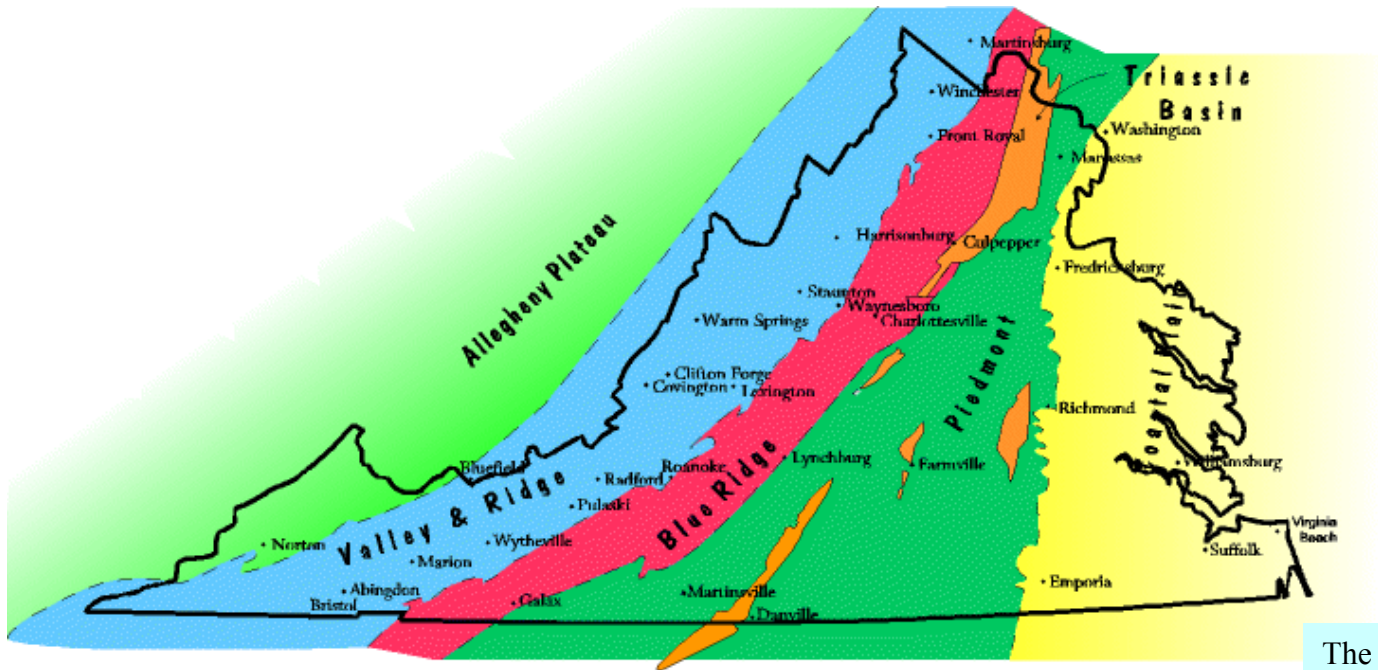


Figure 1. Simplified map of the northern Appalachian orogen showing crustal blocks and terranes (A, Aspy; B, Brookville; BDO, Bras d'Or; BH, Bronson Hill; BRI, Blair River inlier; E, Exploits; G, Gander; H, Humber; K, Kingston; M, Miramichi; N, New River; ND, Notre Dame);



The modern Blue Ridge is an overturned anticline. That is, the rocks have been arched up into a fold, and then shoved over toward the west (left) so that the rocks on the western flank are now no longer right side up (follow the red dashed line).

Notice the Blue Ridge thrust fault at the base; the [Blue Ridge province](#) has been moved westward from its site of origin, perhaps as far east as Richmond. Below the Blue Ridge fault Cambrian and Ordovician sediments ("Lower Paleozoic Sedimentary Rocks") of the Valley and Ridge extend eastward under the Blue Ridge and piedmont provinces ([cross section](#)).

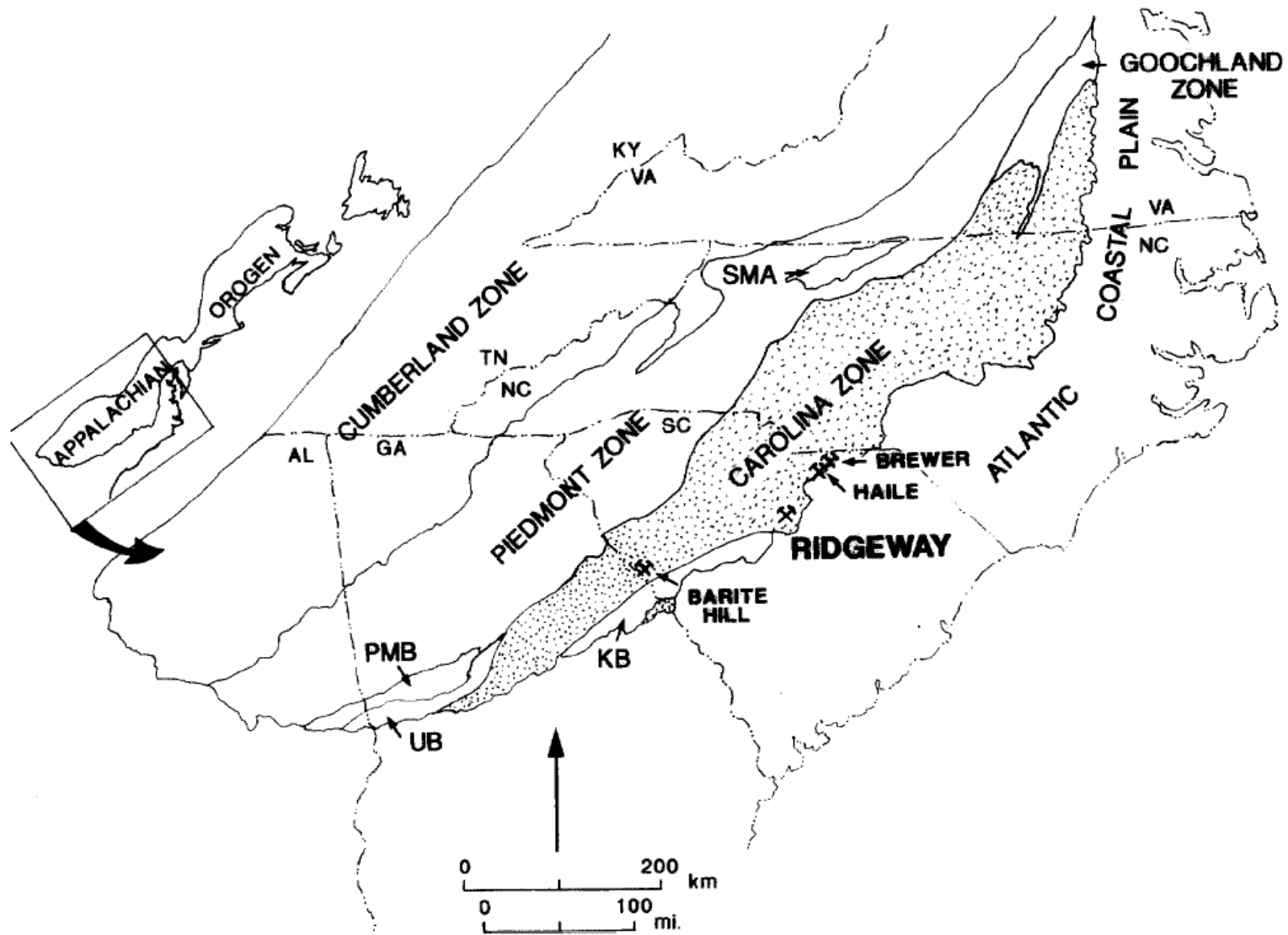
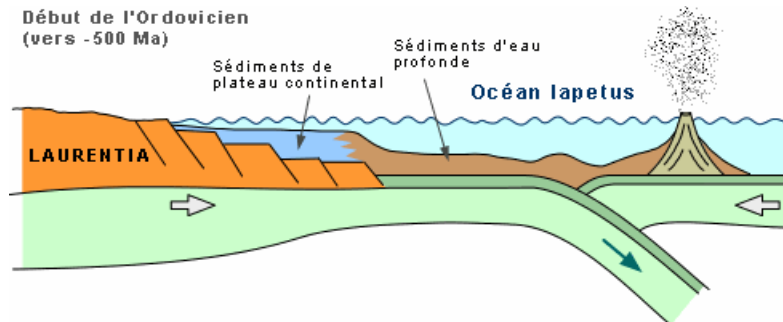


Figure 1. Tectonostratigraphic zonal map of the Southern Appalachian orogen (after Hibbard and Samson, 1995). Locations of the four gold mines active in South Carolina during the late twentieth century are posted on the map. The mines are located within the Carolina zone, whose largest subunit is the Carolina terrane subduction volcanic arc.

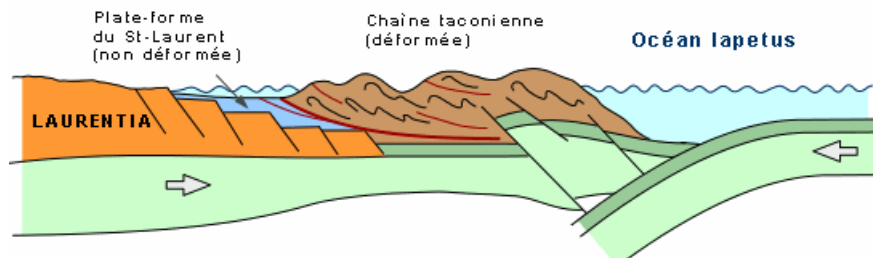
Terány skupiny IV. Zahrnuje v jižní části Apalačů **karolinské břidličné pásmo** a **avalonský sled** v severních Apalačích. **Avalonský terán** obsahuje menší **dílčí fragmenty** je označován jako **avalonský složený terán** nebo jako **avalonský superterán**. Obsahuje **ruly, vulkanické a sedimentární horniny** pozdně prekambričského až spodnopaleozoického stáří, které jsou již dlouho považovány za africko-gondwanské. **Avalonský terán** má jinou přesilurskou historii než terán Gander a západnější terány, obzvláště **odlišnou kambrickou trilobitovou faunu**.

Terán V. Jedná se o sekvenci kambrických-ordovických drob a břidlic usazenou poblíž okraje kontinentu. V Apalačích se označuje jako terán **Meguma** a podobné horniny jsou známy i ze severozápadní Afriky z **mauritanid**. Terán Meguma obsahuje kambro-ordovický **flyš**.

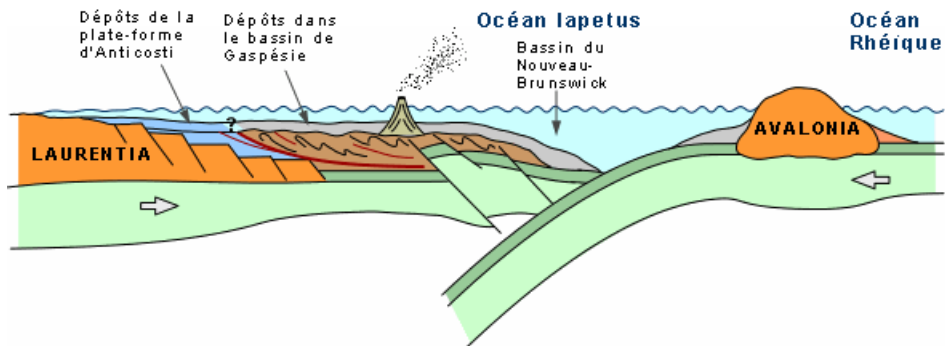
Début de l'Ordovicien
(vers -500 Ma)



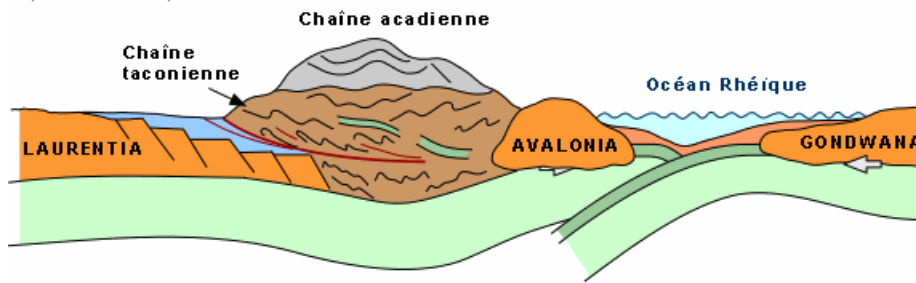
Milieu de l'Ordovicien
(vers -460 Ma)



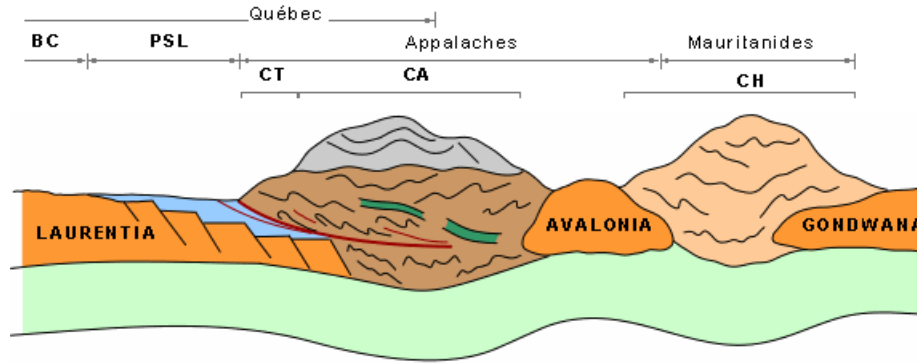
Milieu du Silurien
(vers -420 Ma)



Milieu du Dévonien
(vers -390 Ma)



Fin du Carbonifère
(vers -310 Ma)

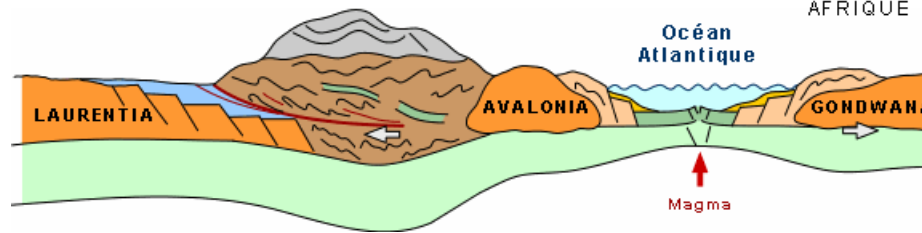


BC Bouclier canadien
Jurassique
(vers -160 Ma)

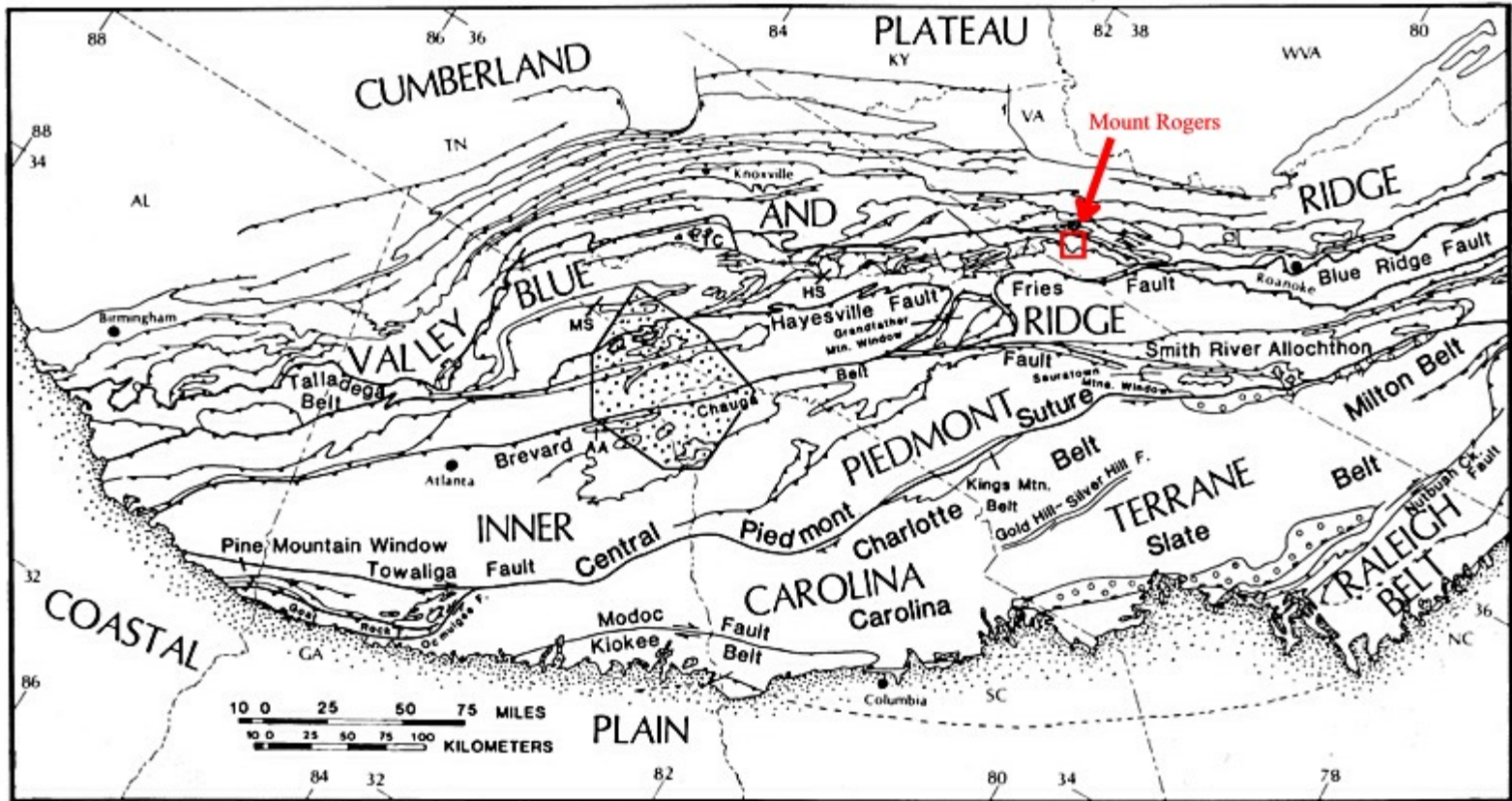
CT Chaîne taconienne

AMÉRIQUE

EUROPE-
AFRIQUE



Newfoundland	S. Appalachians
No correlative	Valley and Ridge Tectonic Province
Western Newfoundland foreland	Jefferson terrane
Fleur de Lys	Potomac terrane
Notre Dame Bay ophiolite/arc terrain	Bel Air/Chopawamsic ophiolitic arc terranes
Burlington Granodiorite	Occoquan pluton
IAPETUS SUTURE	
Exploits	Inner Piedmont terrane
Davidsville/GRUB line	Juliette terrane
Gander Lake	Uchee or Charlotte terranes
Avalonia	Carolina terrane
RHEIC SUTURE	
No correlative	Suwannee



Hopson et al., 1989

Figure 1. Lithotectonic map of the southern Appalachians. The stippled area is the region shown in Figure 2. The Hayesville fault separates the Blue Ridge into eastern and western segments. MS, Murphy Syncline; AA, Alto allochthon; TC, Tuckaleechee Cove window; HS, Hot Springs Window.

Figure 3c.

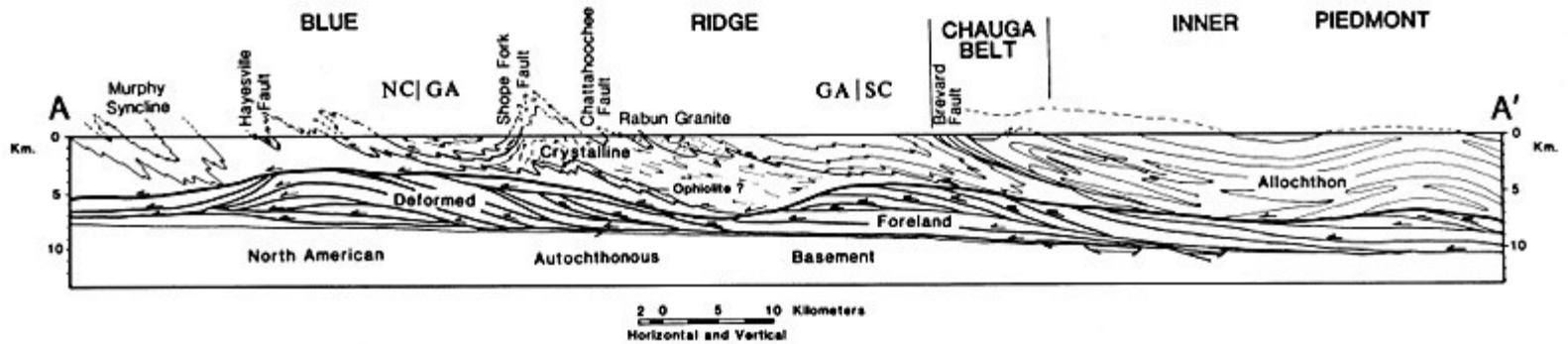
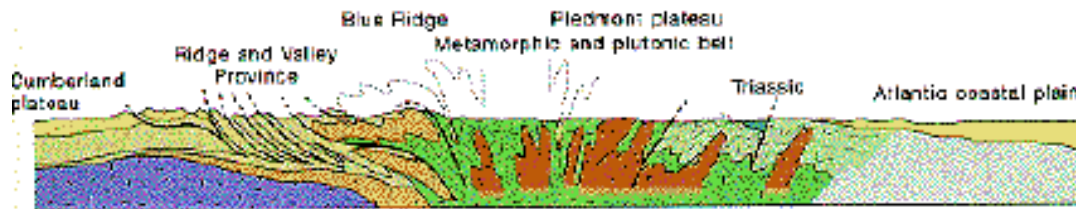
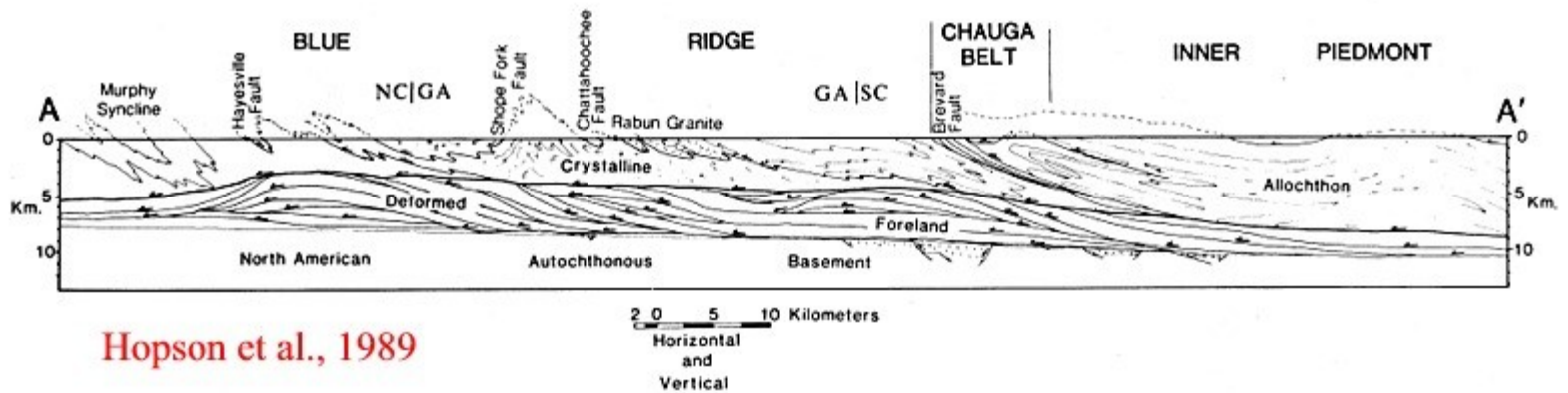
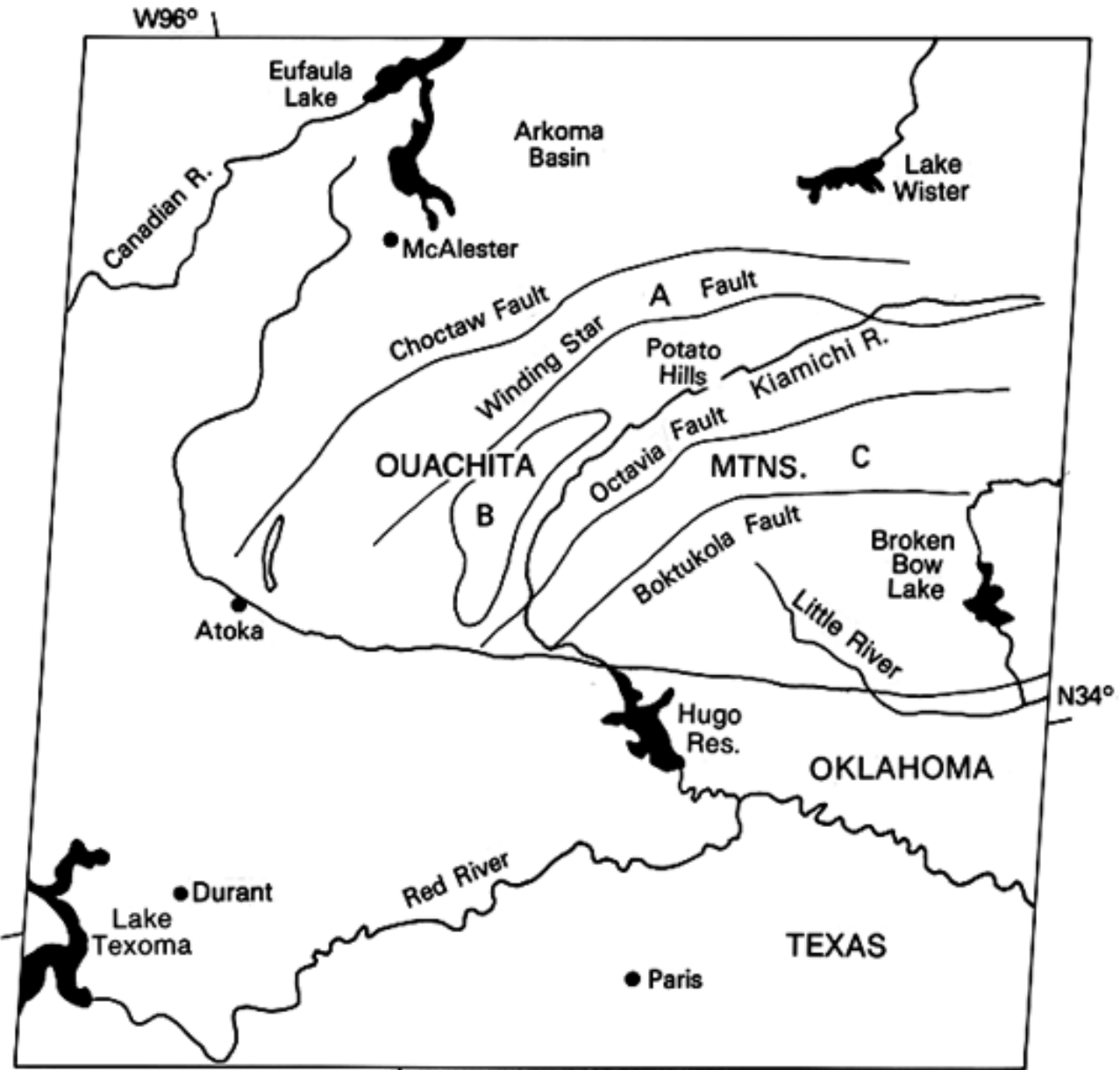


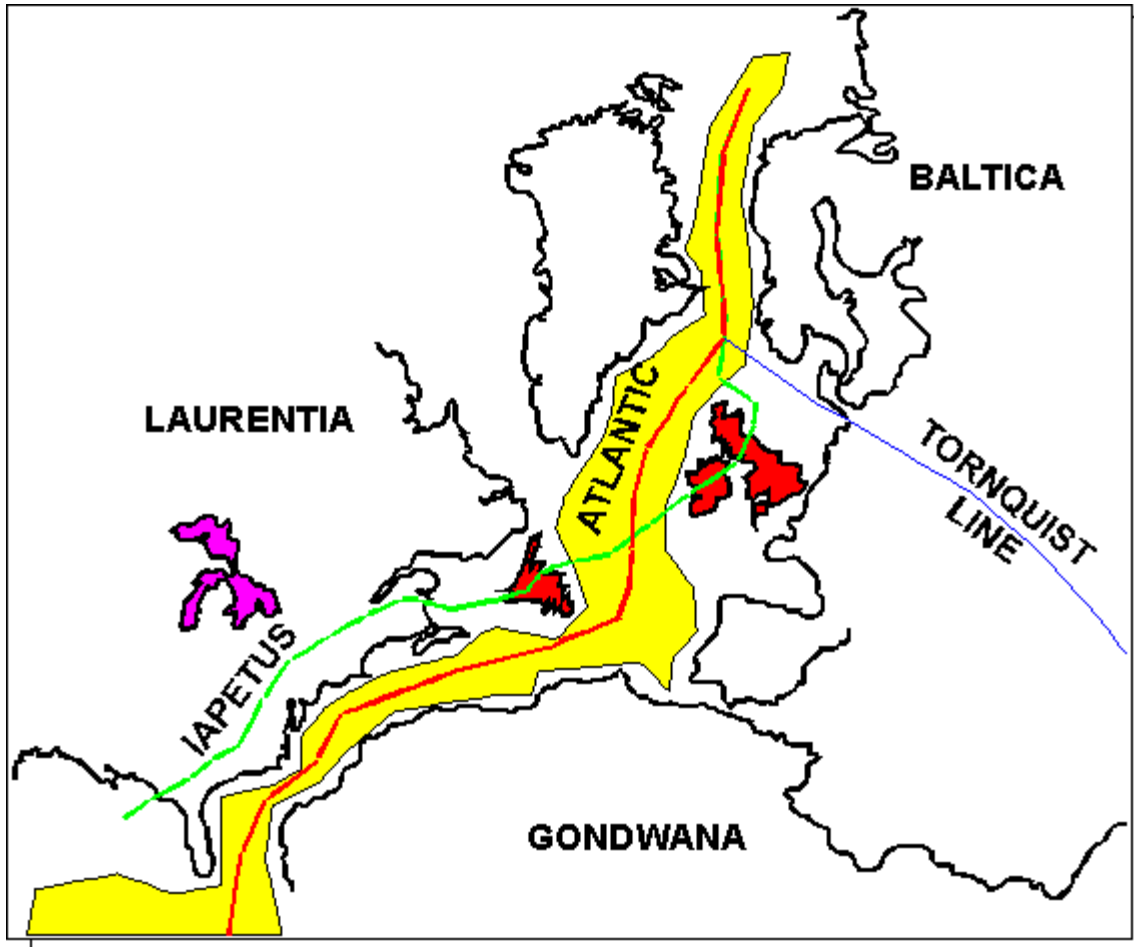
Figure 3d.



Hopson et al., 1989

**Pásmo Quatchita-Marathon – alleghanská fáze, kolize s
jihoamerickou částí Gondwany**

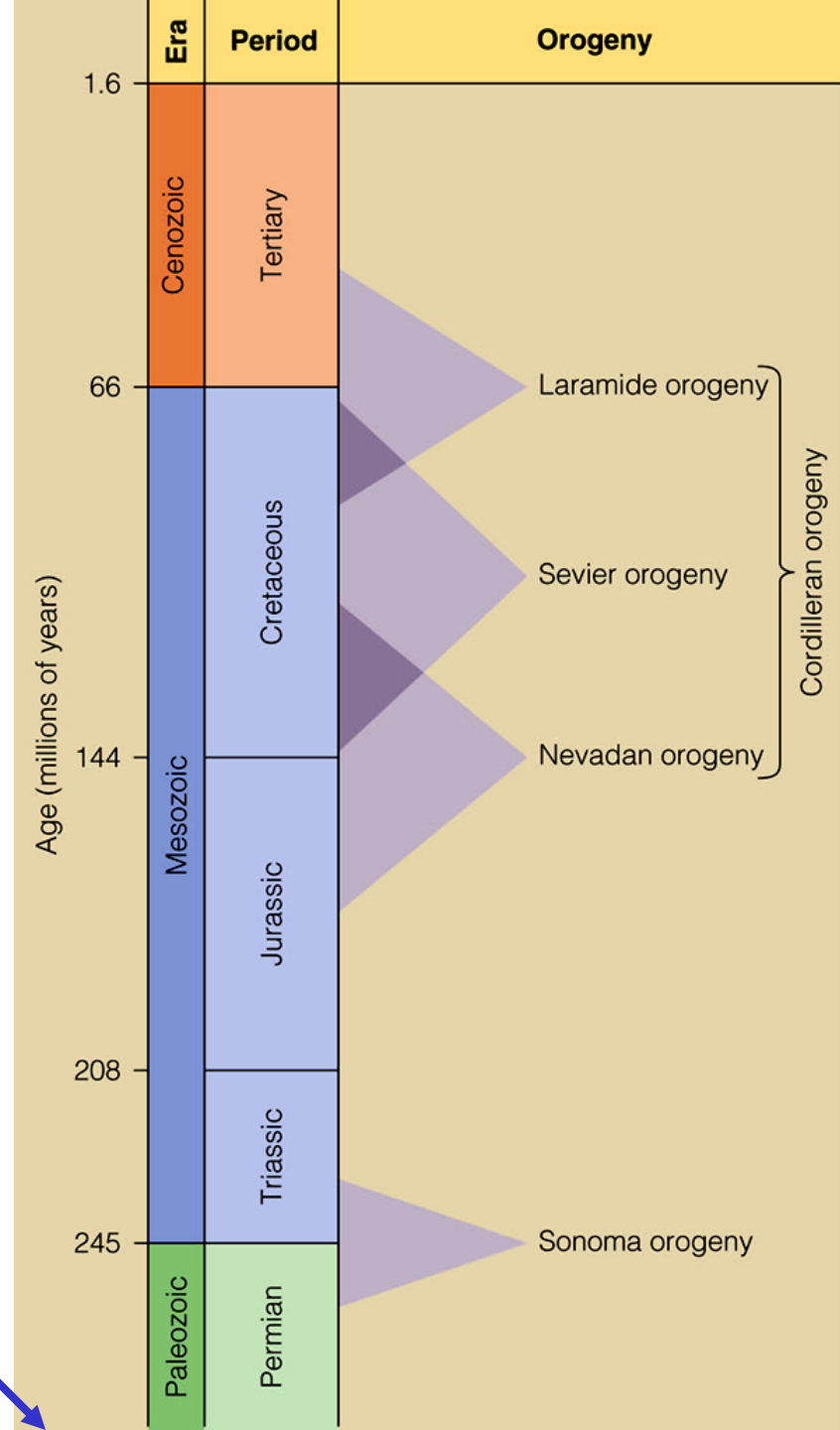
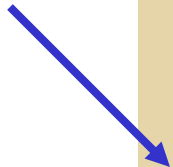




KORDILERY

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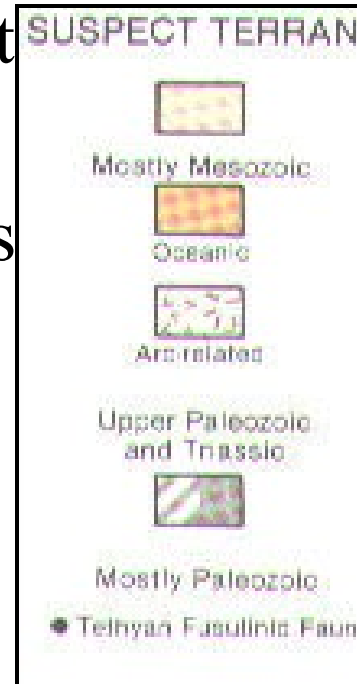
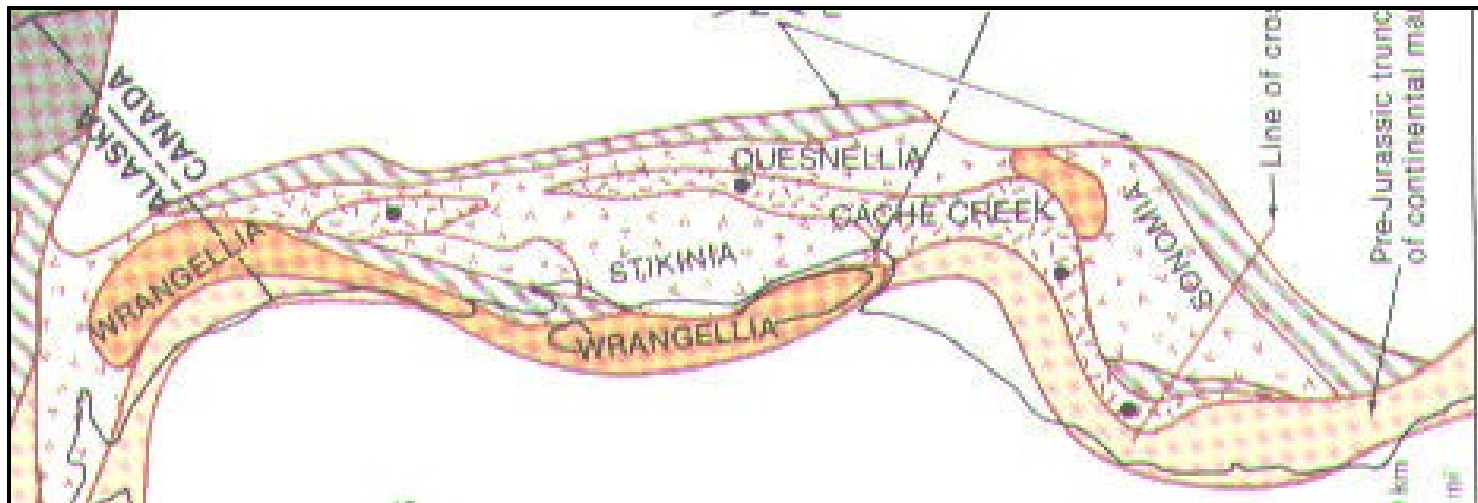
**Antler Orogeny
in Devonian**





The Western Collage

- Cordillera an collage of microplates and arcs
 - accreted during the Paleozoic and Mesozoic
 - terrains have different rock types and fossil assemblages that cannot be correlated
 - suspect terrains--fault-bounded regions that be correlated



S

Paleozoic Passive Margin

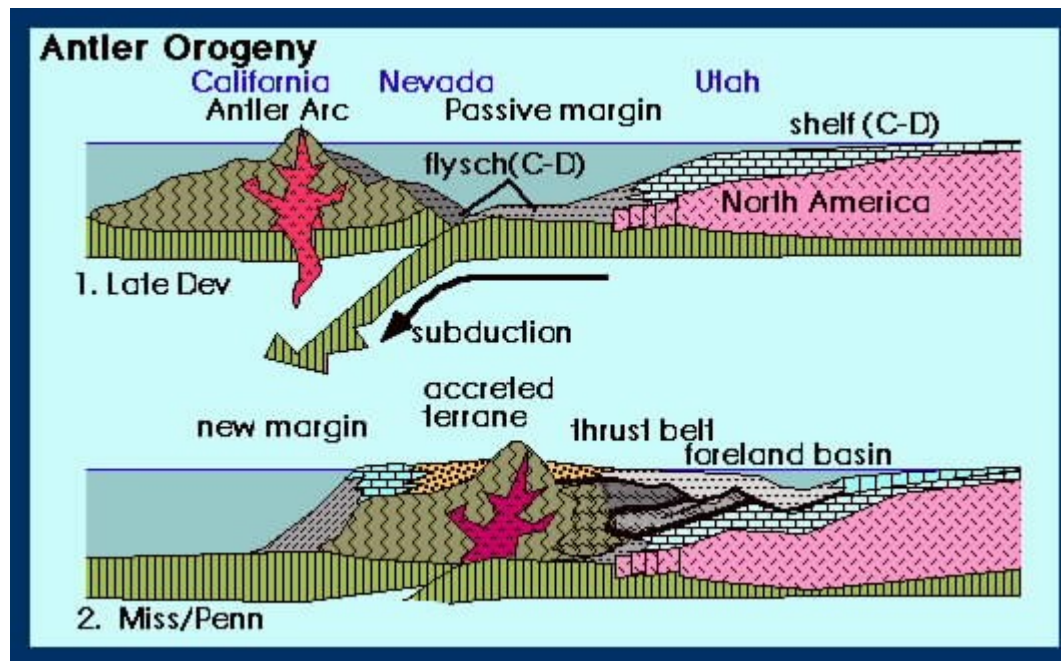
- Existed in Late Precambrian and Early Paleozoic
- Craton and cratonic basin deposits
- Miogeocline continental shelf deposits
- An arc formed in the Ordovician

Antler Orogeny

(300 (300-375 Ma) 375 Ma)

- Late Devonian - Early Mississippian
- Collision of the arc with a passive margin
- Roberts Mountain Allochthon thrust over the passive margin
- A series of foreland basins formed in eastern Nevada

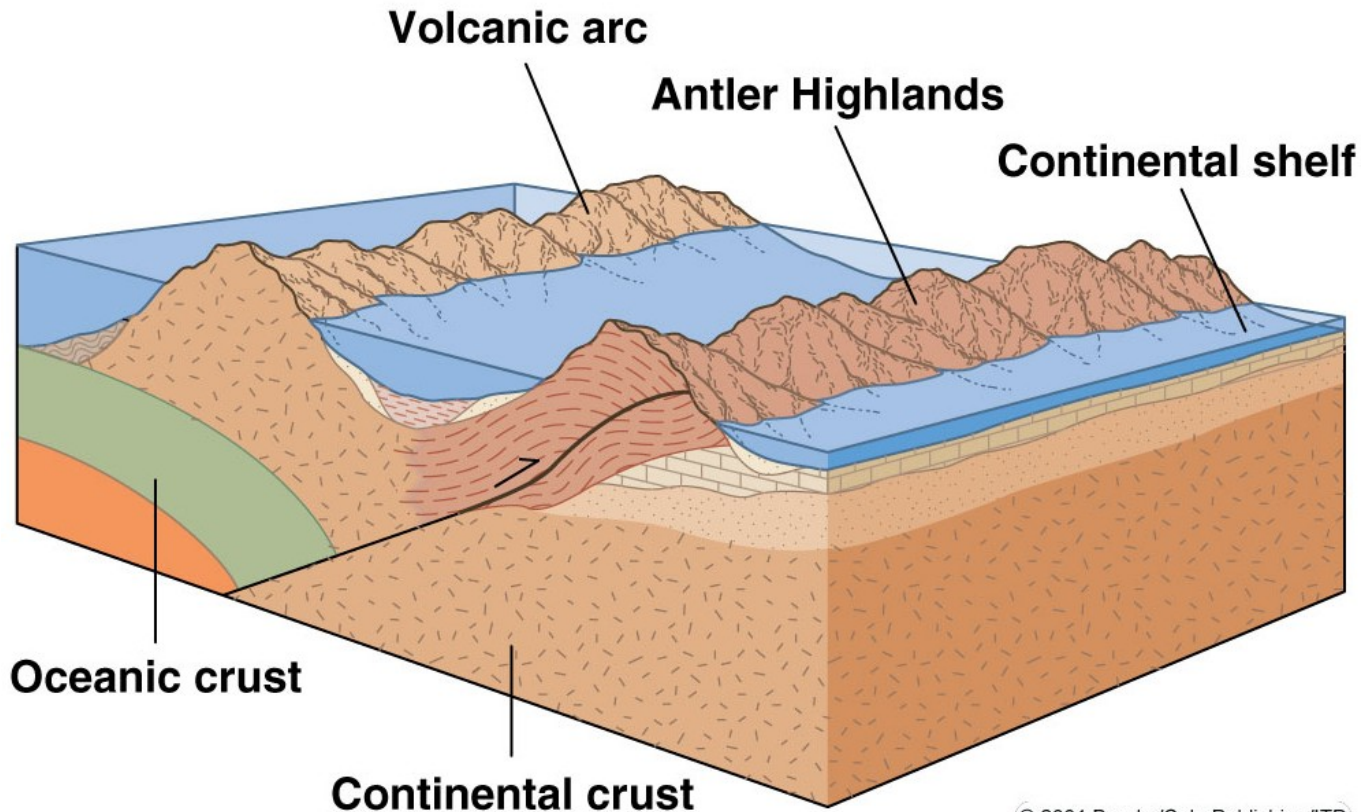
Orogeneze: antlerská, kolize klamathského ostrovní oblouku v devonu a spodním karbonu



20 Western craton (Cordilleran) was a passive margin Late Proterozoic to Early Paleozoic.

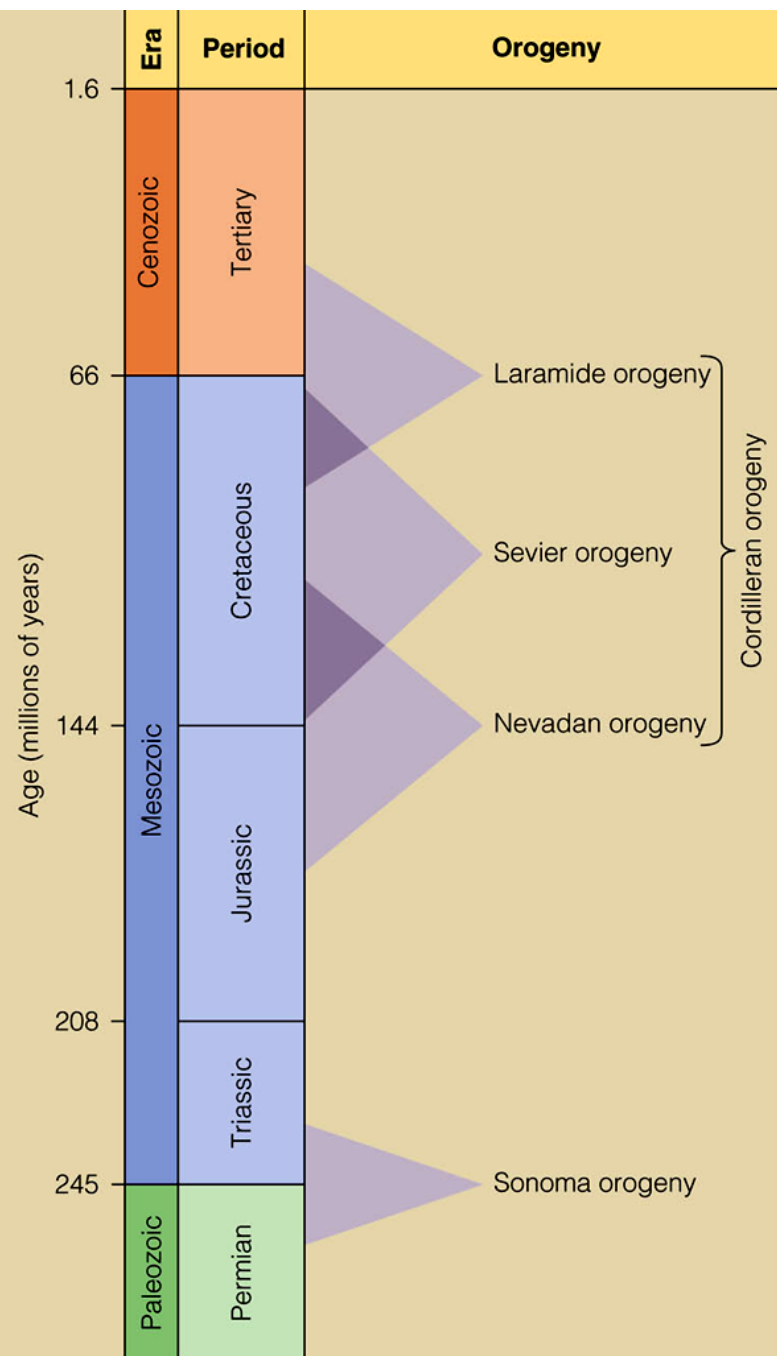
Beginning in the Middle Paleozoic - an island arc formed off the western margin of the craton.

Antler Orogeny – collision of island arc with craton – Late Devonian/Early Mississippian.



From the Antler Orogeny, the western margin remains an active margin.

Mesozoic



Cordilleran Orogeny components – Jurassic – Tertiary Periods

Nevadan – emplacements of large granite batholiths, Sierra Nevada, Idaho, S. California, Coast Ranges.

Sevier – Shallower subduction angle = igneous activity moves eastward.

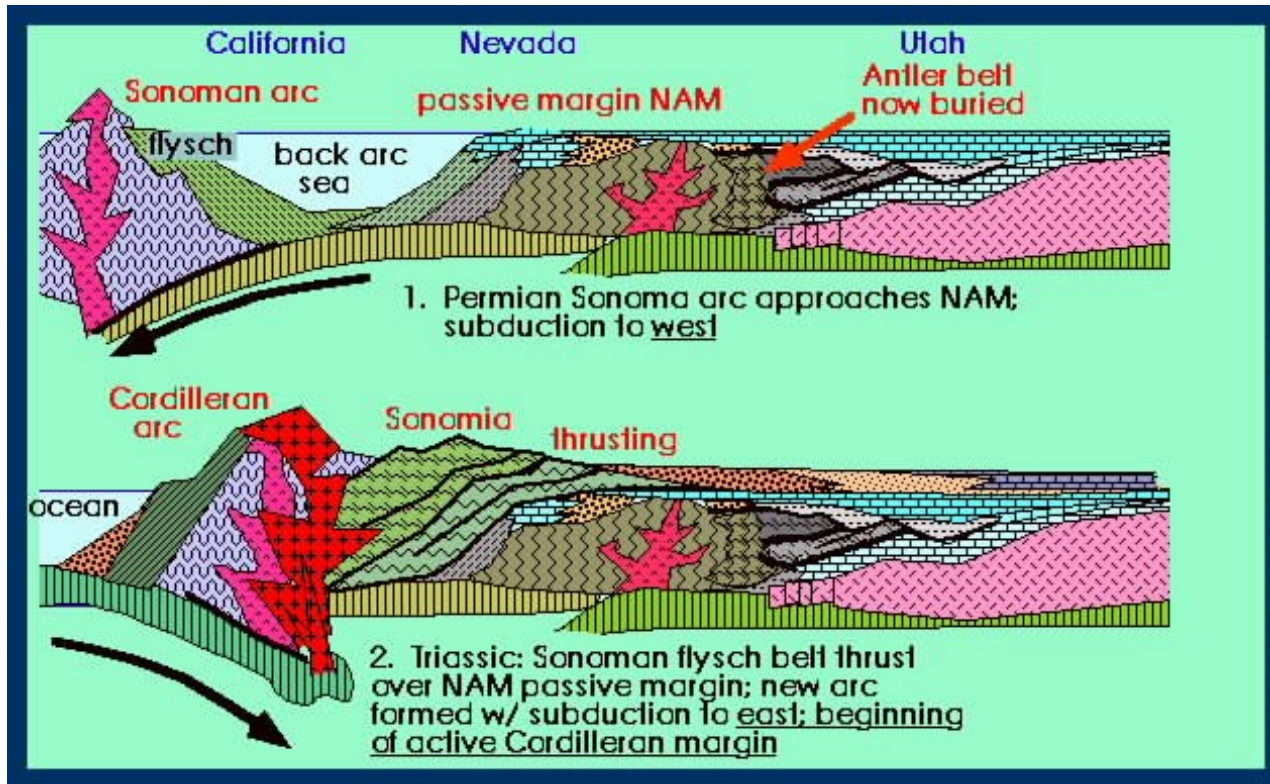
Laramide – “Rocky Mts.” – east of Sevier Orogeny – Cretaceous – Tertiary Periods

Sonoma Orogeny (200-280 Ma)

- Permo -Triassic
- Collision of Arc With a Passive Margin
- Island Arc Terrains Were Accreted
- Golconda Allochthon
 - Thrust Partly Over Roberts Mountain

Allochthon

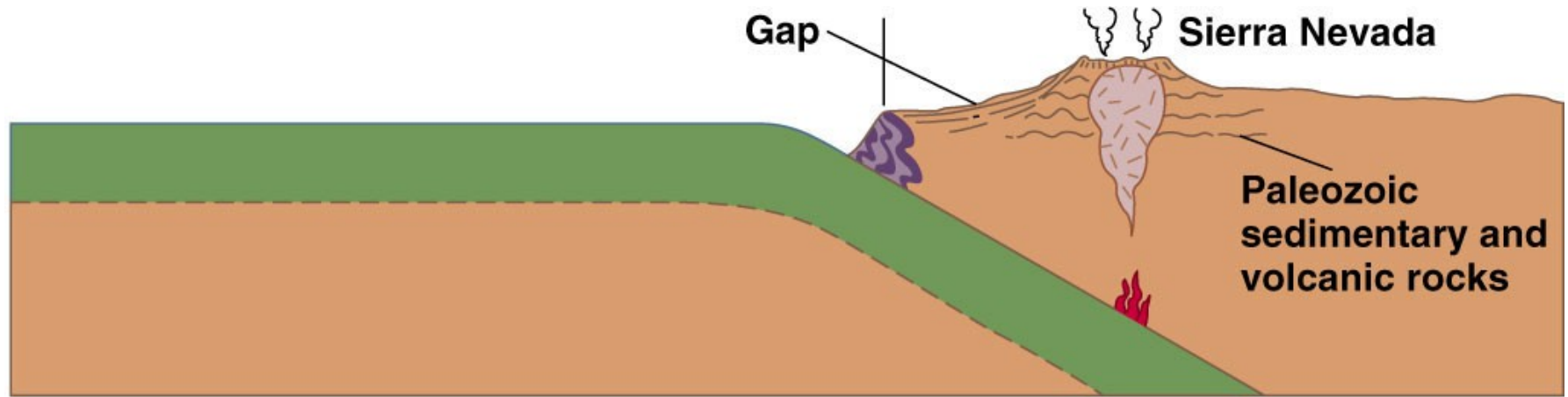
sonomská, kolize dalšího ostrovního oblouku na konci paleozoika



Nevadan Orogeny (140-150 Ma)

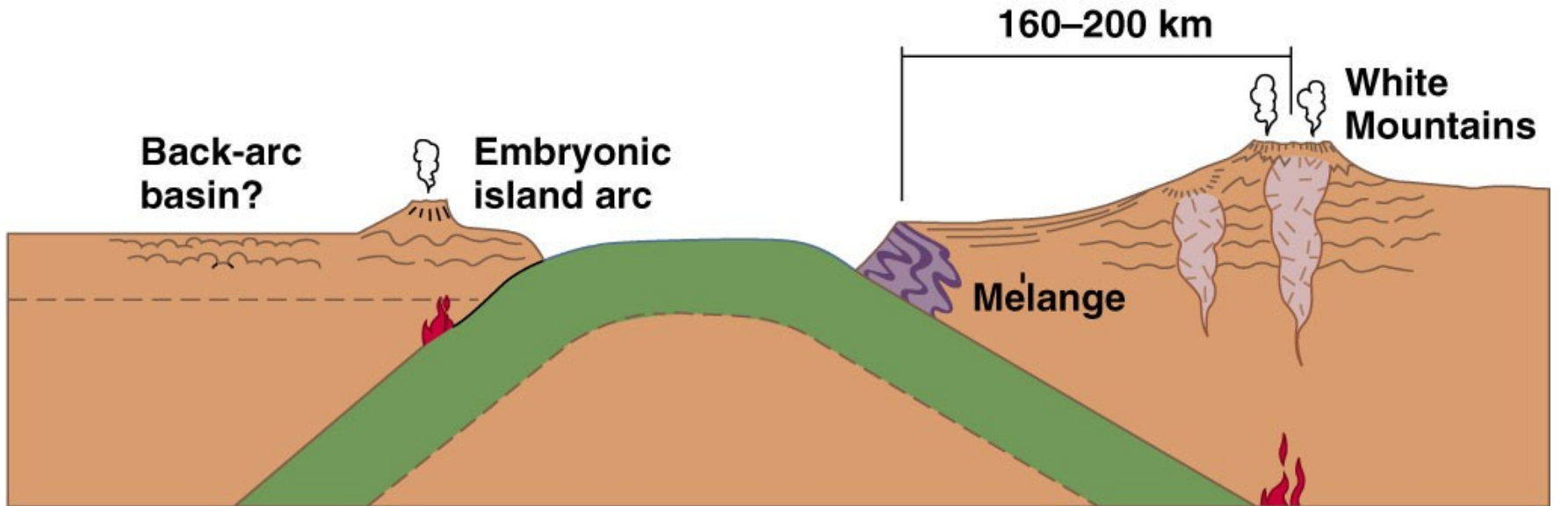
- Several Upper Jurassic Arcs Collided
- Cretaceous Franciscan Fm in the accretionary prism
- Great Valley Sequence filled an elongated forearc basin
- Sierra Nevada was the root zone of the arc

After the Sonoman Orogeny...



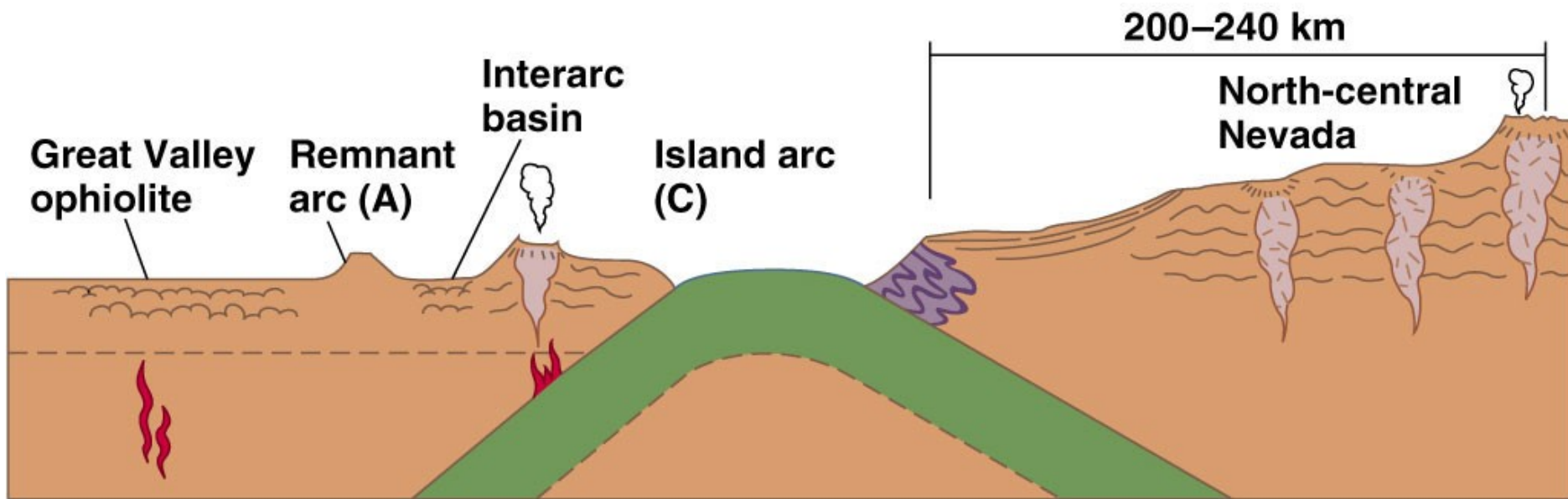
1. Late Triassic–Early Jurassic

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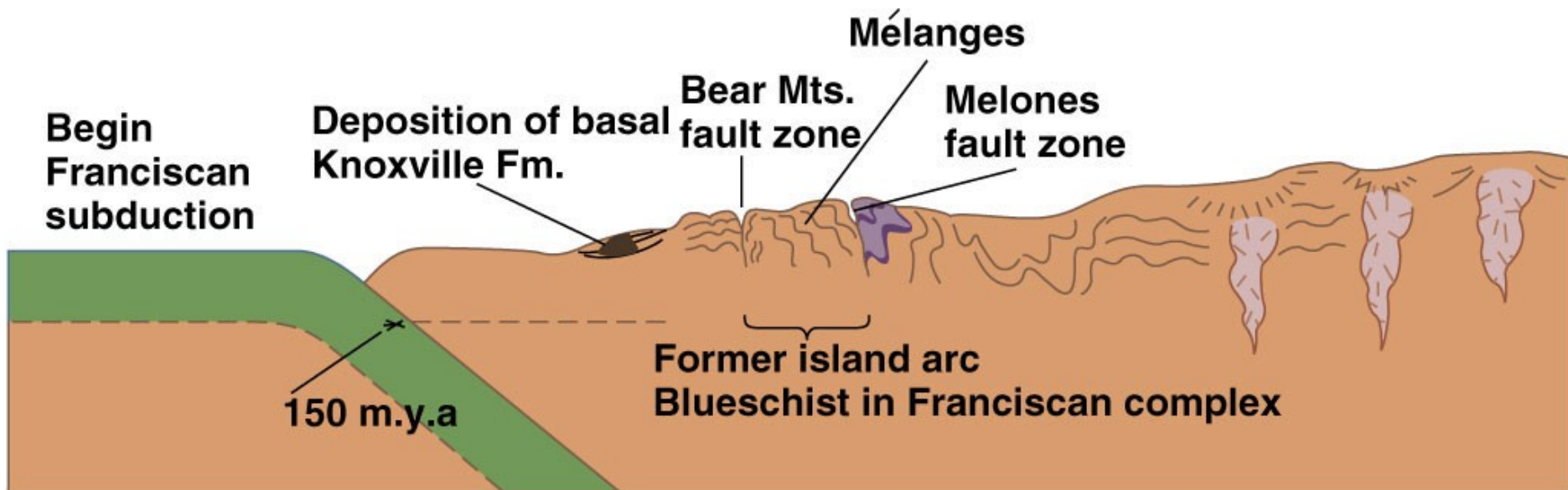
2. Middle Jurassic

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3. Early Late Jurassic

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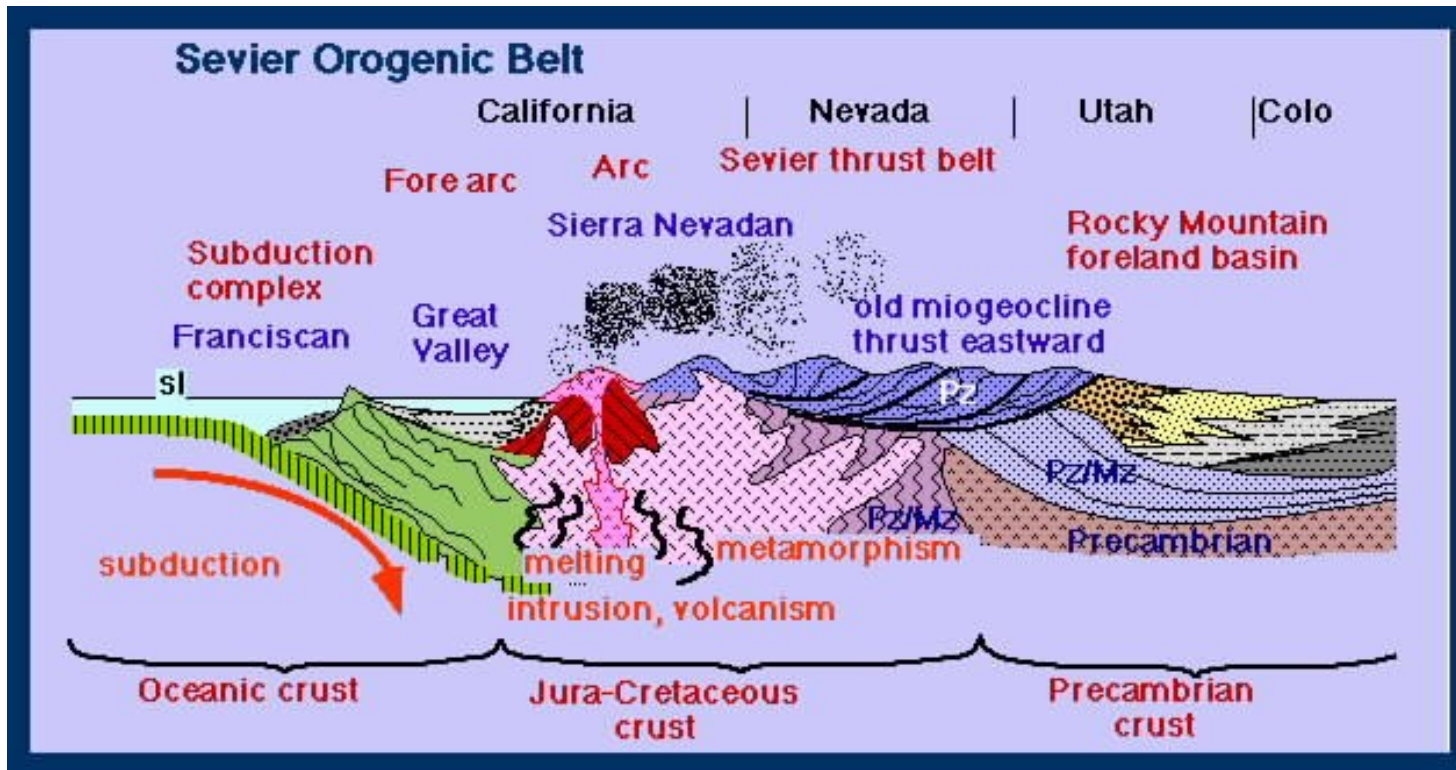
4. Late Jurassic–Nevadan orogeny

© 2001 Brooks/Cole Publishing/ITP

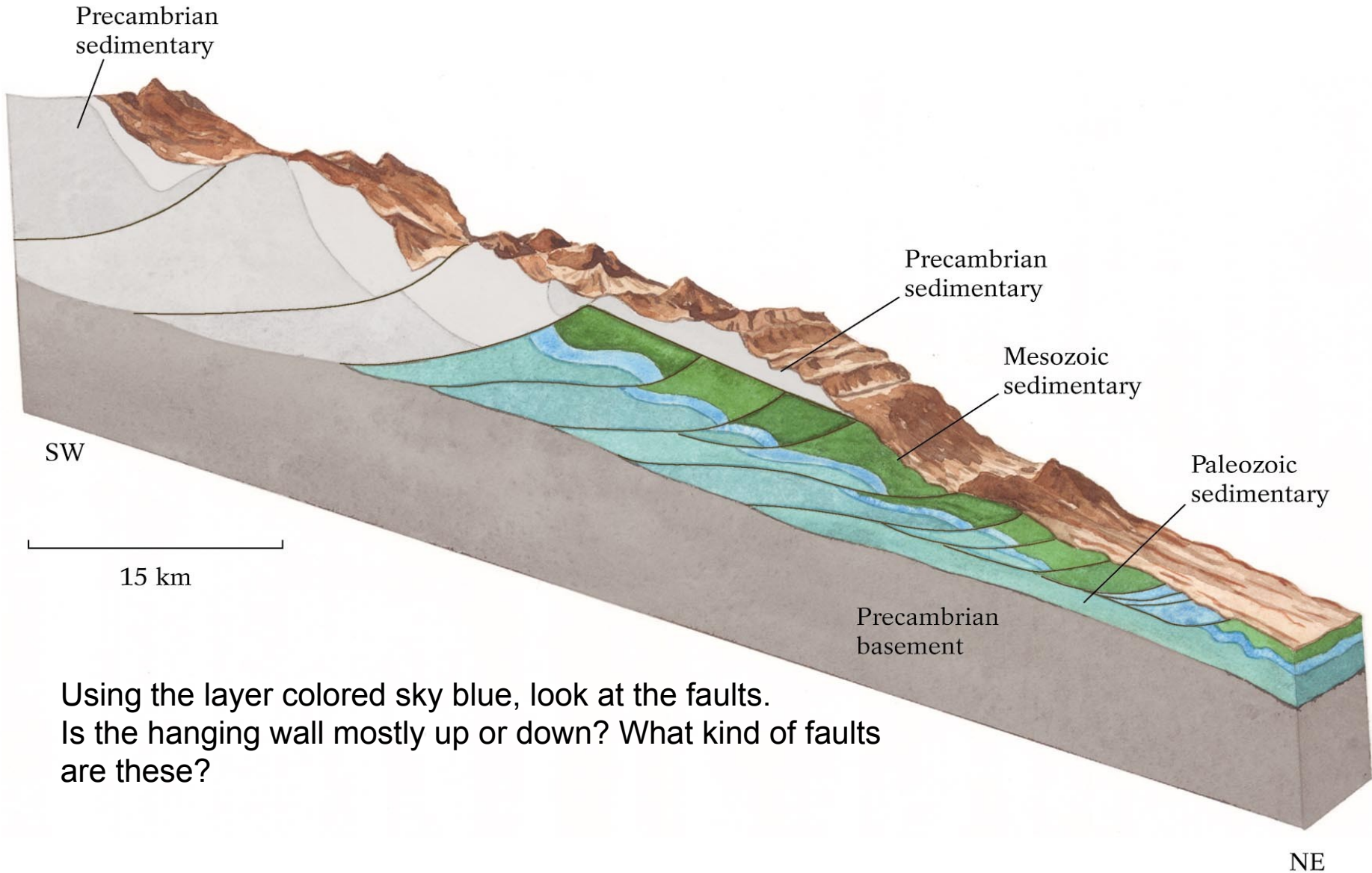
Sevier Orogen (80-130 Ma) 130 Ma)

- Fold-thrust belt behind the arc
- Eastward directed thrusts
- Prominent retro arc basins to the East
- Late Jurassic to Late Cretaceous
- Batholithic intrusions
- Great Valley Sequence
- Franciscan Formation

In general the Sevier orogeny defines a more western compressional event that took advantage of weak bedding planes in overlying [Paleozoic](#) and [Mesozoic](#) sedimentary rock. advantage of As the crust was shortened, pressure was transferred eastward along the weak sedimentary layers, producing “[thin-skinned](#)” [thrust faults](#) that generally get younger to the east. In contrast, the Laramide orogeny produced “basement-cored” uplifts that often took preexisting faults that formed during [rifting](#) in the Late Precambrian during the breakup of the supercontinent [Rodinia](#) or during the [Ancestral Rocky Mountains](#) orogeny.[1]



Sevier thin-skinned deformation



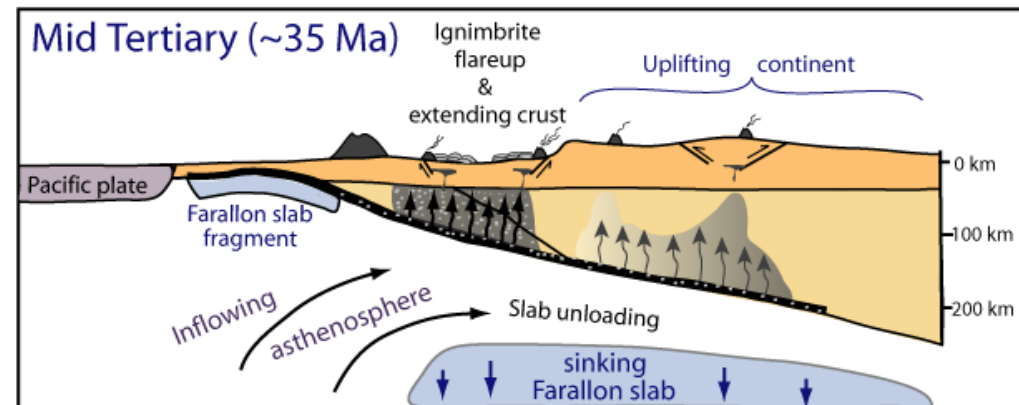
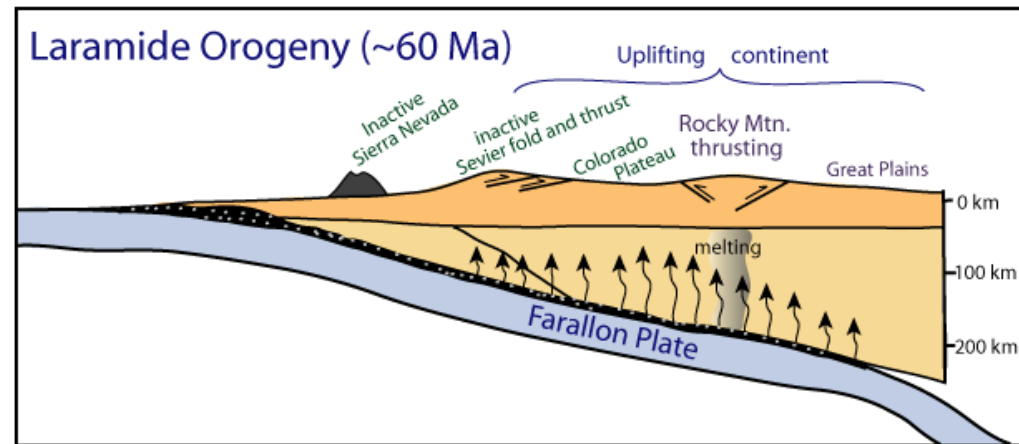
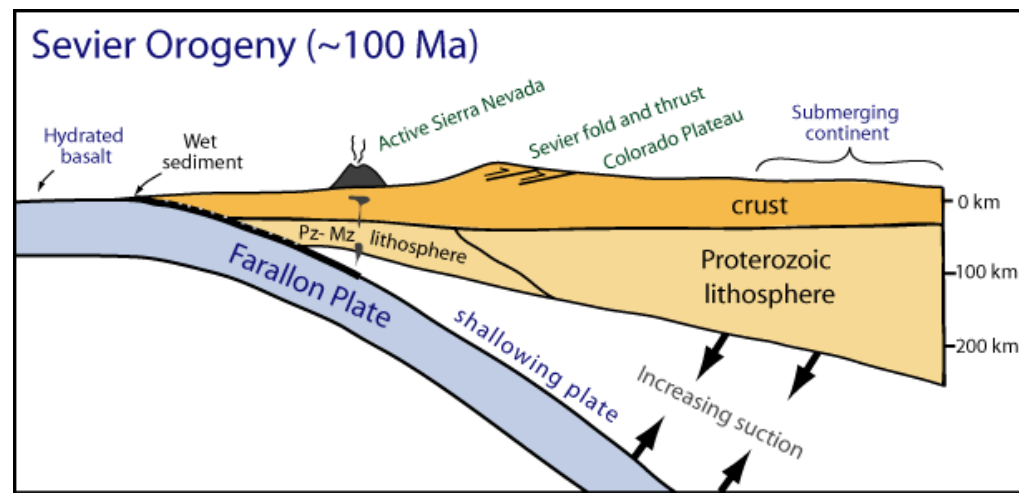
Using the layer colored sky blue, look at the faults.
Is the hanging wall mostly up or down? What kind of faults
are these?

Laramide Orogeny (50- 80 Ma)

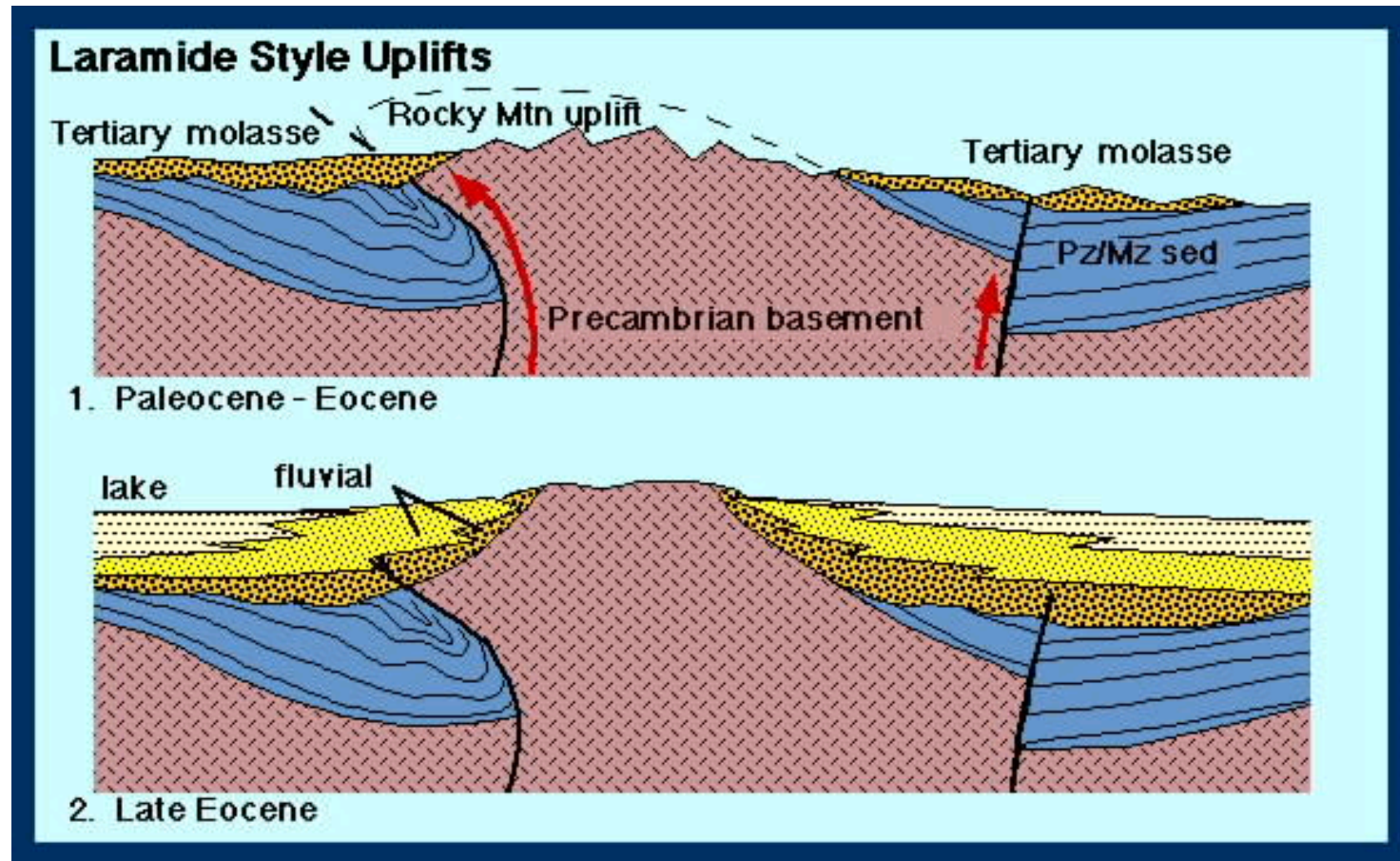
- Late Cretaceous - Early Eocene
- Deformation shifted eastward following magmatism
- Westward directed thrusts
- Formation of major mineral belts

The Laramide orogeny was caused by subduction of a plate at a shallow angle

**Asthenospheric upwelling at 54-50 Ma
(older than in the south)**

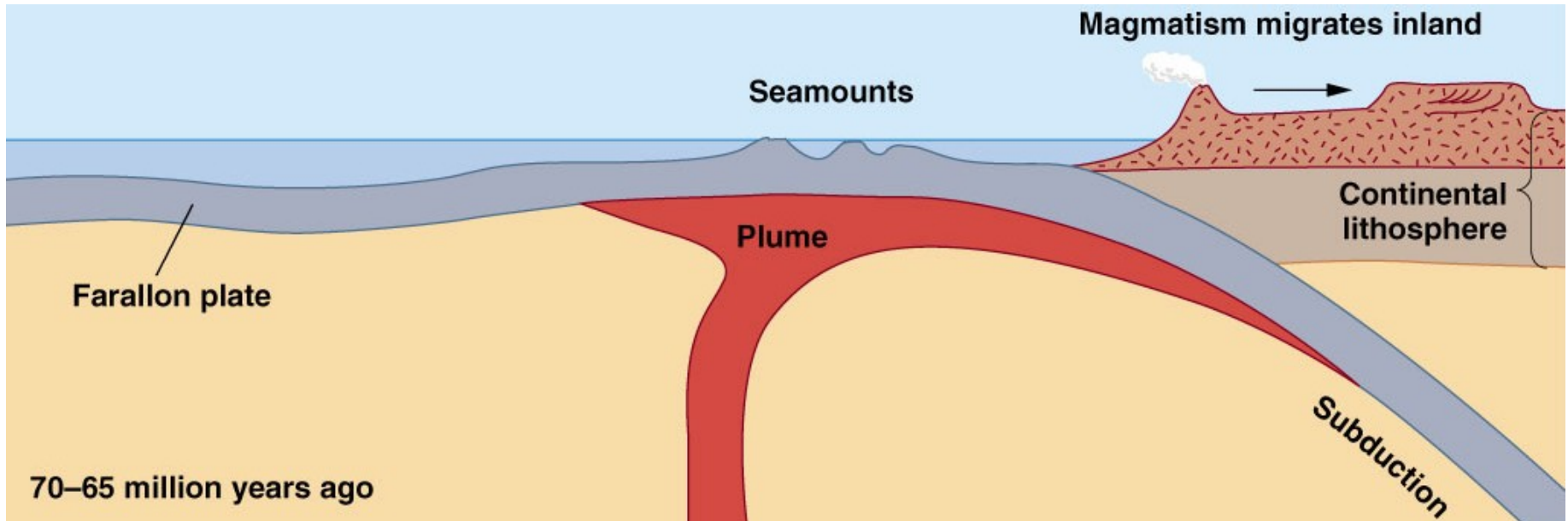


Laramide Involves Basement!



Laramide orogeny

- took place as the Farallon plate, buoyed up by a mantle plume subducted beneath North America at a decreasing angle and igneous activity shifted inland

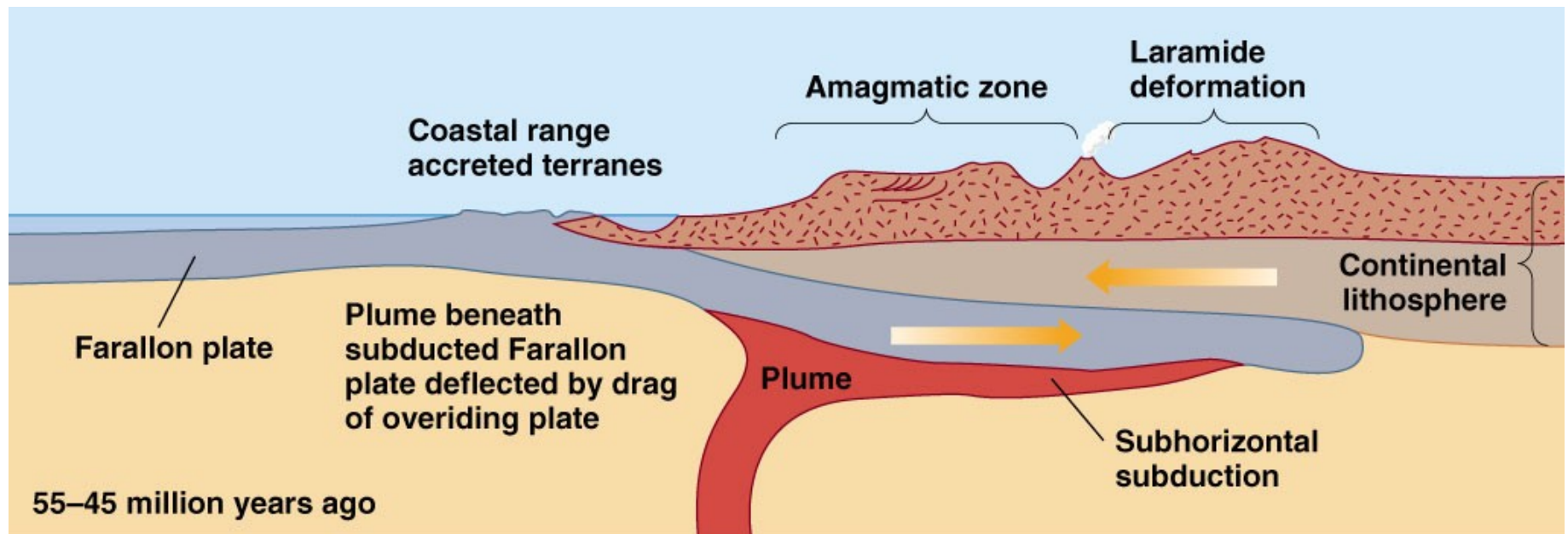


Change to Shallow Subduction

- By Early Tertiary time,
 - the westward-moving North American plate
 - had overridden the part of the Farallon plate,
 - above the head of the mantle plume
- The lithosphere
 - immediately above this plume
 - was buoyed up,
 - accounting for a change
 - from steep to shallow subduction

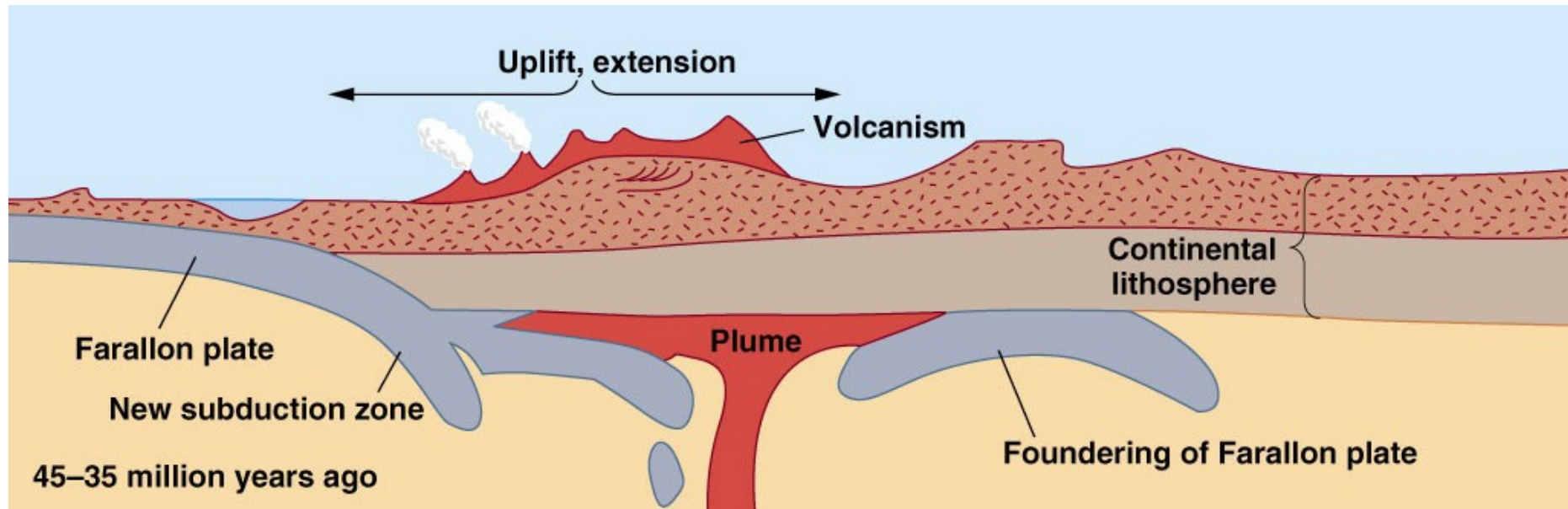
Igneous Activity Ceased

- With nearly horizontal subduction, igneous activity ceased and the continental crust was deformed mostly by vertical uplift



Renewed Igneous Activity

- Disruption of the oceanic plate by the mantle plume marked the onset of renewed igneous activity

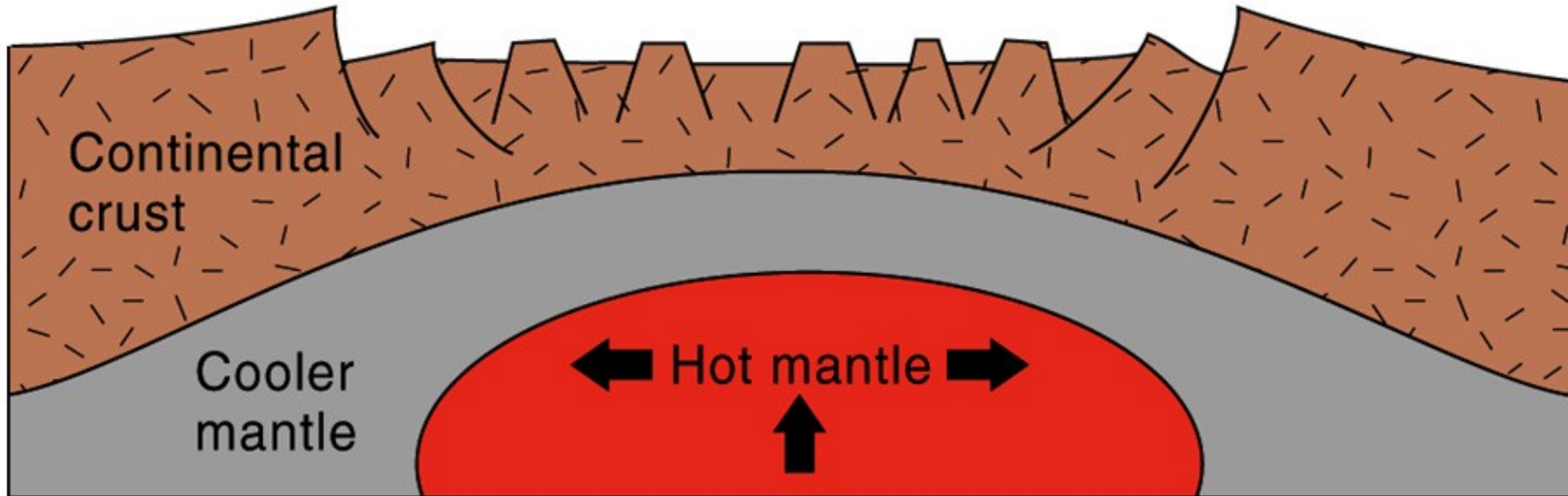


Basin and Range orogeny

The mechanism responsible for formation of the North American Basin and Range province

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← Extension →



Basin and Range



THE CANADIAN CORDILLERA: Geology and Tectonic Evolution

Continued from Page 18

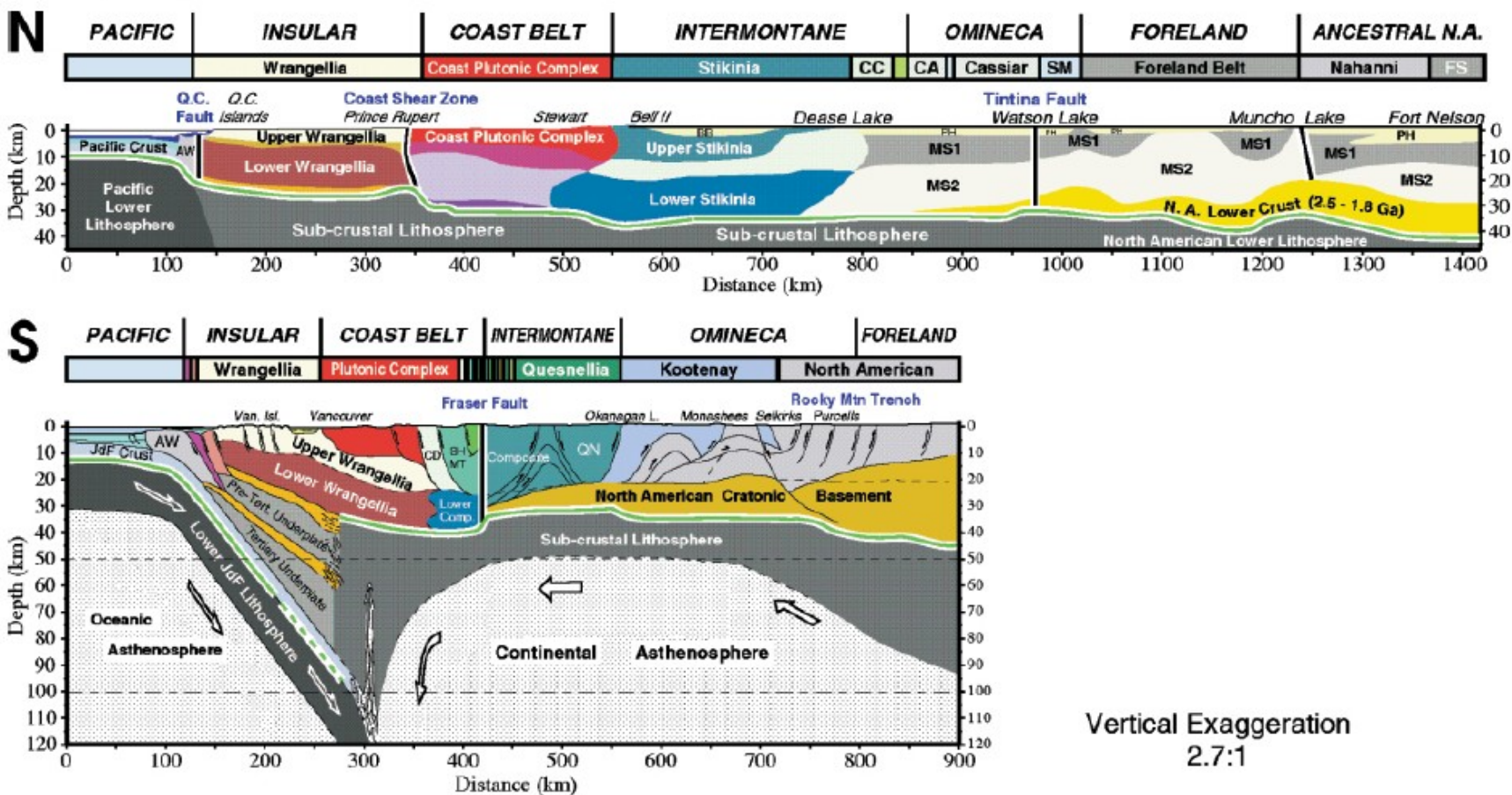
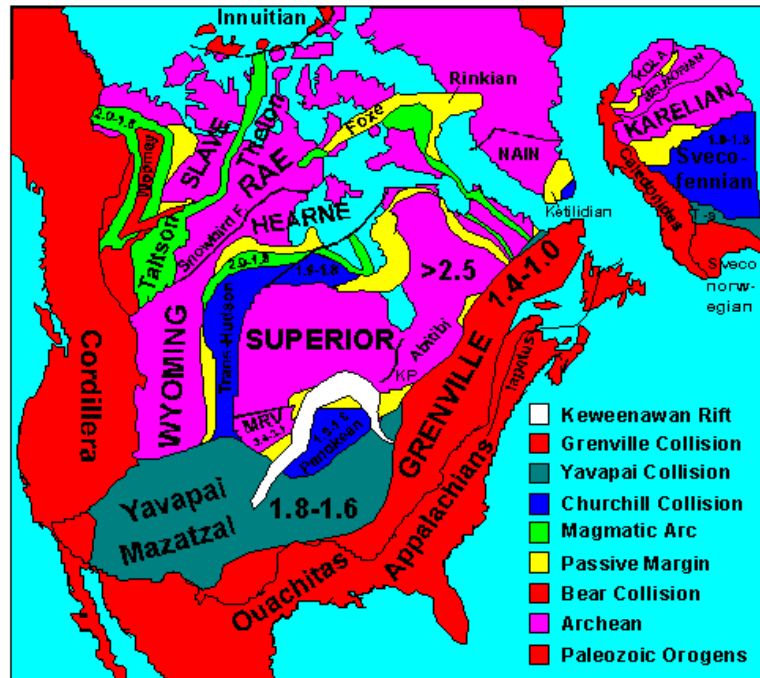


Figure 2. Comparison of interpreted simplified lithospheric structures along the northern (N) and southern (S) Cordilleran Lithoprobe transects, whose locations are shown by the red lines in Figure 1. In both profiles, the heavy green line is the crust-mantle boundary (Moho). AW accreted wedge; AX Alexander terrane; BB Bowser Basin; CA Cassiar terrane; CC Cache Creek terrane; CD Cadwallader terrane; FS, Fort Simpson (a Precambrian terrane in the craton); KO Kootenay terrane; MS1 undivided Precambrian (1200-800 Ma) metasedimentary rocks; MS2 undivided Precambrian (1800-1200 Ma) metasedimentary rocks; MT-SH undivided Methow and Siuksan terranes; QN Quesnel terrane; SM Slide Mountain terrane; ST Stikine terrane; WR Wrangellia. Most terrane descriptions are in Table 2. Figure modified by P.T.S. Hammer from Clowes and Hammer, 2000.

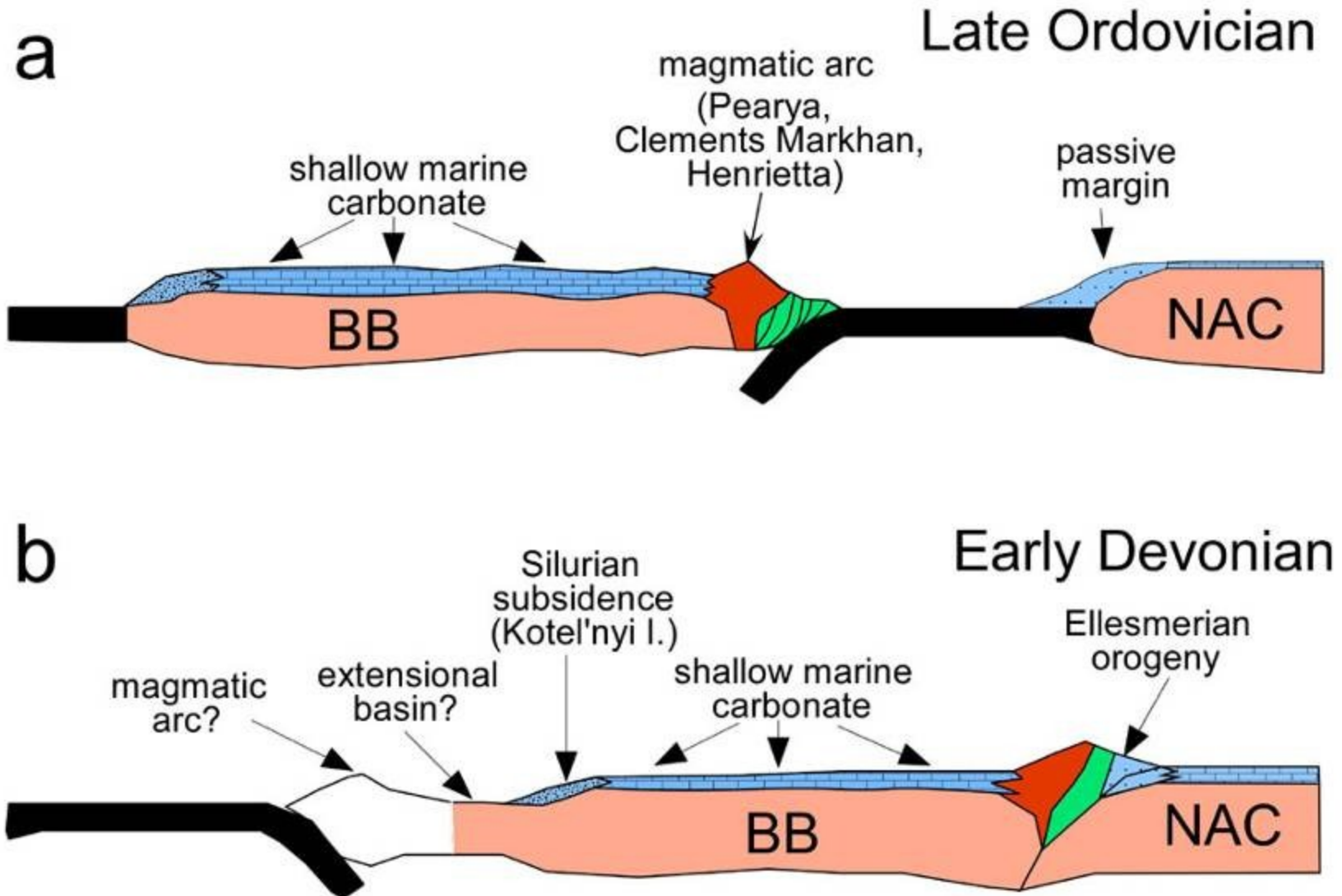
Franklinsko – inuitská orogeneze

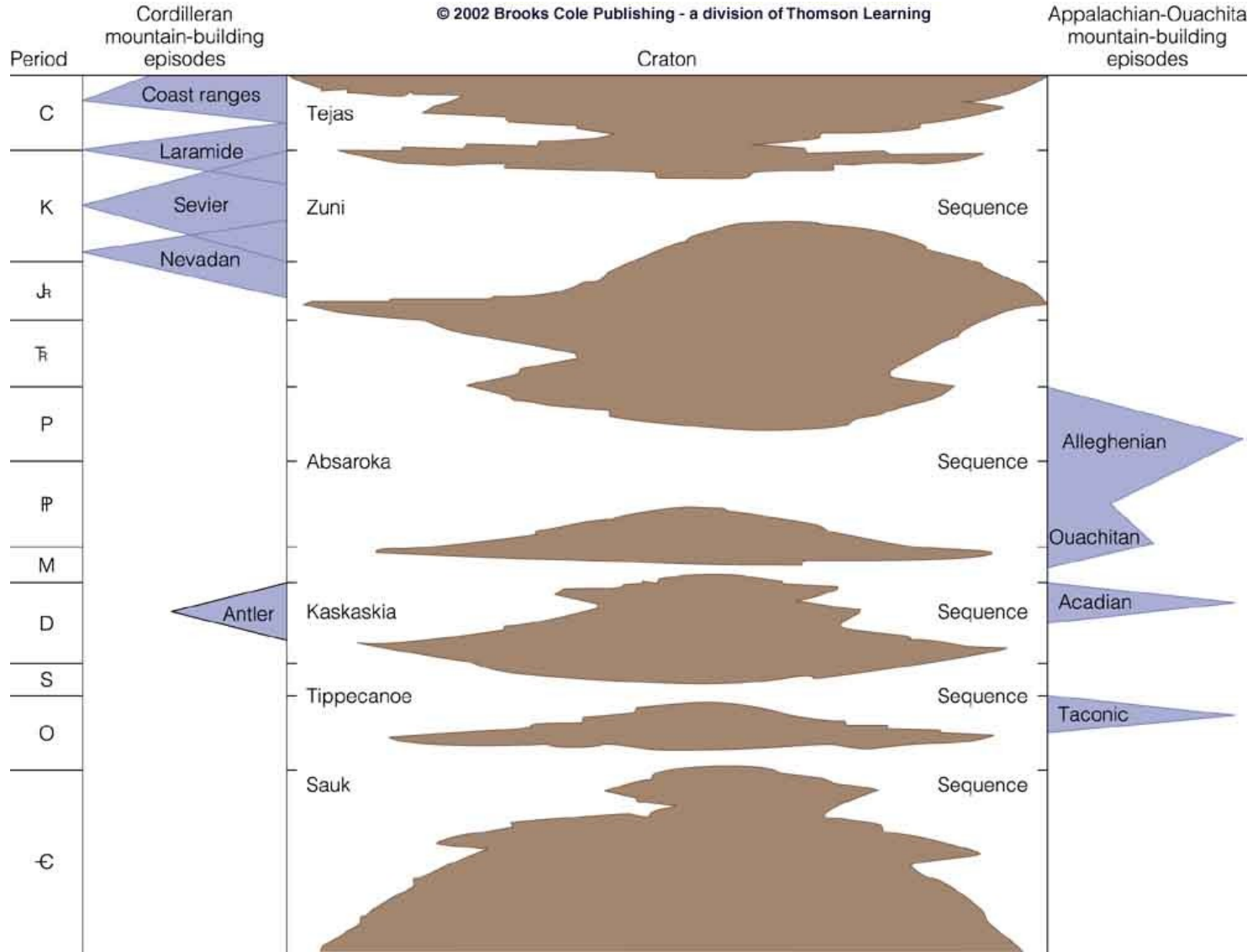
The **Franklinian orogeny**, in the northwestern Canada (Plafker & Berg, 1994), could be a result of collision of the Verkhoyanskian part of Siberia with the North Slope-Chukotkan part of Laurentia. According to Okulitch (1998), the suturing in the Canadian Islands occurred during Ordovician-Silurian time.

The **Innuitian Orogeny** started in the earliest Middle Devonian and may be linked to plate movements that also emplaced an exotic Bennett-Barrovia terrane on the northern edge of the region.



The late Silurian-early Devonian collision of the Bennett-Barrovia block





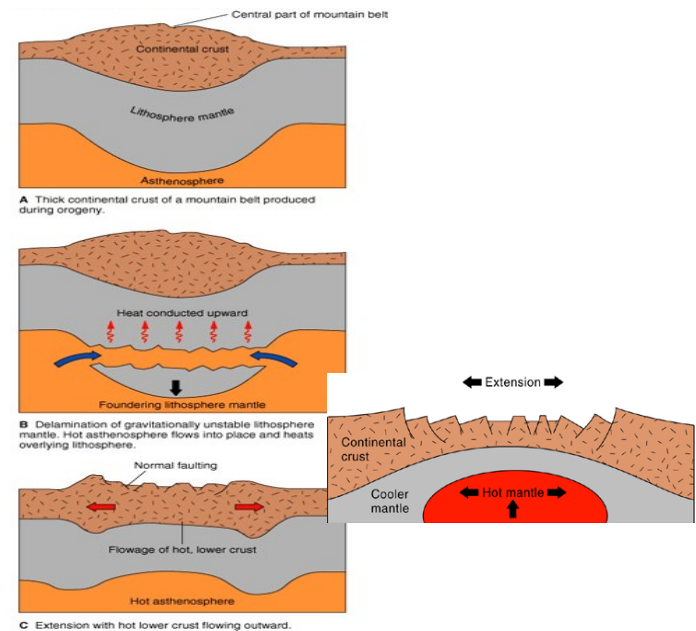


Post Laramide Events

- Cenozoic extension
- Basin and Range formation
- Cenozoic magmatism
- Widespread volcanism

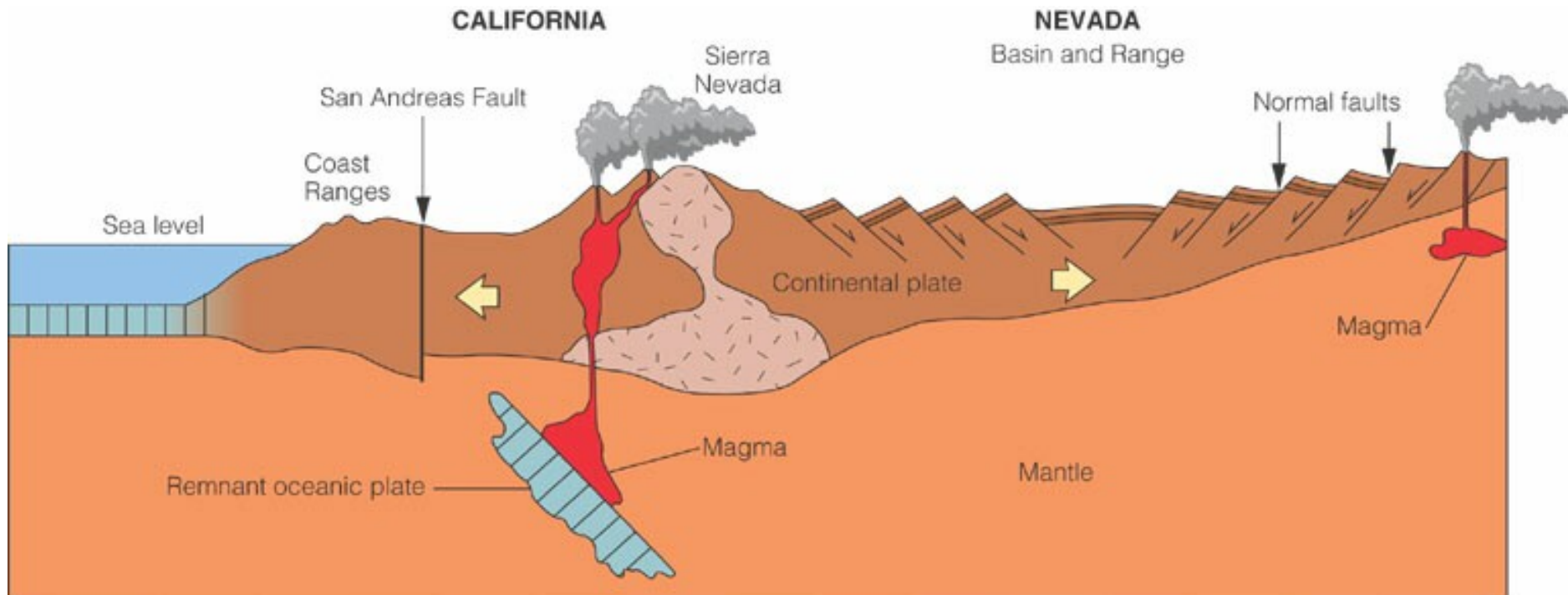
Evolution of Mountain Belts

- Basin-and-Range province of western North America may be the result of *delamination*
 - Overthickened mantle lithosphere beneath old orogenic mountain belt may break off and sink (*founder*) into asthenosphere
 - Resulting inflow of hot *asthenosphere* can stretch and thin overlying crust, producing normal faults under tension



Basin and Range Province

- Generalized cross section of the Basin and Range Province
 - ranges are bounded by faults

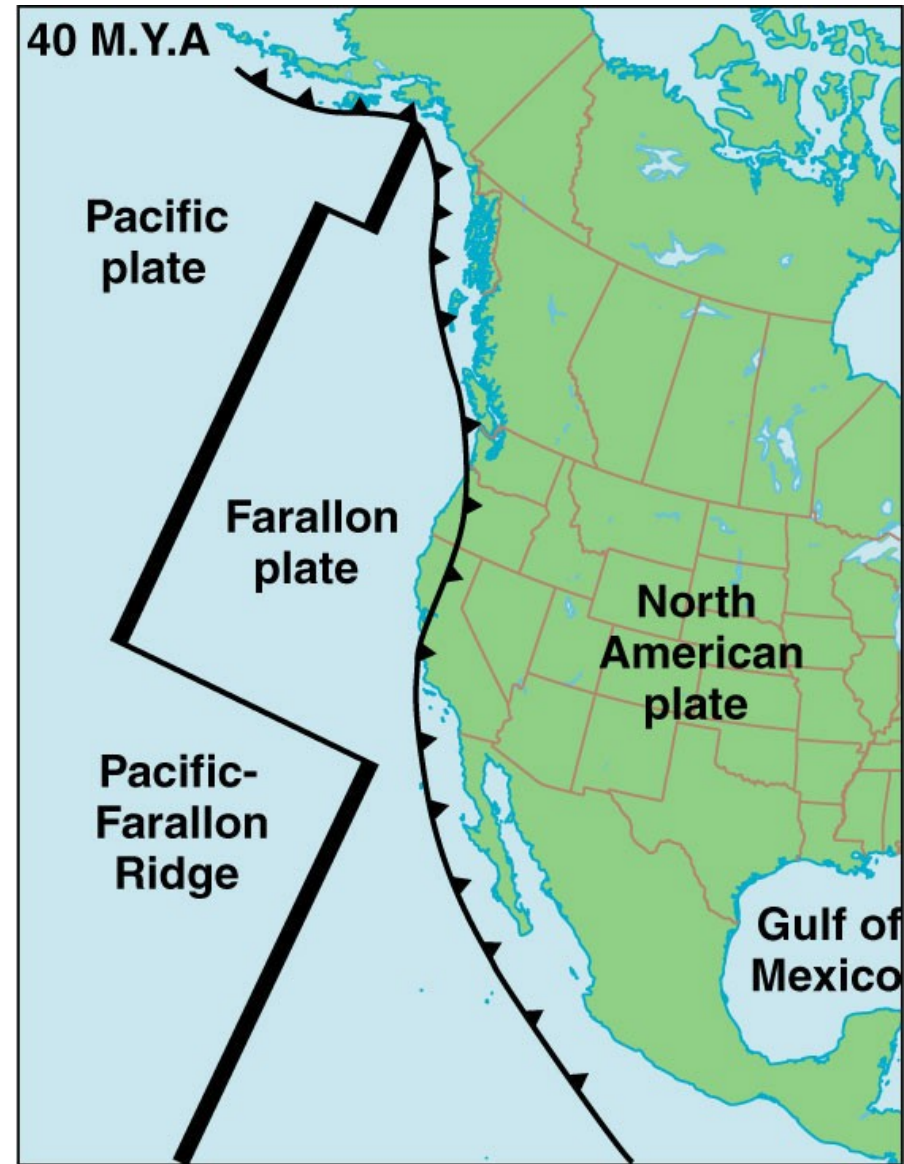


Cordillera Evolved

- After Laramide deformation, Cordillera continued to evolve with large-scale block-faulting, extensive volcanism and vertical uplift and deep erosion
 - Basin and Range
- During about the first half of the Cenozoic Era, a subduction zone was present along the entire western margin of the Cordillera, but now most of it is a transform plate boundary

Pacific Coast

- Before the Eocene,
 - the entire Pacific Coast was a convergent plate boundary
 - Farallon plate was consumed at a subduction zone
 - stretched from Mexico to Alaska



(a)

Change from Subduction

- As the North American Plate overrode the Pacific–Farallon Ridge, its margin became transform faults
 - the San Andreas
 - and the Queen Charlotte – alternating with subduction zones



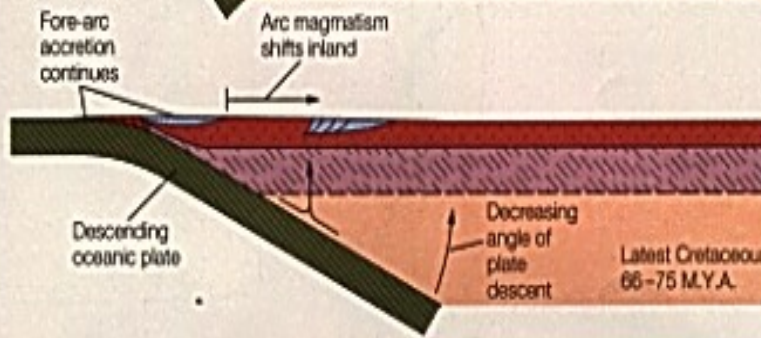
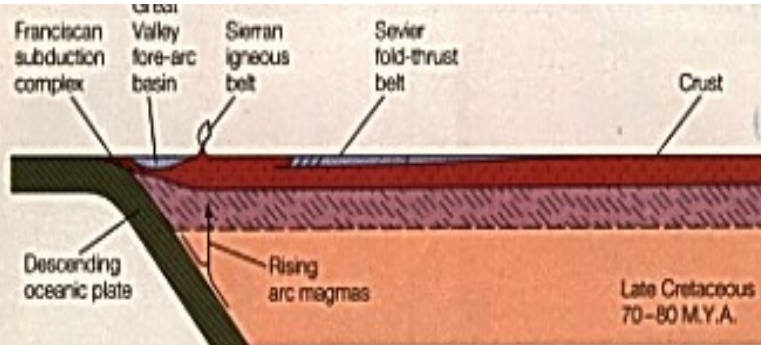
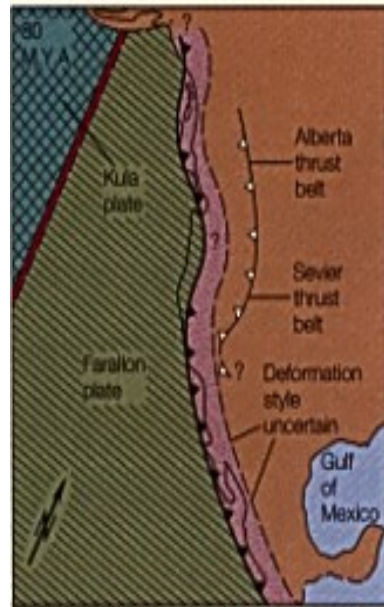
(b)

Extending the San Andreas Fault

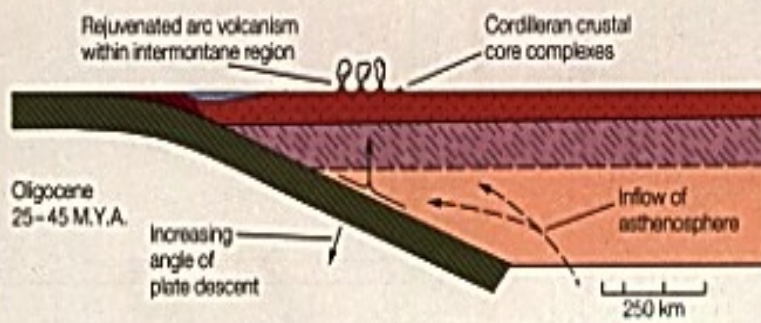
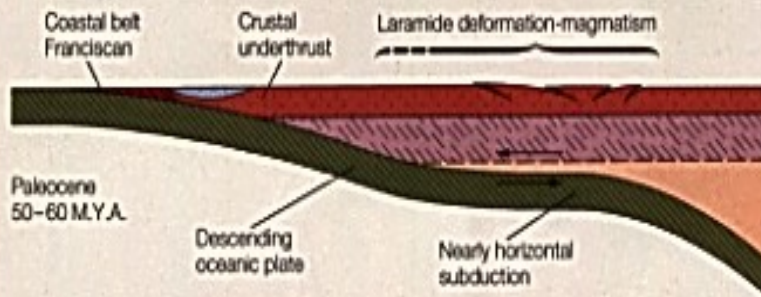
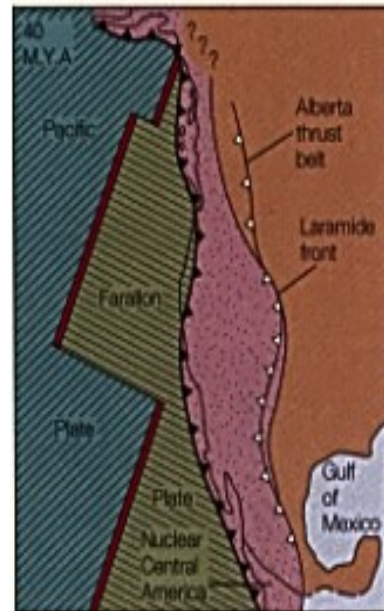
- Further overriding of the ridge extended the San Andreas Fault and diminished the size of the Farallon–Plate remnants
- Now only two small remnants of the Farallon plate exist
 - the Juan de Fuca and Cocos plates



Late Cretaceous



Eocene



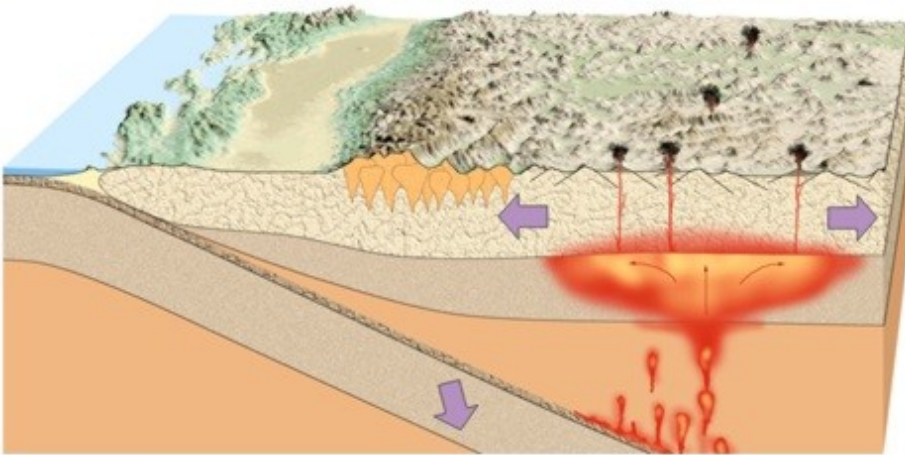
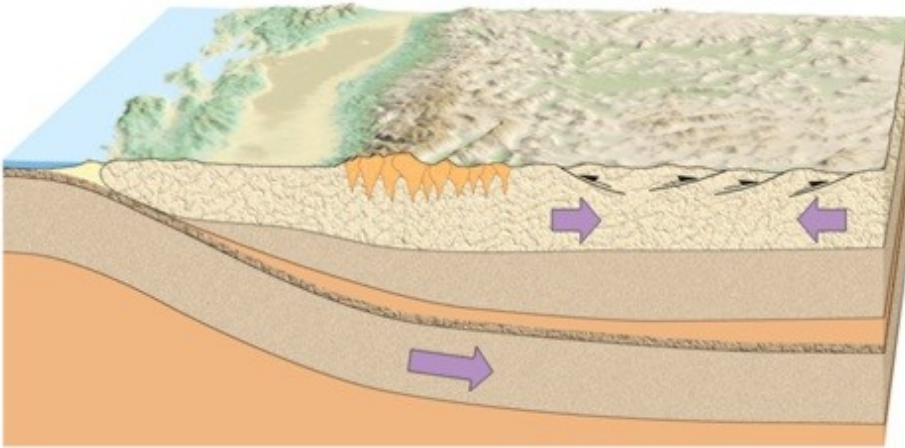
The Tetons (a range), as seen looking west across Jackson Hole, WY (a basin)



Colorado Plateau

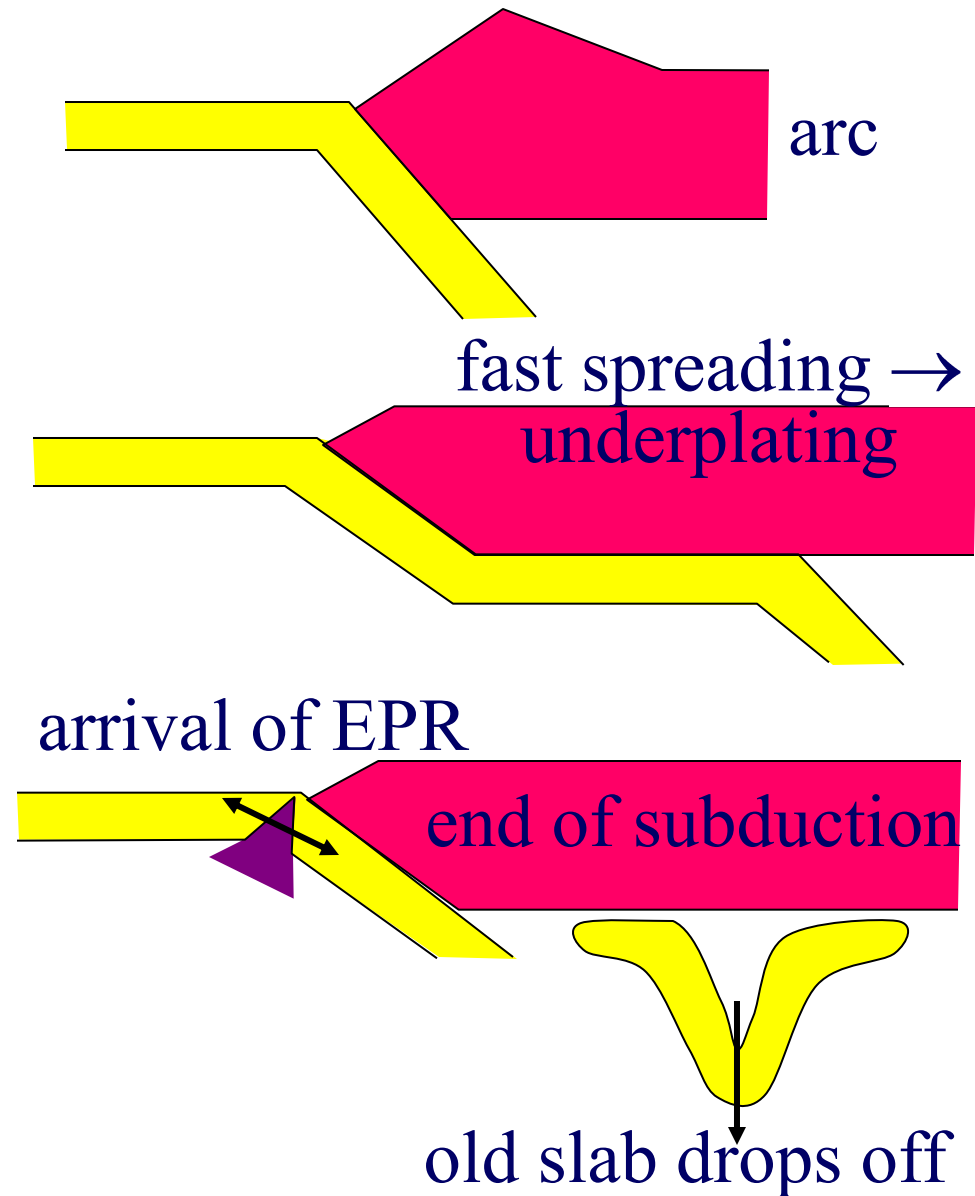


Bottom: Sinking of this oceanic slab allowed for the upwelling of hot material from the asthenosphere. The buoyancy of the warm material caused upwarping and tensional fracturing in the crust above. This event was associated with volcanism and east-west extension of the crust by nearly 150 kilometers.

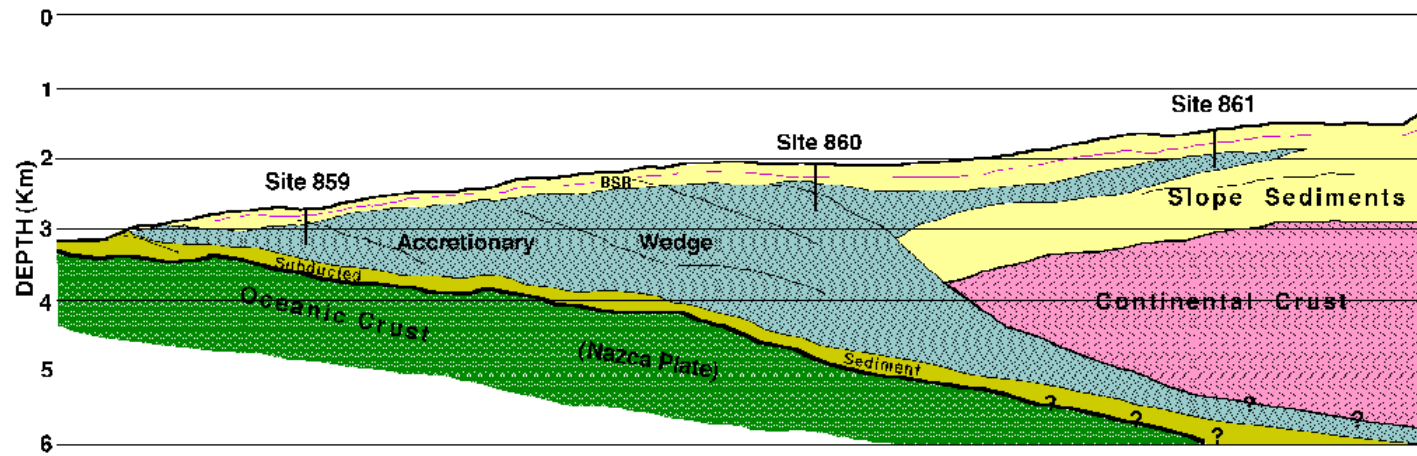
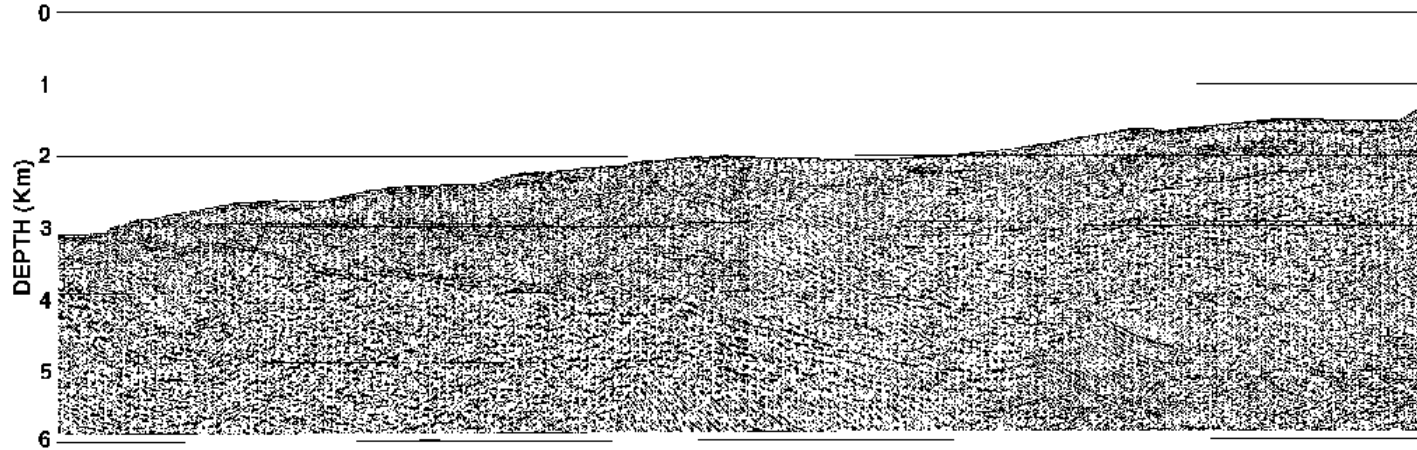


Basin and Range

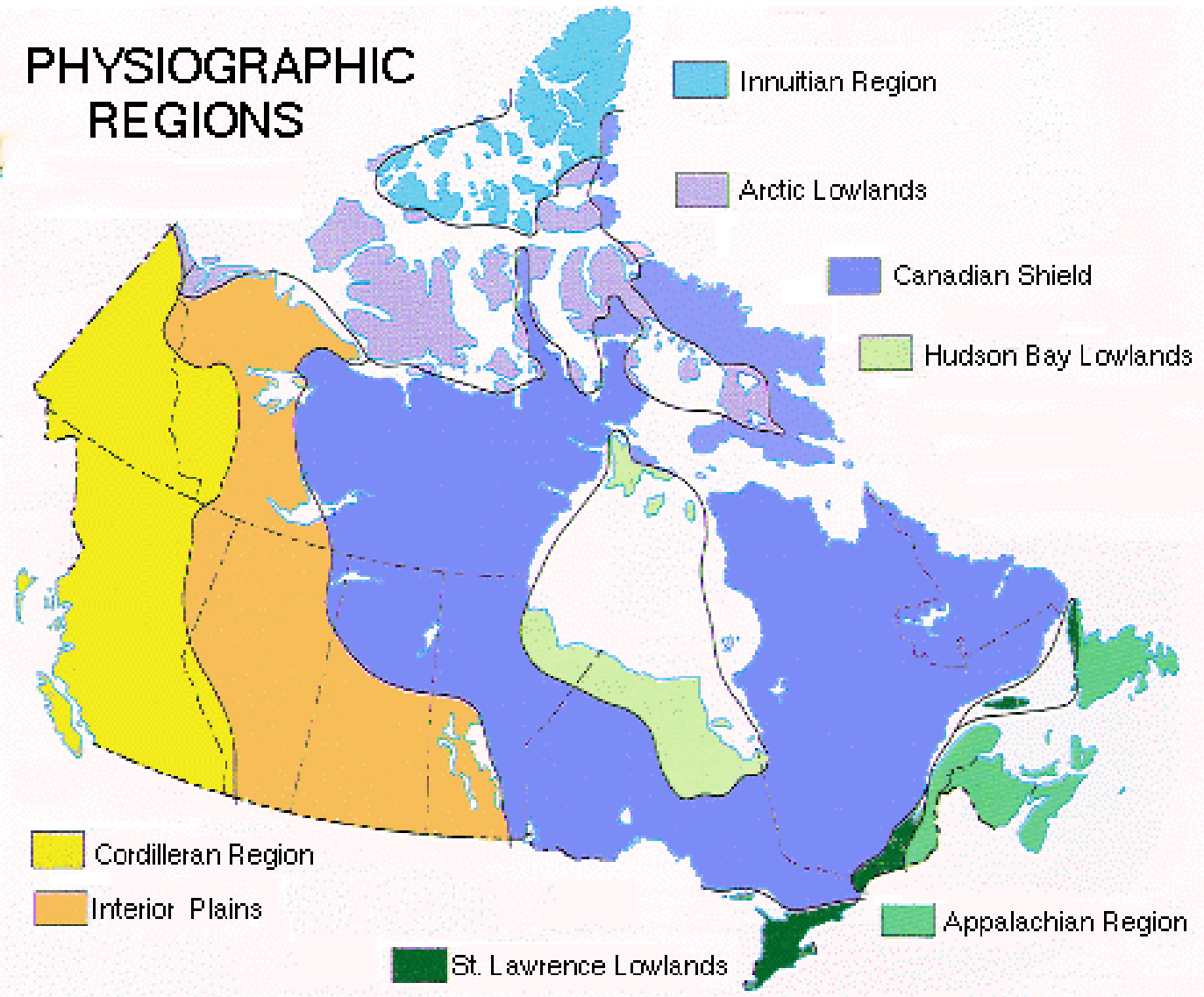
- why is there extension in the Basin and Range in the first place?
- extension started about 25 Ma - same time as the beginning of subduction of the East Pacific Rise
- when slab drops off, mantle oozes around it → huge positive gravity anomaly
- fast uplift of Colorado Plateau (since 5 Ma) as a result of thermal expansion

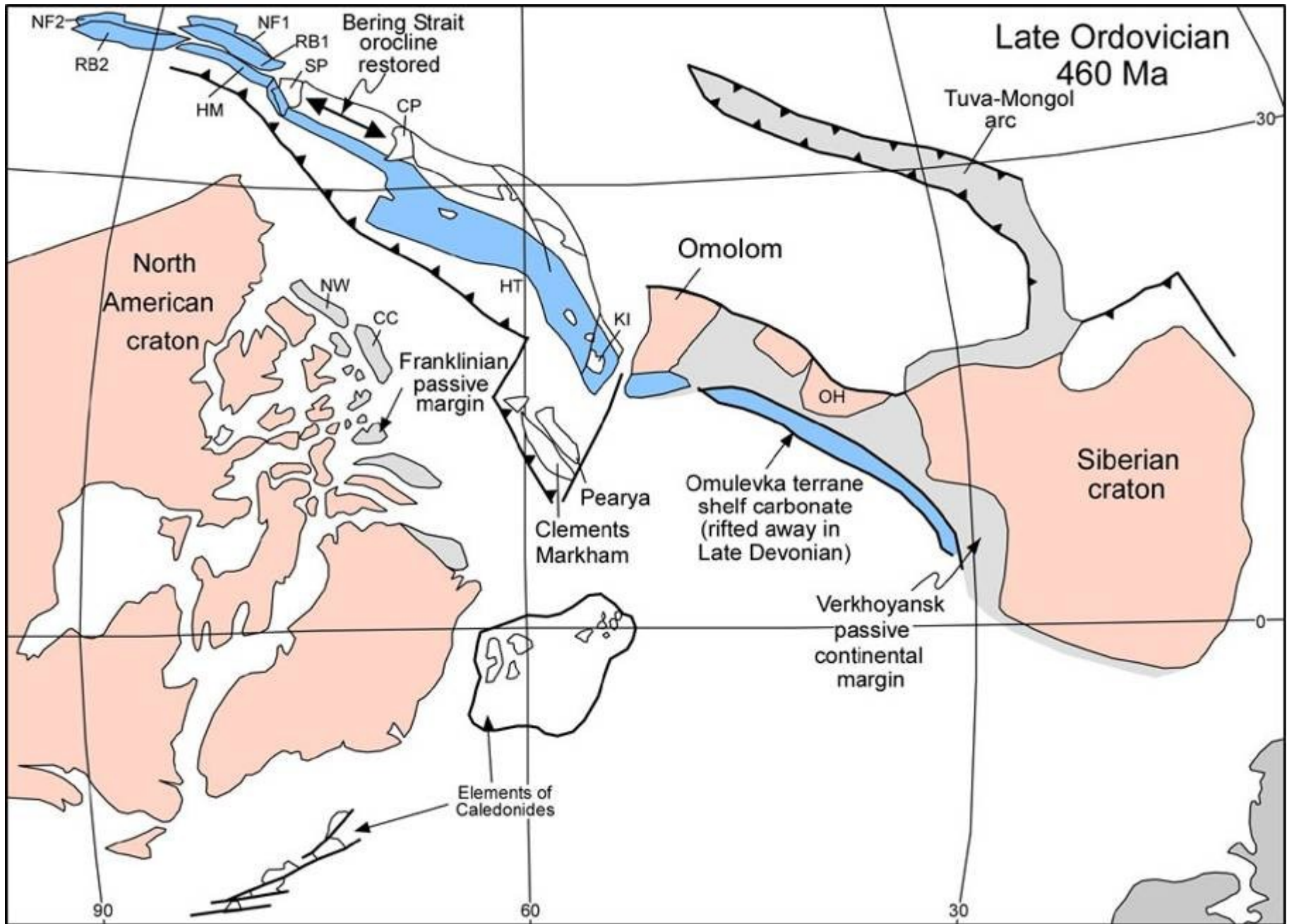


Line 745

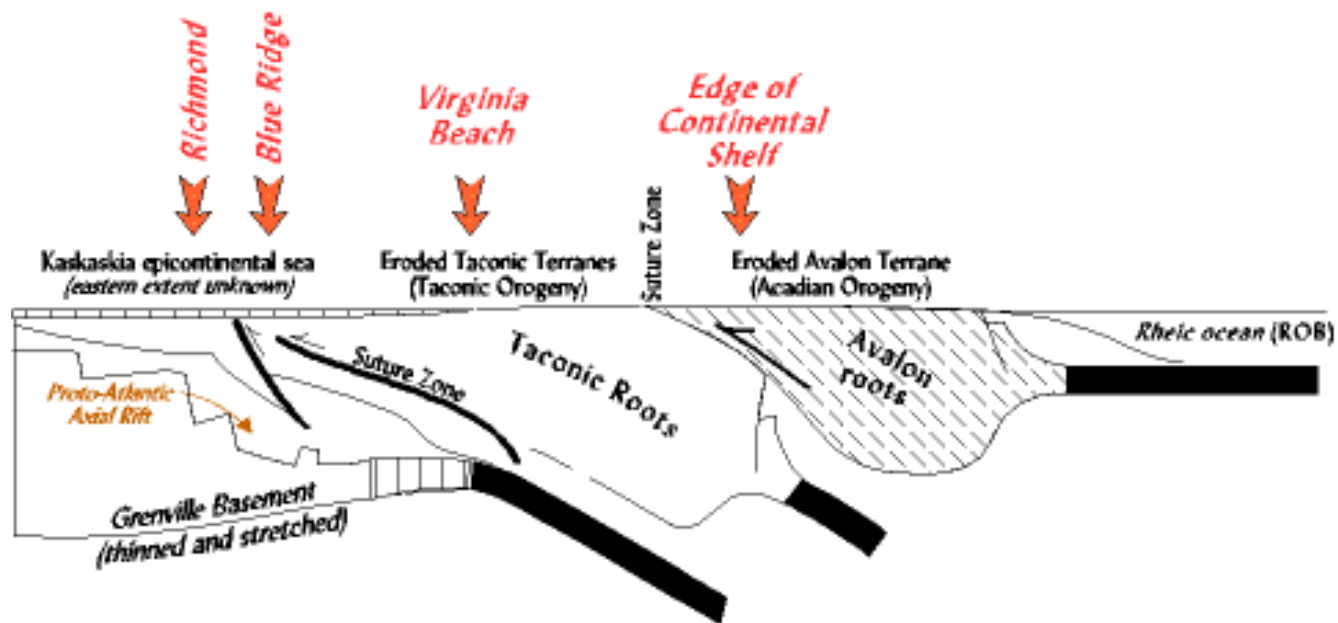


PHYSIOGRAPHIC REGIONS





*Approximate locations of modern geographic features
on the Pre-Alleghenian Mid-Atlantic Region*



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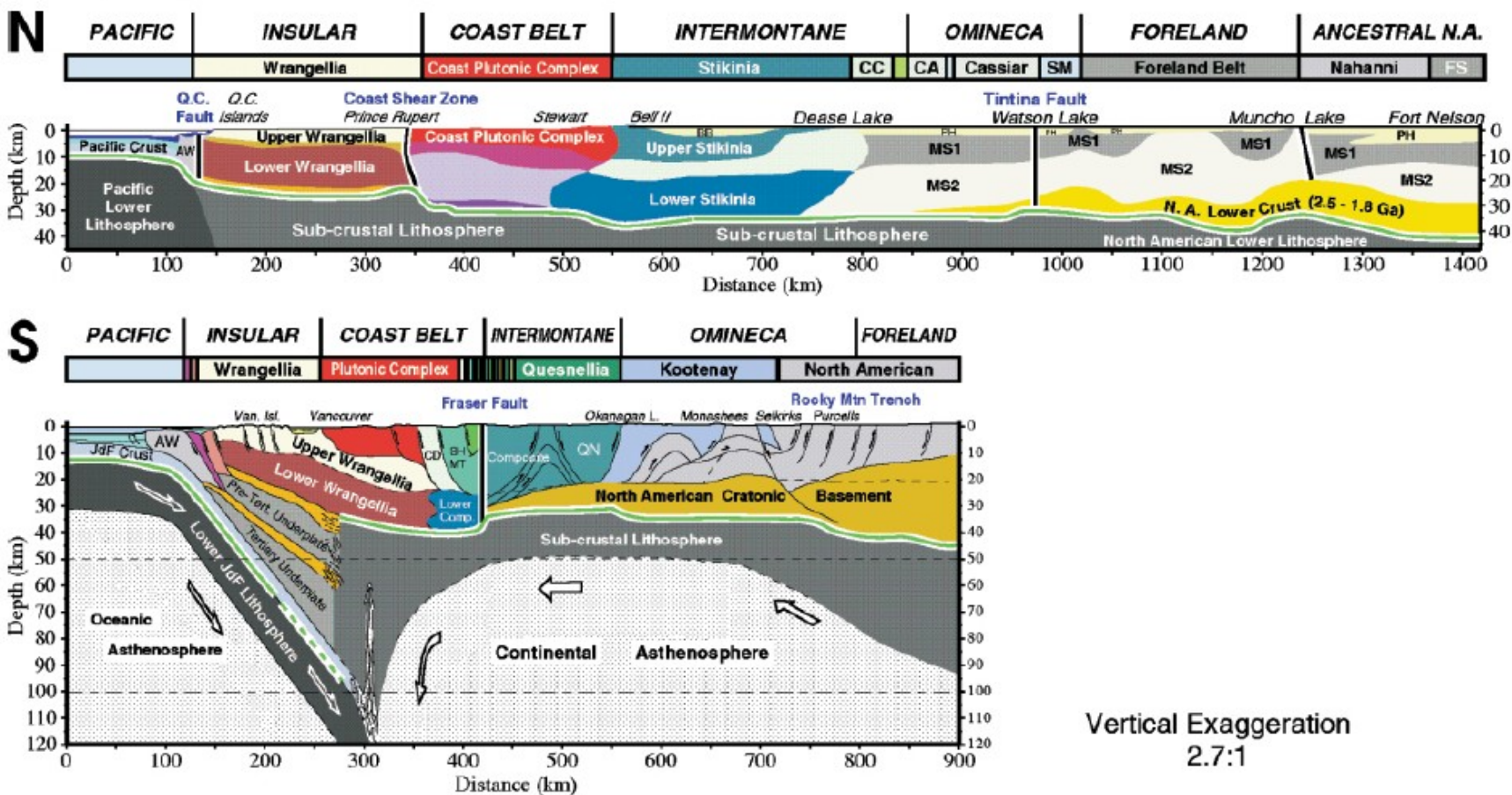


Figure 2. Comparison of interpreted simplified lithospheric structures along the northern (N) and southern (S) Cordilleran Lithoprobe transects, whose locations are shown by the red lines in Figure 1. In both profiles, the heavy green line is the crust-mantle boundary (Moho). AW accreted wedge; AX Alexander terrane; BB Bowser Basin; CA Cassiar terrane; CC Cache Creek terrane; CD Cadwallader terrane; FS, Fort Simpson (a Precambrian terrane in the craton); KO Kootenay terrane; MS1 undivided Precambrian (1200-800 Ma) metasedimentary rocks; MS2 undivided Precambrian (1800-1200 Ma) metasedimentary rocks; MT-SH undivided Methow and Siuksan terranes; QN Quesnel terrane; SM Slide Mountain terrane; ST Stikine terrane; WR Wrangellia. Most terrane descriptions are in Table 2. Figure modified by P.T.S. Hammer from Clowes and Hammer, 2000.

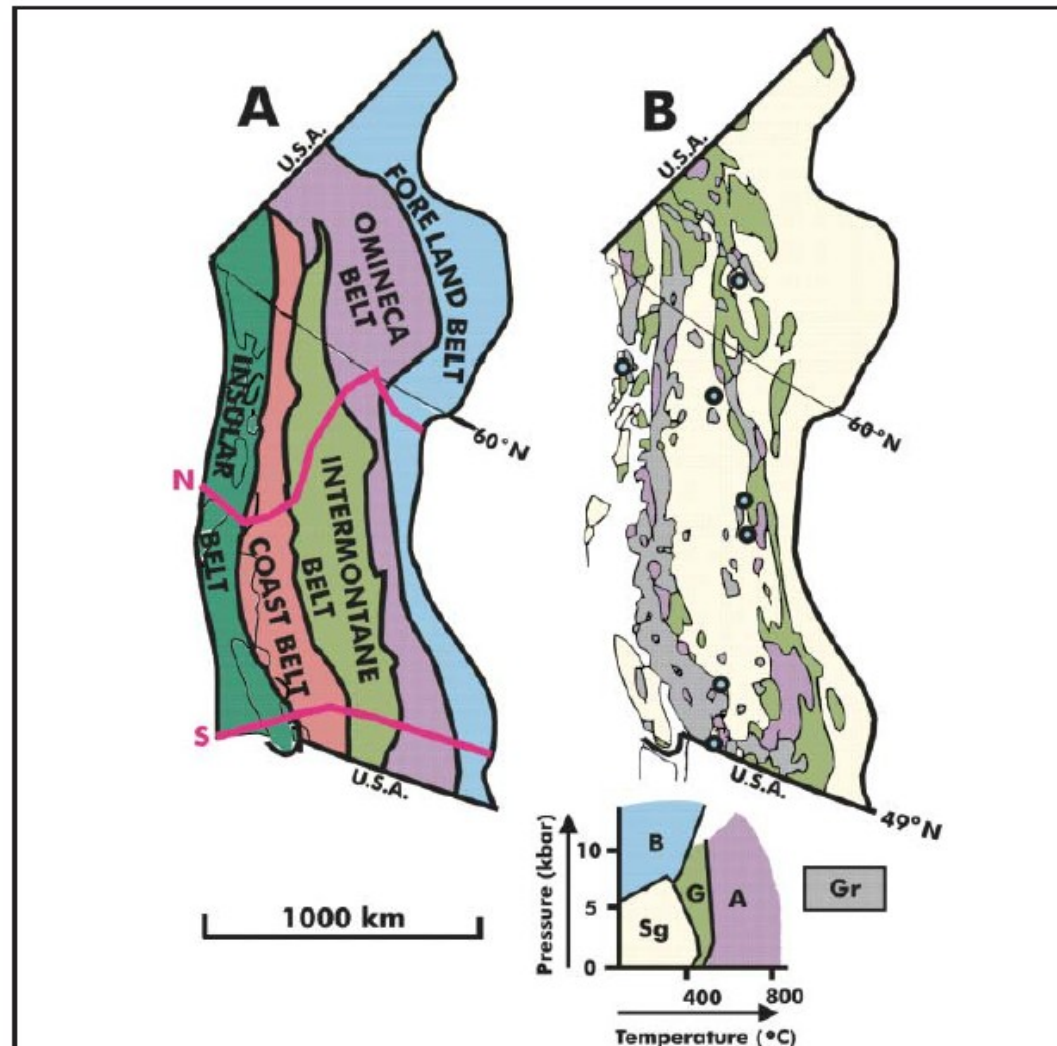


Figure 1. The Canadian Cordillera and adjoining parts of southeastern Alaska showing:
 A Location of the five morphogeological belts (details in Table 1); red lines show approximate positions of the northern (N) and southern (S) Canadian Cordilleran Lithoprobe transects, details of which are in Figure 2.
 B Simplified metamorphic map of the Canadian Cordillera, showing the close correspondence between the distribution of higher grade metamorphic rock facies and granitic rocks and Omineca and Coast belts. The map legend below is a pressure-temperature diagram whose colours correspond with those on the map; metamorphic facies are: Sg subgreenschist; G greenschist; A amphibolite, and B blueschist (blue dots on map); box labelled Gr denotes granitic rock.

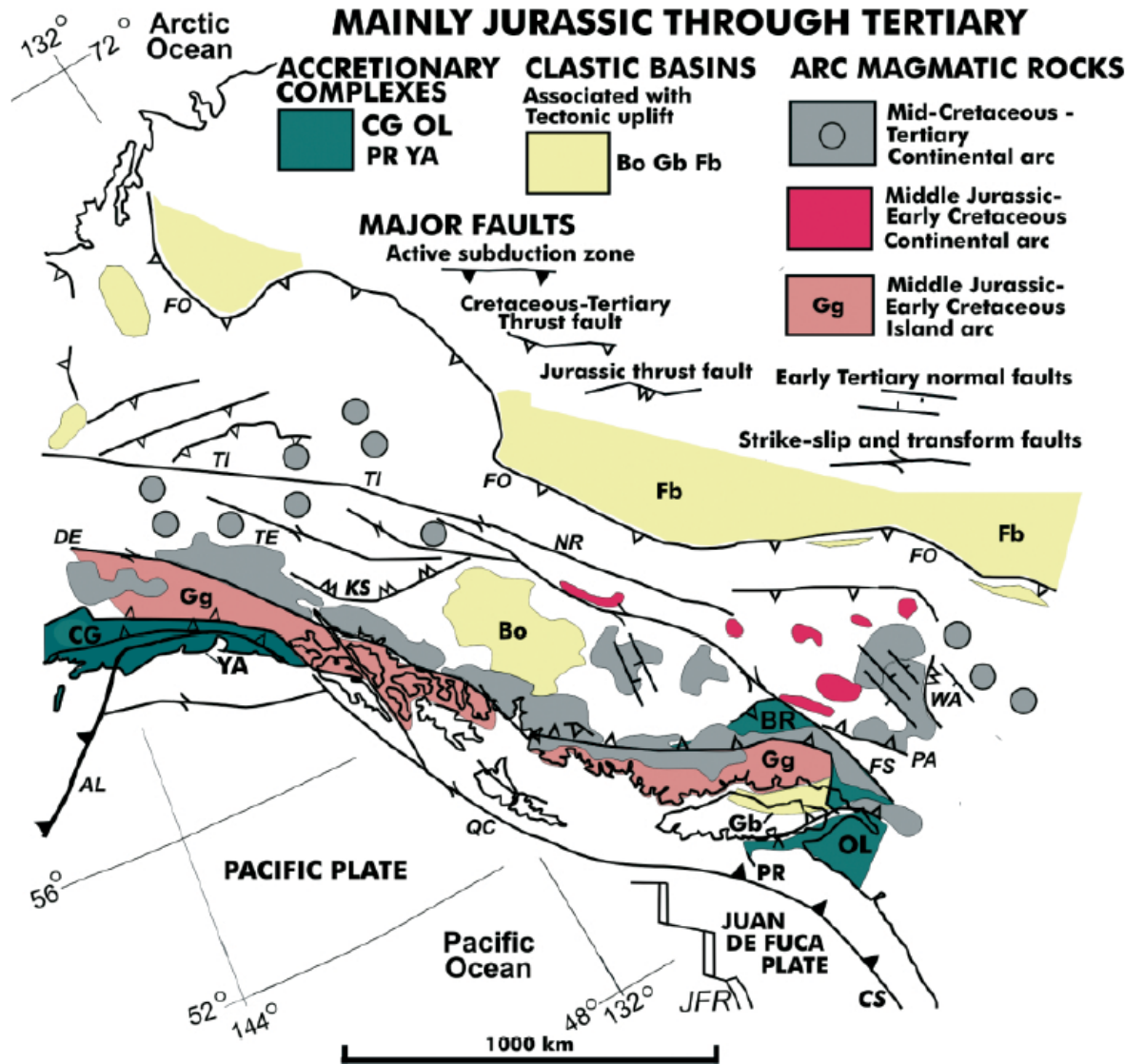


Figure 4. Locations of features of the Canadian Cordillera and adjacent parts of Alaska that formed during the period of terrane accretion and mountain building, mainly from Middle Jurassic through early Tertiary time (~180-40 Ma). (1) Middle Jurassic to Early Cretaceous continental arcs that are emplaced across CC, QN, KO; (2) Middle Jurassic through Early Cretaceous Gravina-Gambier (Gg) island arc, emplaced across WR, AX; (3) Mid-Cretaceous through early Tertiary continental arcs emplaced across all terranes, with exception of the accompanying accretionary complexes; filled circles denote plutons too small to show on the map; (4) clastic sedimentary basins (Bo Bowser Basin; Fb Foreland Basin; Gb Georgia Basin; not shown are basins on the continental shelf filled with material eroded from adjoining, uplifted fold and thrust belts. (5) Major faults include (5a) active subduction zones (AL Aleutian; CS Cascade); (5b) active transform fault (QS Queen Charlotte); (5c) major thrust fault systems of (a) Jurassic age (KS King Salmon; WA Waneta) and (b) mid-Cretaceous and early Tertiary ages (FO Foreland; PA Pasayten); (5d) major strike-slip faults of mainly Late Cretaceous and Tertiary ages (DE Denali; NR Northern Rocky Mountain Trench; TE Teslin; TI Tintina); (5e) early Tertiary normal faults.

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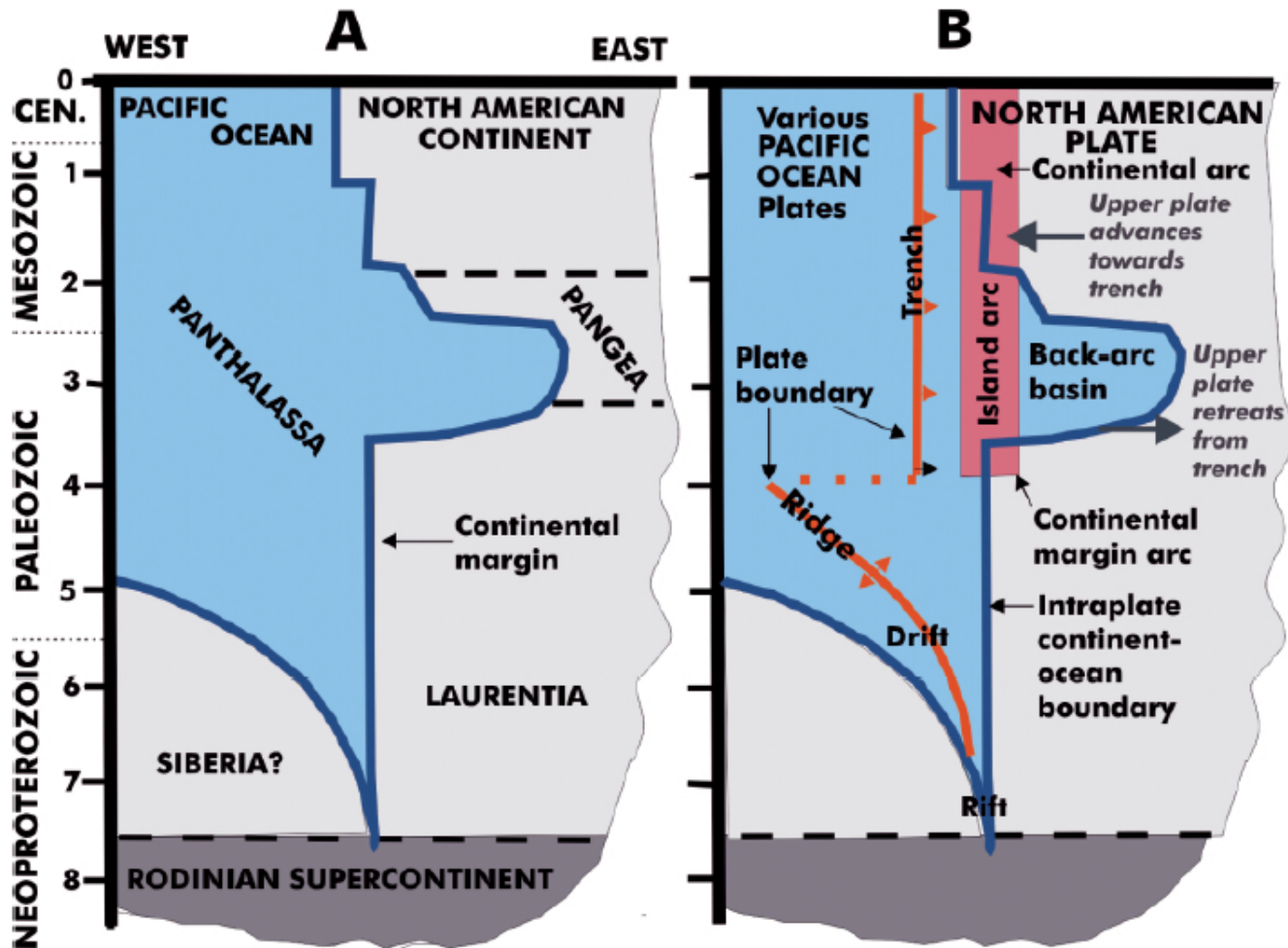


Figure 5: Evolution of the Canadian Cordillera cartooned on space-time diagrams. Horizontal coordinate: west to east (in present geographic coordinates); geographic position (after Early Devonian time) is fixed relative to the trench; vertical coordinate: numbers show age in hundreds of millions of years before present. A gives names of continents and ocean at different times and the names applied to each at those times. B shows features related to the plate tectonic activity at different times.

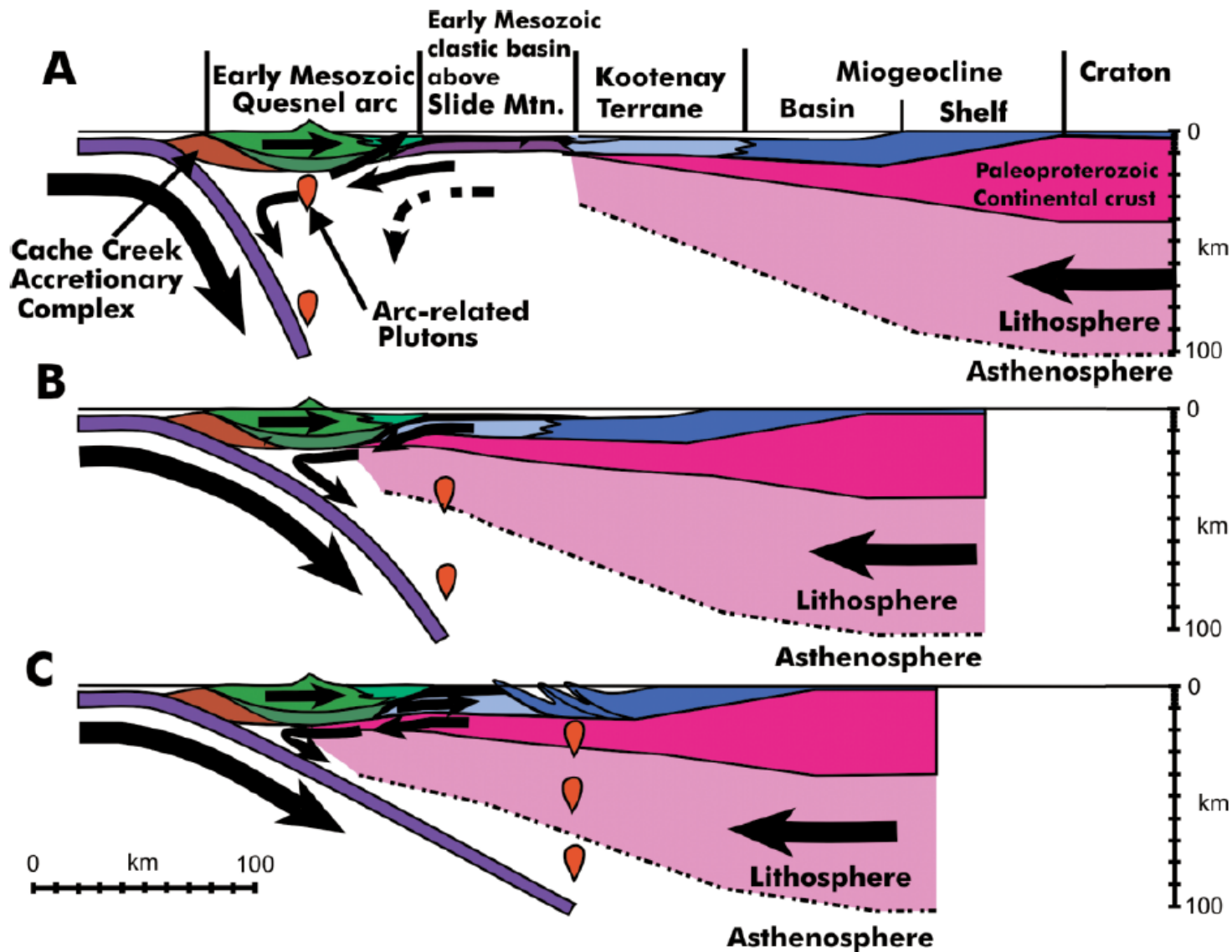


Figure 6: Tectonic wedging and crustal thickening in southeastern British Columbia.

A: Early Jurassic (~185 Ma) island arc (of QN, Figure 3) and its early Mesozoic back-arc basin, (mainly) on top of Slide Mountain terrane; onset of convergence of North America with trench to west and collapse of back-arc basin. B: Late Early Jurassic (~180 Ma) collapsed basin thrust over old continental margin; flattening of subduction zone and initiation of continental margin arc. C: Early Middle Jurassic (~170 Ma) southwest verging deformation occurred as Kootenay terrane was detached from North America and wedged under the old continental margin deposits; North American lithosphere wedged under Quesnel terrane; entrained and consumed in the subduction zone; subduction zone flattened and magmatic arc migrated eastward into the zone of southwest verging deformation.

