Short course "Diamond as a messenger from the Earth's interior: natural samples and experiment" Part 2:

What do we learn from diamonds?

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Outline

- What do we know about diamond formation?
 - Phenocrysts or xenocrysts?
 - Complex growth resorption patterns in natural diamonds
 - Role of oxidation reduction processes
 - Role of carbon saturation in mantle fluids/melts
- Applications
 - Carbon isotopes carbon cycle
 - Craton-formation
 - "window" into the mantle
 - How studies of diamond inclusions help in kimberlite prospecting and exploration

Phenocrysts or xenocrysts?





Taylor et al. (2000)

FIG. 1. HRXCT three-dimensional image of the orligite sensitish US1, created by stacking the 80 two-dimensional

Diamond nucleation and growth by reduction of carbonate melts under high-pressure and high-temperature conditions

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Advanced Materials Laboratory, National Institute for Materials Science 1-1, Namiki, Tsukuba, Ibaraki 305-0044, Japan ABSTRACT

We report for the first time experimental evidence for the nucleation and growth of diamonds from carbonatitic melts by reduction in reactions with silicon metal or silicon carbide. Experiments were carried out in the CaMg(CO₃)₂-Si and CaMg(CO₃)₂-SiC systems at 7.7 GPa and temperatures of 1500–1800 °C. No graphite was added to the run powder as a carbon source; the carbonate-bearing melts supply the carbon for diamond formation. Diamond grows spontaneously from the carbonatitic melt by reducing reactions: CaMg(CO₃)₂ + 2Si = CaMgSi₂O₆ + 2C in the CaMg(CO₃)₂-Si system, and CaMg(CO₃)₂ + 2SiC = CaMgSi₂O₆ + 4C in the CaMg(CO₃)₂-Si system. Our results provide strong experimental support for the view that some natural diamonds crystallized from carbonatitic melts by metasomatic reducing reactions with mantle solid phases.





Phenocrysts or xenocrysts?



Name of kimberlite Emplacement age (Ma) P-type arzburgitic (Ga) P-type Iherzolitic (Ga) FD References E-type (Ga) E-type (Ga) \Diamond Premier 1180 ± 30 ~2.0 1,2,3 ~2.0 ~1.2 519 ~2.0 ~2.0 Venetia 3.4 235 ± 2 ~2.9 ~15 56 Iwaneng 155 \Diamond ~2.6 Klipspringer Finsch 118 ± 3 ~3.3-3.2 1.58 ± 0.05 8, 9, 10 \Diamond Orapa 93.1 ~2.9 0.99 ± 0.05 ⋇ 10, 11 Kimberley pool 95 ~3.3-3.2 2.89 ± 0.06 8, 12 90.4 ~2.9 ~1.1 13 Koffiefontein \Diamond ~1.7 ~1.1 Jagersfontein 86 14 TABLE 2. Ki rlite Ages and Diamond Ages from Slave Province Kimberlites and Diamond Mi Emplacement P-type arzburgitic (Ga) Name of P-type Iherzolitic kimberlite age (Ma) E-type (Ga) FD References Anuri 613 Gahcho Kué 542 2 \diamond ⋇ Snap Lake* 533-535 Victoria Island 256-286 2 \Diamond Jericho 172.3 4 ** Diavik* 55 -35-33 2.2 - 1.85.6 \diamond 35 ± 0.17 7,8 Panda* 53 Kimberlite Ages and Diamond Ages from Kimberlites of the Siberian Crator TAI Name of Emplacement P-type harzburgitic (Ga) P-type Iherzolitic (Ga) kimberlite age (Ma) E-type (Ga) FD References Chomur) 436-421 (Upper Olenek Nakyn 364 4 Udachnaya 361 ± 6 ~3.5 - 3.1 2.9 ± 0.4 $\sim 2.01 \pm 0.06$ ⋇ 1,2, 3 (Daldyn) * Yubileynaya (Alaki 358 ☀ Mii 360 (Malo-Botuoba) 23 Party Congress (Malo-Botuoba) Upper Muna 345 Kharamai 235 Kuoika 128 - 148npilation by Griffin et al. (1999); Inclusion ages: 1 = Pearson et al. (1999), 2 = Parson et al. (1995), 3 = Richard-References: For mberlite ages se son and Harris (19 4

From Gurney et al. (2010)

Archean

Southern Afr

Diamond Min

Proterozoic

Kimberlite Ages and Diamond Age

From Shirey (2013) and Gurney et al. (2010)

Complex growth - resorption patterns in natural diamonds Smart et al. (20

121.6 -4.68 (f) K95-3-4 (e) L97-A2-5 (d) K95-5-1 663.5 .4.50 (i) K95-3-3





Role of carbon saturation in mantle fluids/melts

Diamond may form in Earth's mantle by a variety of processes (Stachel and Luth, 2015):

- recrystallization of the low-pressure graphite polymorph,
- Precipitation from a fluid or melt saturated with carbon,
- by oxidation-reduction reactions involving carbonate or methane.



Oxidation - reduction processes



Carbon saturation in mantle fluids/melts





- Xenoliths P-T below solidus → Diamond growth from fluid
- extremely limited redox buffering capacity of cratonic peridotites → redox reactions cannot produce notable diamond growth

What can we learn from diamonds?

- Carbon cycle
- Formation of cratons
- "Window" into the mantle
- Kimberlite exploration

Carbon isotopes - carbon cycle



Worldwide carbon isotopic composition of diamonds ranges from -41 to +5‰, close to the range in sedimentary rocks.

~72% of diamonds have carbon isotopic composition within of -8 to -2‰ (mean -5‰) = similar to mantlederived rocks (mid-ocean ridge basalts, ocean island basalts, carbonatites, kimberlites).

From Shirey et al. (2013)



Underthrusting of oceanic slabs or upwelling plume magmatism?





Formation of cratons

Mantle residence temperatures for peridotitic, websteritic and eclogitic diamonds based on the nitrogen thermometer shows that diamonds forming in very distinct environments (peridotite vs. eclogite) show identical equilibration temperatures.

-40

-30

Origin of cratonic lithosphere during Archean subduction events



When plate tectonics has started? Insights from diamond ages



Diamonds as a "time-capsule" of ancient processes



From Stachel and Harris (2008) and Helmstaedt and Schulze (1989)

When plate tectonics has started? Insights from diamond ages



Shirey and Shigley (2013)

- Significant difference in age between E- and P-type diamonds: no eclogitic diamonds (Etype) older than 3 billion years
- E-type diamonds (3 Ga) capture the first record of basaltic rock (eclogite is a basalt at high pressure metamorphism) in the mantle keel of the continents.
- This indicates ocean basin closure and continental collision (modern plate tectonics or Wilson Cycle), because
- Basalt is derived from the ocean floor and is incorporated into the mantle keel during collision.
- Mark a transition from a planet dominated by vertical geodynamic processes (plumes) to lateral tectonics and subduction.

When plate tectonics has started? Insights from nitrogen and carbon isotopes





- Witwatersrand diamonds have enriched nitrogen but mantle carbon isotopic compositions.
- This nitrogen values suggest contamination of the mantle by nitrogen-rich Archaean sediments.
- Modern-style plate tectonics operated as early as 3.5 billion years ago

"Window" into the mantle



Composition of fluids in diamonds



Kimberlite prospecting and exploration



- Pyrope Garnet Chromite •
- Ilmenite
- Chrome-diopside (Olivine)
- Heavy minerals, resistant to chemical and • mechanical weathering





McClenaghan & Kjarsgaard, 2007

Diamond grade of a kimberlite depends on:

- How much diamond-bearing peridotite and eclogite are present
- What the diamond grade of the source rocks were
- How well the diamonds were preserved during transportation to the surface

Indicator minerals:

- Recognition of mineral composition known to be associated with diamond
- Confirmation that these are derived from the diamond stability field ("diamond window")
- Assessment of the quantity of highinterest mantle minerals sampled and preserved by kimberlite

Mineral – composition

Mineral abundance

Recognition of mineral composition known to be associated with diamond





Mineral Services

Recognition of mineral composition known to be associated with diamond

Ilmenite



MgO (white)

Diamond Inclusions



Garnets The mantle source :

- Rich in carbon?
- In diamond stability field ("diamond window")?
- 1. Peridotite association vs. Eclogite association
- Peridotite xenoliths:
 Grt harzburgite > Spl harzburgite > Grt lherzolite
- 1. About 32-fold preferential association of carbon for depleted harzburgite over lherzolite



From Stachel and Harris (2008)

Garnets: "diamond window"



Thermobarometery

- Garnet
- Chrome Diopside
- Enstatite

Ni in Garnet thermometer:

- -well-calibrated, reliable
- single grain thermobarometer





Diamond preservation in kimberlite



After Mitchell (1986)

